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Ensemble Filter on dynamically driven RKHS: Application to a Multilayer Quasi-Geostrophic ocean model

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Key Points:

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10	•	An ensemble filter defined on a time evolving RKHS family driven by the system
11		dynamics is presented.
12	•	The corresponding filter enables extending classical Bayesian ensemble filters to effi-
13		ciently solve a 4D Var assimilation problem.
14	•	Assessed on a multilayer quasi-geostrophic simulation for the North Atlantic Ocean
15		basin in an OSSE framework with SSH observations.

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16 Abstract

Estimating the ocean's dynamic state, particularly at the mesoscale, is challenging due to the non-linear nature of ocean processes. Traditional data assimilation methods like optimal interpolation (OI) and four-dimensional variational assimilation (4D-Var) improve ocean state estimates but face trade-offs: OI is computationally efficient but struggles with small-scale features, while 4D-Var is more accurate but computationally demanding. Ensemble methods offer a compromise, with ensemble Kalman filters assuming Gaussianity and particle filters limited by high-dimensional systems.

This article introduces a novel ensemble filter using a Reproducing Kernel Hilbert 24 Spaces (RKHS) framework, embedding observables in time-evolving RKHSs to relax Gaus-25 signity assumptions and provide computational efficiency comparable to 3D-Var while ad-26 dressing 4D-Var problems. A key innovation is a tiled assimilation approach that enhances 27 the ensemble's physical properties over time. Applied to a quasi-geostrophic model of the 28 North Atlantic with synthetic sea surface height (SSH) data, this RKHS-based filter accu-29 rately assimilates observations with non-Gaussian noise, improving both the accuracy and 30 efficiency of ocean state estimations in complex systems. 31

³² Plain Language Summary

Estimating the ocean's dynamic state, such as eddies and currents, is challenging due to the complexity of ocean processes. This paper introduces a new data assimilation method, grounded in a rigorous theoretical framework, that uses an ensemble of simulations to improve these estimates. It builds new trajectories by combining ensemble members in a time-invariant way, enabling efficient assimilation of observations across space and time.

By updating the ensemble at the start of each observation period and simulating forward, the method enhances physical consistency. Tests with a North Atlantic model and synthetic satellite data demonstrate its ability to accurately estimate the ocean's state, even with noisy, non-Gaussian observations, while improving both accuracy and efficiency compared to traditional methods.

43 1 Introduction

Data assimilation methods in ocean modeling aim to integrate sparse observations, 44 both in time and space, of the ocean surface into numerical circulation models to produce 45 accurate estimates of the ocean's state. In this context, a major challenge is the assimi-46 lation of sea surface height (SSH) observations, collected over time from orbiting satellite 47 swaths, to accurately represent ocean mesoscale dynamics. The mesoscale, characterized by 48 features such as eddies, fronts, and jets, is crucial for understanding and modeling global 49 ocean circulation. To address this complex problem, several methods proposed by the data 50 assimilation community (Carrassi et al., 2018; Evensen, 2009; Reich & Cotter, 2015) have 51 been adapted and are now routinely used for various applications. 52

Optimal interpolation (OI) is a geostatistical method used to reconstruct a dense 53 field of relevant oceanic variables from spatially sparse satellite observations by considering 54 predetermined statistical properties (Le Traon et al., 2003). This method integrates ob-55 servations collected at various times, assuming spatiotemporal correlation function between 56 the observations and the estimated field. It is in particular employed to construct the Data 57 Unification and Altimeter Combination System (DUACS) products (Taburet et al., 2019). 58 OI's reliance on statistical interpolation to merge observations from varying times often re-59 sults in an over-smooth degraded representation of small-scale structures. This limitation 60 is particularly evident in the reconstruction of mesoscale features, where finer details are 61 crucial for accurate modeling. Ubelmann et al. (2015) observed that optimal interpolation 62 often misrepresents the transport over time of small-scale eddies. To address this issue, they 63

introduced dynamic interpolation, which offers a more accurate reconstruction of small-scale 64 structure transport between two observations. Assuming an invertible dynamical system, 65 the method involves running the model forward and backward between two observation 66 times, and then reconstructing the sea surface height as the mean of the results from both 67 simulations. Although this method better represents transport in the reconstructions, it 68 still requires two space interpolations to obtain the complete fields needed to initialize the 69 forward and backward simulations. Consequently, the effective resolution remains relatively 70 low. 71

72 Variational methods aim to minimize a cost function that penalizes the distance between the observations and the estimated state. Among these, four-dimensional variational 73 assimilation (4D-Var), assimilates spatially and temporally distributed observations. This 74 method uses optimizations tools, such as adjoint methods (Le Dimet & Talagrand, 1986), 75 which efficiently compute the gradient of the cost function with respect to its parameters by 76 running the dynamics forward and then the tangent linear adjoint system backward. Despite 77 the effectiveness of these methods, the associated computational costs can be prohibitively 78 high, hardly parallelizable, and constructing a numerically exact tangent linear operator 79 can be a challenging task. To address this, neural networks such as 4DVarNet (Fablet et 80 al., 2021) can be use to efficiently solve it after a learning phase. Recently, 4DVarNet has 81 been applied to the reconstruction of the sea surface currents from sea surface height (SSH) 82 and sea surface temperature (SST) observations (Fablet et al., 2024), demonstrating recon-83 structions with a higher effective resolution than DUACS. Machine learning methods offer 84 powerful tools for efficient data assimilation (Cheng et al., 2023). These approaches can be 85 integrated with classical assimilation methods to improve future state predictions (Brajard 86 et al., 2020; Farchi et al., 2021). It can also be used to interpolate partial observations to 87 retrieve the ocean mesoscale dynamic state (Ouala et al., 2018; Manucharyan et al., 2021). 88 In these methods, neural networks are employed to model the system's dynamics, thereby 89 bypassing the need to solve the dynamical system equations directly. However, this class 90 of approach opens the question how to assess, without the help of a dynamical model, the 91 physical quality of the generated predictions. In particular, neural networks introduce gen-92 eralization challenges, which may lead to poor performance when the model encounters new 93 or previously unseen physical phenomena. 94

Another strategy, called *nudging*, introduces a correction term into the dynamics equa-95 tions to fit better with the observations (Hoke & Anthes, 1976). Back-and-forth nudging 96 (BFN) (Auroux & Blum, 2008), involves iteratively solving the forced dynamical equation 97 forward and backward in time until the algorithm converges toward a trajectory. This 98 method circumvents the need for minimization tools used in 4D-Var methods, but it re-99 quires uncontrolled modifications of the dynamical system, and hence of the physical prior 100 we have on the system's dynamics. Le Guillou, Metref, et al. (2021) apply BFN to data from 101 the Surface Water Ocean Topography (SWOT) mission to assimilate a quasi-geostrophic 102 model, demonstrating the potential of SWOT data to enhance BFN performance. It can be 103 remarked, that BFN incurs computational costs due to the necessity of running the model 104 multiple times until convergence. This makes it impractical as it is for application to com-105 plex models. To address this limitation, Le Guillou, Lahaye, et al. (2021) proposed a hybrid 106 approach that combines BFN with 4D-Var. In this approach, BFN is used to assimilate bal-107 anced motions in a quasi-geostrophic model, while 4D-Var is employed to assimilate internal 108 tides in a shallow water model. 109

To manage computational complexity in high-dimensional ocean models, one effective approach is the use of ensemble Kalman filter (EnKF) methods, which operate locally in the phase space (Evensen, 2009; Carrassi et al., 2018), in the sense that the solution is assumed to be well characterized by the span of an ensemble of realizations of the dynamical model relatively close from each other. The use of Kalman filters in nonlinear systems is justified by their ability to provide the best linear unbiased estimator within a Gaussian framework. EnKFs assimilate observations sequentially as they become available, assuming

both the state estimation error and observation error to be Gaussian. They approximate 117 the state estimate and its uncertainty using an ensemble of realizations, where the analyzed 118 ensemble members are constructed as linear combinations of the forecast ensemble. However, 119 this superposition principle lacks theoretical justification for highly nonlinear systems and 120 it has been shown that EnKF do not converge toward the true filtering distribution (Le 121 Gland et al., 2011). For a general review on the asymptotic results of the EnKF in a 122 Gaussian context, interested readers may refer to (A. N. Bishop & Del Moral, 2023). Despite 123 this limitation, EnKFs have been successfully applied in various oceanographic studies, 124 demonstrating their efficiency in estimating the ocean state by assimilating SST or SSH data 125 (Leeuwenburgh et al., 2005; Wan et al., 2010). A significant limitation of classical EnKFs 126 is their sequential nature, which leads to correct the model at the final observation time of 127 a given time window (and not at the window's initial time). To address this, smoothing 128 methods based on forward-backward assimilation of trajectories from data collected over the 129 whole time window have been developed. Among these, ensemble extension of the 4DVar 130 techniques (4DEnVar) have been proposed (Lorenc, 2003; Liu et al., 2008; Desroziers et al., 131 2014). Such techniques replace the exact numerical tangent linear operator by an ensemble 132 approximation, leading to much cheaper and practical methods (Yang et al., 2015). Other 133 methods use time-constant weights to reconstruct trajectories, and formulate a cost function 134 of this weight vector (Hunt et al., 2004; Bocquet & Sakov, 2014). Connections between these 135 methods and 4DEnVar can be established. However, the assumption of constant weights 136 over time lacks clear theoretical justification in a classical setting. 137

Alternatively, particle filters (Doucet et al., 2001; Gordon et al., 1993; Van Leeuwen et 138 al., 2019) relax the Gaussianity and linearity assumptions of Kalman filters by representing 139 the state distribution with an ensemble of particles, each associated with a weight. Unlike 140 EnKFs, particle filters do not alter the ensemble's physical properties by creating members 141 as linear combinations of existing members. However, particle filters are subject to the 142 curse of dimensionality, meaning that in high-dimensional state spaces, a huge number of 143 particles is required to accurately represent the state distribution. This makes particle filters 144 computationally expensive and challenging to apply to high-dimensional systems. 145

Reproducing Kernel Hilbert Spaces (RKHS) have been employed in data assimilation 146 methods to address system non-linearity and to relax the inherent assumptions of Kalman 147 filters. Mauran et al. (2023) propose a reformulation of the Ensemble Transform Kalman 148 Filter (ETKF) (C. H. Bishop et al., 2001; Hunt et al., 2007), operating within a RKHS 149 framework that uses feature maps (i.e. the mapping from the phase space onto the RKHS) 150 defined from the ensemble state space and observations. Gottwald and Reich (2021) intro-151 duced an ensemble Kalman filter designed to simultaneously retrieve the state variable and 152 learn the dynamical propagator, using random feature map bases. Both approaches aim 153 to determine weight vectors whose size scales with the state space, making them compu-154 tationally expensive for high-dimensional systems. This issue of increasing computational 155 cost with system dimensionality is also observed in other kernel-based assimilation methods 156 (Luo, 2019; Zhang et al., 2022). 157

RKHS-based assimilation methods share similarities with the decomposition method introduced by Ubelmann et al. (2021), where the state variable of the dynamical system is decomposed into a reduced space, and optimal interpolation is performed directly on the reduced state vector to define the analyzed state. In RKHS approaches, the reduced state vector corresponds to the weight vector in the RKHS, and the ensemble filter operates directly in this space.

The method proposed in this study relies on the theoretical framework developed in (Dufée et al., 2024), which differs from usual RKHS based estimation. In this setting the dynamical system is embedded within a family of time-varying reproducing kernel Hilbert spaces (RKHS). This framework enables establishing an evolution law for the RKHS bases across time. The RKHS are transported along time by the dynamical system, defining an isometry between the kernel evaluations over time. This approach enables the use of embar-

rassingly simple data assimilation schemes, reconstructing the trajectory as a time-invariant 170 linear combination of the ensemble members. These schemes rely on a fully justified super-171 position principle allowing to compare the trajectories at initial time, leading consequently 172 to effectively solve a 4D-Var problem with a time series of observations at the computational 173 cost of a three-dimensional variational (3D-Var) methods, in which time evolution is not 174 considered. Compared to the other RKHS schemes proposed in the literature our method 175 works on a small dimensional vector of function evaluations corresponding to the ensem-176 ble dimension. The small dimension of these evaluation vectors allows for direct matrix 177 inversion in the weight space making our approach applicable to high-dimensional systems. 178 Similar to particle filters, each ensemble member in the RKHS are assigned with (a signed) 179 weight, representing the state distribution. 180

This method has been applied to a multilayer quasi-geostrophic system in the North Atlantic basin, using synthetic observations generated from real-world satellite tracks, following an observing system simulation experiment (OSSE) methodology. We show that, within this framework, the proposed method leads to improved reconstruction resolution.

The paper is structured as follows: section 2 provides a brief overview of the RKHSs, the Koopman operator and how with the combination of these two tools a RKHS family transported by the dynamics is defined. Section 3 focuses on assimilation methods. Section 4 introduces the experimental framework. Finally, section 5 describes the results obtained.

¹⁸⁹ 2 Theoretical framework

This section introduces the RKHS theory, the Koopman operator, and the framework established by Dufée et al. (2024), which defines a family of RKHSs, built from the transport by the dynamical system of an initial kernel at the time instant t_0 .

2.1 RKHS

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A Reproducing Kernel Hilbert Space (RKHS), \mathcal{H} , is a Hilbert space of functions f: $E \to \mathbb{C}$, where E is a non-empty set, equipped with a inner product $\langle \cdot, \cdot \rangle_{\mathcal{H}}$. A map k: $E, E \to \mathbb{C}$ is defined as being a reproducing kernel of \mathcal{H} if it has the two following properties:

- membership property: $\forall x \in E, k(\cdot, x) \in \mathcal{H},$
- reproducing property: $\forall x \in E, \forall f \in \mathcal{H}, \langle f, k(\cdot, x) \rangle_{\mathcal{H}} = f(x).$

The reproducing property implies that the kernel satisfies $k(x,y) = \overline{k(y,x)}$, where $\overline{\bullet}$ denotes the complex conjugate. The RKHS is fully and uniquely determined by its kernel k(Aronszajn, 1950). The functions $k(\cdot, x)$ for $x \in E$ are usually referred to as feature maps. It can be shown that their span is dense in $(\mathcal{H}, \|\cdot\|_{\mathcal{H}})$. The integral operator associated with the kernel k, is defined as $\mathcal{L}_k : L^2(E, \nu) \to L^2(E, \nu)$, such that for any $f \in L^2(E, \nu)$,

$$\left(\mathcal{L}_{k}f\right)\left(x\right) = \int_{E} k\left(x,y\right)f\left(y\right)\nu\left(dy\right) , \qquad (1)$$

where L^2 denotes the space of square-integrable functions, and ν is a measure on E. The inner product in $L^2(E,\nu)$ for f and g in $L^2(E,\nu)$ is given by

$$\langle f,g\rangle_{L^2(E,\nu)} = \int_E f(y)\,\overline{g(y)}\nu\,(dy) \ . \tag{2}$$

Consequently, the inner product between a function $f \in L^{2}(E,\nu)$ and a kernel function $k(\cdot, x)$ becomes:

$$\langle f, k(\cdot, x) \rangle_{L^{2}(E,\nu)} = \int_{E} f(y) \overline{k(y,x)} \nu(dy) = \int_{E} k(x,y) f(y) \nu(dy) = (\mathcal{L}_{k}f)(x) .$$
(3)

The range of \mathcal{L}_k is assumed to be dense in $L^2(E,\nu)$, implying that its eigenvalues are strictly positive. By Mercer's theorem as k is a bounded, continuous, symmetric, positive semi-definite kernel, \mathcal{L}_k eigenfunctions $\{\varphi_i\}_{i\in\mathbb{N}}$ and eigenvalues $\{\mu_i\}_{i\in\mathbb{N}}$ define the RKHS \mathcal{H} as

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$$\mathcal{H} = \left\{ f \in L^2(E,\nu), f = \sum_{i=1}^{\infty} a_i \varphi_i \ s.t. \ \sum_i \frac{|a_i|^2}{\mu_i} < \infty \right\} . \tag{4}$$

This last expression highlights the fact that \mathcal{H} is a space of functions with fast decaying coefficients within $L^2(E,\nu)$.

218 2.2 Dynamical system

Let consider a dynamical system defined by a continuous flow Φ_t , operating on functions $\mathbf{X} \in L^2 \left(\mathbb{R}^+ \times \Omega_x, \mathbb{R}^d \right)$ included in an invariant phase space manifold, with Ω_x representing the spatial support. The state of the system at time $t \in \mathbb{R}^+$, denoted by $\mathbf{X}(t) = \mathbf{X}_t$ is given by $\mathbf{X}_t = \Phi_t (\mathbf{X}_0)$, where \mathbf{X}_0 is the initial condition. In differential form the flow is given by the system:

$$\begin{cases} \frac{\partial \mathbf{X}_{t}}{\partial t} = \mathcal{M}(\mathbf{X}_{t}) \\ \mathbf{X}(0) = \mathbf{X}_{0} \end{cases},$$
(5)

where X_t belongs to Ω a compact (invariant) subset of the space of square-integrable functions $L^2(\Omega_x, \mathbb{R}^d)$. The nonlinear operator \mathcal{M} is assumed to be differentiable with continuous derivatives. An observable of this system is defined as a square-integrable function that takes as input a function from Ω and outputs a real valued vector. In the following, we assume that the dynamical system is invertible, and that there exists a measure ν preserved by this dynamical system.

Remark 1 Let us point out that in our data assimilation setting, the invertibility condition
 will be required only on the assimilation time window. Existence of invariant measures for
 continuous system is ubiquitous.

The Koopman operator U_t is an infinite-dimensional linear operator mapping $L^2(\Omega, \nu)$ into itself. It propagates the observables over time, such that for every observable f and every \mathbf{X} in Ω , we have $U_t f(\mathbf{X}) = f(\Phi_t(\mathbf{X}))$.

2.3 Dynamical system embedded in a RKHS

With all the previous ingredients, we are now able to present the framework proposed by Dufée et al. (2024), on which our assimilation schemes are built.

Let Ω_0 denote the ensemble of initial conditions, and define $\Omega_t := \Phi_t(\Omega_0)$ as the ensemble of states (or snapshots) at time $t \in \mathbb{R}^+$. Since the system is invertible, for any snapshot $X_t \in \Omega_t$, there exists a unique initial condition $X_0 \in \Omega_0$ associated with it. Moreover in the present framework, we consider that for each state X_t , the current time t is known, in such a way it can be uniquely associated with its initial condition X_0 . Given an initial kernel $k_0 : \Omega_0 \times \Omega_0$ it is then possible to define a positive Hermitian kernel $k_t : \Omega_t \times \Omega_t$ such that,

$$\forall \mathbf{X}_t, \mathbf{Y}_t \in \Omega_t, \ k_t \left(\mathbf{X}_t, \mathbf{Y}_t \right) = k_0 \left(\Phi_t^{-1} \left(\mathbf{X}_t \right), \Phi_t^{-1} \left(\mathbf{Y}_t \right) \right) \ . \tag{6}$$

Let $(\mathcal{H}_0, \langle \cdot, \cdot \rangle_{\mathcal{H}_0})$ and $(\mathcal{H}_t, \langle \cdot, \cdot \rangle_{\mathcal{H}_t})$ represent the RKHSs associated with k_0 and k_t , respectively. A RKHS family $\mathcal{W} = (\mathcal{H}_t)_{t \geq 0}$ can be constructed by considering the RKHSs along time.

The operator $\mathcal{U}_t : \mathcal{H}_0 \to \mathcal{H}_t$ is defined as:

$$\forall \mathbf{X}_{0} \in \Omega_{0}, \ \mathcal{U}_{t}k_{0}\left(\cdot, \mathbf{X}_{0}\right) = k_{t}\left(\cdot, \Phi_{t}\left(\mathbf{X}_{0}\right)\right) \ . \tag{7}$$



Figure 1: Feature maps transported over time from t_0 to t_k by the dynamical system, with the observations collected between t_0 and t_k represented at the center of the figure in the form of satellite swath.

This operator propagates the features maps over time. It can be seen as a useful tool to define the kernel k_t at any time instant. Operator \mathcal{U}_t is closely related to the adjoint of the Koopman operator U_t , know as the Perron Frobenius operator. It follows from equation (6) that \mathcal{U}_t defines an isometry between \mathcal{H}_0 and \mathcal{H}_t .

The theory presented here operates on infinite ensembles of functions, given that both Ω and \mathcal{H} are infinite dimensional spaces. However, in practice, we are constrained to consider only a finite number of trajectories, denoted as $\{X^{(i)}, i = 1, ..., p\}$. Figure 1 illustrates the transported feature maps associated with the ensemble members, which will be employed to solve the data assimilation problem in section 3.2.

In ensemble data assimilation methods, a superposition principle is most of the time assumed, which in theory is not applicable for non-linear dynamical systems. However, this principle can be rigorously justified within the RKHS family described previously (Dufée et al., 2024). Considering an observable $g_x : \Omega_t \to \mathbb{R}^d$ such that for $t \ge 0$, $\forall \mathbf{X}_t \in \Omega_t, \forall x \in \Omega_x, g_x(\mathbf{X}_t) = \mathbf{X}_t(x)$. Assuming g_x is smooth enough to belong to \mathcal{H}_t , the kernel reproducing property allows the following decomposition:

$$g_x\left(\mathbf{X}_t\right) = \langle g_x\left(\cdot\right), k_t\left(\cdot, \mathbf{X}_t\right) \rangle_{\mathcal{H}_t} . \tag{8}$$

Since $k_t(\cdot, \boldsymbol{X}_t)$ is included in \mathcal{H}_t it admits a decomposition such that

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$$k_t(\cdot, \boldsymbol{X}_t) = \sum_{i=1}^p \boldsymbol{\beta}_i k_t\left(\cdot, \boldsymbol{X}_t^{(i)}\right) , \qquad (9)$$

where β is a vector of length p. At the continuous level, the sum involved in (9) is infinite, however with the finite number of ensemble members considered, it reduces to a finite sum approximation. By the inner product bilinearity and the kernel reproducing property it follows that:

$$g_x\left(\mathbf{X}_t\right) = \sum_{i=1}^p \beta_i g_x\left(\mathbf{X}_t^{(i)}\right) \,. \tag{10}$$

As this holds for every $x \in \Omega_x$, it can be concluded that:

$$\boldsymbol{X}_{t} = \sum_{i=1}^{p} \boldsymbol{\beta}_{i} \boldsymbol{X}_{t}^{(i)} , \qquad (11)$$

at condition that all the observables constituting the state vector – or at least the variables we aim at estimating – belong to \mathcal{H}_t . This superposition principle allows the construction of new trajectories as linear combinations of previously computed ones, with the linear combination being constant over time due to equation (6). As further described in the following sections this property leads to the setting of very appealing data assimilation methods.

Remark 2 From a machine learning perspective, the RKHS family can be viewed as latent spaces defined from the dynamics, where the weights in the latent spaces are equal to those in the physical space, and characterize a given trajectory.

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In the next section we provide a brief recap on ensemble data assimilation techniques.

²⁸⁸ 3 Assimilation methods

3.1 Classical methods

Although the previous setting is expressed in full generality within functional spaces, for the sake of simplicity, we will describe the assimilation techniques in the following sections using finite-dimensional spaces.

Assimilation methods aim to provide the best estimate of the state of a dynamical system from observations. The observations $Y \in \mathbb{R}^m$ are assumed to be related to the state of the system $X \in \mathbb{R}^n$, as follows,

$$\boldsymbol{Y} = H\left(\boldsymbol{X}\right) + \boldsymbol{\epsilon} \,, \tag{12}$$

where $H: \mathbb{R}^m \to \mathbb{R}^n$, with usually $n \gg m$, denotes the observation operator, $\epsilon \in \mathbb{R}^m$ 297 represent an additive centred Gaussian observation noise with covariance matrix $\boldsymbol{R} \in \mathbb{R}^{m \times m}$. 298 Classical assimilation methods aim to solve the ill-posed problem $H(\mathbf{X}) \simeq \mathbf{Y}$, seeking to 299 find the state that best fits the observations. For sake of simplicity, H will be assumed 300 to be linear in the following, and thus represented through matrix $\mathbb{H} \in \mathbb{R}^{m \times n}$ such that 301 $H(\mathbf{X}) = \mathbb{H}\mathbf{X}$. To have a well-posed problem and therefore a unique solution, one searches 302 for the closest solution to the forecasted state X^{f} . Consequently, assimilation methods 303 generally seek to minimize a cost function as follows 304

$$J(\mathbf{X}) = \frac{1}{2} \left\| \mathbb{H}\mathbf{X} - \mathbf{Y} \right\|_{\mathbf{R}^{-1}}^{2} + \frac{1}{2} \left\| \mathbf{X} - \mathbf{X}^{f} \right\|_{\mathbf{P}^{-1}}^{2}, \qquad (13)$$

where $\mathbf{P} \in \mathbb{R}^{n \times n}$ is the state variable covariance matrix, and as usual $\| \bullet \|_{\mathbf{A}}$ stands for the norm $(\bullet^T \mathbf{A} \bullet)$, with upperscript, T, denoting the transpose operator.

Cancellation of the cost function gradient gives the following expression of the optimal analyzed state

$$\boldsymbol{X}^{a} = -\left(\nabla^{2} J\left(\boldsymbol{X}\right)\right)^{-1} \left(\nabla J\left(0\right)\right) = \left(\mathbb{H}^{T} \boldsymbol{R}^{-1} \mathbb{H} + \boldsymbol{P}^{-1}\right)^{-1} \left(\mathbb{H}^{T} \boldsymbol{R}^{-1} \boldsymbol{Y} + \boldsymbol{P}^{-1} \boldsymbol{X}^{f}\right) .$$
(14)

³¹¹ Using the Sherman-Morrison-Woodbury formula, the previous equation can be rewritten,

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$$\boldsymbol{X}^{a} = \boldsymbol{X}^{f} - \boldsymbol{G} \left(\mathbb{H} \boldsymbol{X}^{f} - \boldsymbol{Y} \right) , \qquad (15)$$

³¹³ with the Kalman gain expressed as

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$$\boldsymbol{G} = \boldsymbol{P} \mathbb{H}^T \left(\boldsymbol{R} + \mathbb{H} \boldsymbol{P} \mathbb{H}^T \right)^{-1} \,. \tag{16}$$

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3.1.1 Ensemble Square root Filter (ESRF)

Ensemble methods consider p realizations (or members) in the state space, whose mean and covariance matrix represent the estimated state and covariance. Note that in highdimensional problems, such as those encountered in ocean dynamics, we have $p \ll m \ll n$, with typical values of $p \sim 50$, $m \sim 10^3$, $n \sim 10^6$. The ensemble matrix X is defined as a $n \times p$ matrix. Each columns i of X contains a member of the ensemble noted $X^{(i)}$, with $1 \le i \le p$. We briefly present in the following the ensemble square root filter (Whitaker & Hamill, 2002), which will serve as a reference in this study. Its name comes from the fact that the state covariance matrix is expressed through its square root, the anomaly matrix $A = [(X^{(1)} - \overline{X}), (X^{(2)} - \overline{X}), \dots, (X^{(p)} - \overline{X})]$, where \overline{X} denotes the ensemble mean, which represents the estimated state. The empirical covariance matrix is then expressed as:

$$\boldsymbol{P} = \frac{1}{p-1} \boldsymbol{A} \boldsymbol{A}^T \,. \tag{17}$$

With this decomposition, $A \in \mathbb{R}^{n \times p}$ is stored instead of $P \in \mathbb{R}^{n \times n}$.

The analyzed covariance matrix is expressed as the inverse of the Hessian of cost function J defined in equation (13). Using the Sherman-Morrison-Woodbury formula the covariance matrix update reads

$$\boldsymbol{P}^{a} = (\boldsymbol{I} - \boldsymbol{G} \mathbb{H}) \boldsymbol{P}^{f} , \qquad (18)$$

where G is the gain matrix expressed in equation (16). Exponents \bullet^f and \bullet^a stand for forecast and analyzed, respectively. Substituting the gain into the covariance matrix formula leads to,

$$\frac{1}{p-1}\boldsymbol{A}^{a}\left(\boldsymbol{A}^{a}\right)^{T} = \boldsymbol{A}^{f}\left(\boldsymbol{I}-\frac{1}{p-1}\left(\mathbb{H}\boldsymbol{A}^{f}\right)^{T}\left(\boldsymbol{R}+\mathbb{H}\frac{1}{p-1}\boldsymbol{A}^{f}\left(\boldsymbol{A}^{f}\right)^{T}\mathbb{H}^{T}\right)^{-1}\left(\mathbb{H}\boldsymbol{A}^{f}\right)\right)\frac{1}{p-1}\left(\boldsymbol{A}^{f}\right)^{T}.$$
(19)

Applying the Sherman–Morrison–Woodbury formula the previous equation can be written as

$$\boldsymbol{A}^{a}\left(\boldsymbol{A}^{a}\right)^{T} = \boldsymbol{A}^{f}\left(\boldsymbol{I} + \frac{1}{p-1}\left(\mathbb{H}\boldsymbol{A}^{f}\right)^{T}\boldsymbol{R}^{-1}\left(\mathbb{H}\boldsymbol{A}^{f}\right)\right)^{-1}\left(\boldsymbol{A}^{f}\right)^{T}.$$
(20)

Using the square-root matrix $\boldsymbol{S} \in \mathbb{R}^{p \times p}$ defined as

$$\boldsymbol{S} = \left(\boldsymbol{I} + \frac{1}{p-1} \left(\mathbb{H}\boldsymbol{A}^{f} \right)^{T} \boldsymbol{R}^{-1} \left(\mathbb{H}\boldsymbol{A}^{f} \right) \right)^{-1/2} \,. \tag{21}$$

³⁴² The analyzed anomaly is expressed from the forecasted anomaly, as:

$$\boldsymbol{A}^{a} = \boldsymbol{A}^{f} \boldsymbol{S} . \tag{22}$$

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The analyzed ensemble mean is defined using equation
$$(15)$$
,

$$\overline{\boldsymbol{X}^{a}} = \overline{\boldsymbol{X}^{f}} - \boldsymbol{G} \left(\overline{\mathbb{H}} \overline{\boldsymbol{X}^{f}} - \boldsymbol{Y} \right) , \qquad (23)$$

and the analyzed ensemble members are obtained by adding to the analyzed mean (23) the
 analyzed anomaly (22),

$$\boldsymbol{X}^{a} = \overline{\boldsymbol{X}^{a}} + \boldsymbol{A}^{a} \tag{24}$$

$$= \overline{X^a} + A^f S . \tag{25}$$

It is important to note that (25) defines the update of the ensemble in the state space. In ocean models, as indicated previously the state is typically very large. To address this issue it is preferable to work in the weight space.

351 3.1.2 ESRF in the weight space

We now express directly the ESRF in the weight space. To that end, let β^f and β^a denote the $p \times p$ weight matrices representing the forecast and analyzed ensembles, such that $X^f = X^f \beta^f$ and $X^a = X^f \beta^a$. The *i*th column of the weight matrix contains the coefficients to construct the *i*th member of the ensemble from the forecasted ensemble members. It appears clearly from its definition, that β^f is the identity matrix. This formulation is similar to decompositions found in the literature, such as in Lorenc (2003), Hunt et al. (2004), and Liu et al. (2008).

Injecting $P^{f} = A^{f} \left(A^{f} \right)^{T}$ in the gain definition leads to

$$\boldsymbol{G} = \frac{1}{p-1} \boldsymbol{A}^{f} \left(\boldsymbol{A}^{f} \right)^{T} \mathbb{H}^{T} \left(\boldsymbol{R} + \mathbb{H} \frac{1}{p-1} \boldsymbol{A}^{f} \left(\boldsymbol{A}^{f} \right)^{T} \mathbb{H}^{T} \right)^{-1} .$$
(26)

Writing $\mathbf{A} = \mathbf{X} (\mathbf{I} - 1/p)$, where 1 denotes the $p \times p$ matrix filled with one, the previous equation reads then

$$G = \frac{1}{p-1} \boldsymbol{X}^{f} \left(\boldsymbol{I} - \mathbb{1}/p \right) \left(\boldsymbol{I} - \mathbb{1}/p \right)^{T} \left(\boldsymbol{X}^{f} \right)^{T} \mathbb{H}^{T} \left(\boldsymbol{R} + \mathbb{H} \frac{1}{p-1} \boldsymbol{A}^{f} \left(\boldsymbol{A}^{f} \right)^{T} \mathbb{H}^{T} \right)^{-1} .$$
(27)

A gain matrix in the weight space, denoted $G_w \in \mathbb{R}^{p \times m}$, can thus be defined such that $G = X^f G_w$, with,

$$\boldsymbol{G}_{w} = \frac{1}{p-1} \left(\mathbb{H}\boldsymbol{A} \right)^{T} \left(\boldsymbol{R} + \frac{1}{p-1} \mathbb{H}\boldsymbol{A}^{f} \left(\mathbb{H}\boldsymbol{A}^{f} \right)^{T} \right)^{-1} .$$
(28)

³⁶⁷ Considering the mean weight vectors $\overline{\beta^{f}}$ and $\overline{\beta^{a}}$, such that $\overline{X^{f}} = X^{f}\overline{\beta^{f}}$ and $\overline{X^{a}} = X^{f}\overline{\beta^{a}}$, ³⁶⁸ which correspond to the column averages of the associated weight matrices. As β^{f} is the ³⁶⁹ identity, $\overline{\beta^{f}}$ is a *p* length vector filled with $\frac{1}{p}$. The equation (23) can thus be rewritten in ³⁷⁰ the weight space:

$$\overline{\beta^{a}} = \overline{\beta^{f}} - G_{w} \left(\mathbb{H} X^{f} \overline{\beta^{f}} - Y \right).$$
(29)

As $X^a = X^f \beta^a$ and $A^f = X^f (\beta^f)'$, where \bullet' denotes the right matrix multiplication of the operand by the matrix I - 1/p, equation (25) can be rewritten

$$\boldsymbol{X}^{f}\boldsymbol{\beta}^{a} = \boldsymbol{X}^{f}\overline{\boldsymbol{\beta}^{a}} + \boldsymbol{X}^{f}\left(\boldsymbol{\beta}^{f}\right)'\boldsymbol{S},\tag{30}$$

³⁷⁵ which yields the weight update

$$\boldsymbol{\beta}^{a} = \overline{\boldsymbol{\beta}^{a}} + \left(\boldsymbol{\beta}^{f}\right)' \boldsymbol{S} . \tag{31}$$

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3.2 Ensemble filter on RKHS family

The formalism proposed by Dufée et al. (2024), as detailed in section 2, has the 378 advantage to be directly formulated in terms of piece-wise in time, linear combination of 379 trajectories. It naturally enables the estimation of an optimal trajectory from time-series 380 of observations. More precisely, this ensemble filter (EnF) defined on the RKHS family, 381 operates on the weights, within the RKHS family, assuming the observation operator is well 382 defined on the RKHS family through the smooth evaluation operator assumption (8). In 383 the following, we will assume without loss of generality that the kernel is real-valued. The 384 filter, referred to as RKHS-EnF in the following, is built from a 4DVar-like cost function, 385 accounting for time-series of observations, and incorporating a regularization term based on 386 the norm of the kernel function in the initial RKHS \mathcal{H}_0 : 387

$$J(\mathbf{X}) = \frac{1}{2} \sum_{j=0}^{J} \left\| \mathbb{H}_{t_j} \mathbf{X}_{t_j} - \mathbf{Y}_{t_j} \right\|_{\mathbf{R}_{t_j}^{-1}}^2 + \frac{1}{2} \left\| k_0 \left(\cdot, \mathbf{X}_0 - \overline{\mathbf{X}_0^f} \right) \right\|_{\mathcal{H}_0}^2 , \qquad (32)$$

where \mathbb{H}_{t_j} , X_{t_j} , Y_{t_j} stand respectively for the observation matrix, the ensemble matrix and the observation vector at time t_j . Given β the weight vector such that $X_{t_j} = \sum_{i=1}^p \beta_i X_{t_j}^{(i)}$, and using equations (9), the second term of the right-hand side can be expressed as

$$\left\|k_0\left(\cdot, \mathbf{X}_0 - \overline{\mathbf{X}_0}\right)\right\|_{\mathcal{H}_0}^2 = \left\|\sum_{i=1}^p \left(\boldsymbol{\beta}_i - \overline{\boldsymbol{\beta}_i^f}\right) k_0\left(\cdot, \mathbf{X}^{(i)}\right)\right\|_{\mathcal{H}_0}^2.$$
(33)

³⁹³ By the kernel reproducing property, we obtain

$$\left\|k_0\left(\cdot, \boldsymbol{X}_0 - \overline{\boldsymbol{X}_0}\right)\right\|_{\mathcal{H}_0}^2 = \left(\boldsymbol{\beta} - \overline{\boldsymbol{\beta}^f}\right)^T \boldsymbol{K} \left(\boldsymbol{\beta} - \overline{\boldsymbol{\beta}^f}\right) , \qquad (34)$$

where K is the kernel matrix that contains the evaluation of the initial kernel between every member's pair of the ensemble, $K_{i,j} = k_0 \left(X_0^{(i)}, X_0^{(j)} \right)$, for $1 \le i, j \le p$. The second member of the cost function penalizes hence the weight anomaly in the RKHS.

The previous cost function can thus be rewritten as a function of the weight vector in the RKHS, thereby formulating the assimilation problem directly in the weight space:

$$J_{\text{RKHS}}\left(\boldsymbol{\beta}\right) = \frac{1}{2} \sum_{j=0}^{J} \left\| \sum_{i=1}^{p} \boldsymbol{\beta}_{i} \mathbb{H}_{t_{j}} \boldsymbol{X}_{t_{j}}^{\left(i\right)} - \boldsymbol{Y}_{t_{j}} \right\|_{\boldsymbol{R}_{t_{j}}^{-1}} + \frac{1}{2} \left\| \boldsymbol{\beta} - \overline{\boldsymbol{\beta}^{f}} \right\|_{\boldsymbol{K}}^{2} .$$
(35)

Let us define V such that $\mathbb{H}_t X_t = V_t k$ and $V = (V_{t_j})_{j \in [\![1,J]\!]}$ where $k = (k_0 (X_0, X^{(i)}))_{i \in [\![1,p]\!]}$. By the reproducing property, we have $k = K\beta$. Denoting $Y = (Y_{t_j})_{j \in [\![1,J]\!]}$ and $\tilde{R} = I_J \otimes (R_{t_j})_{j \in [\![0,J]\!]}$, where \otimes represents the Kronecker product, the cost function can be finally compactly expressed as:

$$J_{\text{RKHS}}\left(\boldsymbol{\beta}\right) = \frac{1}{2} \left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{\beta} - \boldsymbol{Y}\right)^{T} \widetilde{\boldsymbol{R}}^{-1} \left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{\beta} - \boldsymbol{Y}\right) + \frac{1}{2} \left(\boldsymbol{\beta} - \overline{\boldsymbol{\beta}^{f}}\right)^{T} \boldsymbol{K} \left(\boldsymbol{\beta} - \overline{\boldsymbol{\beta}^{f}}\right) .$$
(36)

Remark 3 In equation (35), a linear observation operator is considered through the matrix \mathbb{H}_{t_j} . However, the expression (36) remains valid for non-linear observation operators, at the condition that the observables associated with these observations belong to \mathcal{H}_t . This is the reason of introducing explicitly the notation $\mathbb{H}_t \mathbf{X}_t = \mathbf{V}_t \mathbf{K} \boldsymbol{\beta}$. This linear combination representation is a direct consequence of the reproducing property. The fact that these coefficients are constant in time ensues from the framework proposed by Dufée et al. (2024) presented in section 2.3.

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3.2.1 Updating the kernel matrix

⁴¹⁴ The inverse of the kernel matrix $K \in \mathbb{R}^{p \times p}$ can be interpreted as the weight covariance ⁴¹⁵ matrix, $P_w^f = (K)^{-1}$. We notice that due to the RKHS property expressed in equation (4), ⁴¹⁶ K is invertible. To simplify the notation, let us denote $B^f = K^{-1/2}$ as the square root ⁴¹⁷ matrix of P_w^f , so that $P_w^f = B^f (B^f)^T$. Similarly, we denote the analyzed covariance ⁴¹⁸ square root matrix as $B^a = (P_w^a)^{1/2}$. In practice, the square root of the covariance matrix ⁴¹⁹ is computed using an eigenvalue decomposition. Analogous to the ESRF, we can then define ⁴²⁰ the square root matrix S_{RKHS} ,

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$$\boldsymbol{B}^{a} = \boldsymbol{B}^{f} \boldsymbol{S}_{\text{RKHS}} , \qquad (37)$$

422 the EnF gain,

$$\boldsymbol{G}_{\text{RKHS}} = \boldsymbol{P}_{w}^{f} \left(\boldsymbol{V} \boldsymbol{K} \right)^{T} \left(\widetilde{\boldsymbol{R}} + \left(\boldsymbol{V} \boldsymbol{K} \right) \boldsymbol{P}_{w}^{f} \left(\boldsymbol{V} \boldsymbol{K} \right)^{T} \right)^{-1} , \qquad (38)$$

⁴²⁴ and the covariance matrix, which is updated according to:

$$\boldsymbol{P}_{w}^{a} = \left(\boldsymbol{I} - \boldsymbol{G}_{\text{RKHS}}\left(\boldsymbol{V}\boldsymbol{K}\right)\right)\boldsymbol{P}_{w}^{f} . \tag{39}$$

426 Substituting the gain into this equation yields:

$$\boldsymbol{B}^{a}\left(\boldsymbol{B}^{a}\right)^{T} = \boldsymbol{B}^{f}\left(\boldsymbol{I} - \left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{B}^{f}\right)^{T}\left(\widetilde{\boldsymbol{R}} + \left(\boldsymbol{V}\boldsymbol{K}\right)\boldsymbol{B}^{f}\left(\boldsymbol{B}^{f}\right)^{T}\left(\boldsymbol{V}\boldsymbol{K}\right)^{T}\right)^{-1}\left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{B}^{f}\right)\right)\left(\boldsymbol{B}^{f}\right)^{T}.$$
(40)

⁴²⁸ Upon applying the Sherman–Morrison–Woodbury formula, the above equation can be rewrit-⁴²⁹ ten as:

$$\boldsymbol{B}^{a} \left(\boldsymbol{B}^{a}\right)^{T} = \boldsymbol{B}^{f} \left(\boldsymbol{I} + \left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{B}^{f}\right)^{T} \widetilde{\boldsymbol{R}}^{-1} \left(\boldsymbol{V}\boldsymbol{K}\boldsymbol{B}^{f}\right)\right)^{-1} \left(\boldsymbol{B}^{f}\right)^{T} .$$
(41)

⁴³¹ As $B^{a} (B^{a})^{T} = B^{f} S_{\text{RKHS}} (S_{\text{RKHS}})^{T} (B^{f})^{T}$, the RKHS EnF square root matrix is finally ⁴³² identified as:

$$\boldsymbol{S}_{\text{RKHS}} = \left(\boldsymbol{I} + \left(\boldsymbol{V} \boldsymbol{K} \boldsymbol{B}^{f} \right)^{T} \widetilde{\boldsymbol{R}}^{-1} \left(\boldsymbol{V} \boldsymbol{K} \boldsymbol{B}^{f} \right) \right)^{-1/2} \,. \tag{42}$$

⁴³⁴ The matrix S_{RKHS} is then used to update the covariance square root matrix, from which ⁴³⁵ the analyzed covariance matrix can be derived. Both the gain G_{RKHS} and transition matrix ⁴³⁶ S_{RKHS} are defined for assimilating an observation time series, with these matrices being ⁴³⁷ weighted by the inverse of the kernel matrix.

438 3.2.2 Updating the weights

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It is important to note that unlike the ESRF, the RKHS-based approach does not 439 directly provide a formula for updating the ensemble weight anomalies. The matrix S_{RKHS} 440 does not update the anomalies in the same direct manner. However, since the anomaly 441 can be treated as a centered random vector with a non-singular covariance matrix, it is 442 possible to apply a whitening transformation to this vector. This is achieved by multiplying 443 it by the square root of its covariance matrix inverse, $K^{1/2}$, resulting in a decorrelated vector 444 (Koivunen & Kostinski, 1999; Friedman, 1987). To transform this vector into one that aligns 445 with the analyzed covariance matrix P_w^a , a subsequent coloring transformation is applied. 446 This transformation, being the inverse of the whitening process, involves multiplying the 447 vector by the square root of the analyzed covariance matrix $(\mathbf{P}_w^a)^{1/2}$. Consequently, the 448 updated weight anomaly can be expressed as follows: 449

$$(\boldsymbol{\beta}^{a})' = (\boldsymbol{\beta}^{f})' (\boldsymbol{P}_{w}^{f})^{-1/2} (\boldsymbol{P}_{w}^{a})^{1/2} = (\boldsymbol{\beta}^{f})' \boldsymbol{K}^{1/2} (\boldsymbol{B}^{f} \boldsymbol{S}_{\text{RKHS}}) .$$

$$(43)$$

Given that $B^f = K^{-1/2}$, it follows that the previous equation can be simplified to

$$(\boldsymbol{\beta}^{a})' = \left(\boldsymbol{\beta}^{f}\right)' \boldsymbol{S}_{\text{RKHS}} .$$
(44)

The ensemble mean can be then updated in a manner consistent with classical ensemble Kalman filters:

$$\overline{\beta}^{a} = \overline{\beta}^{f} - \boldsymbol{G}_{\text{RKHS}} \left(\boldsymbol{V} \boldsymbol{K} \overline{\beta}^{f} - \boldsymbol{Y} \right) , \qquad (45)$$

and the ensemble weights are updated as:

$$\boldsymbol{\beta}^{a} = \overline{\boldsymbol{\beta}^{a}} + \left(\boldsymbol{\beta}^{f}\right)' \boldsymbol{S}_{\text{RKHS}} , \qquad (46)$$

 $_{458}$ similarly as in (31).

A similar approach can be considered in the ESRF context to express the ensemble update. Given that $A^a = A^f S$, the transition matrix reads $S = (A^f)^{-1} A^a$. Within the ESRF framework, where the anomaly is assumed to be the square root of the covariance matrix, the update of the anomaly can also be interpreted as a whitening transformation followed by a coloring transformation. In the ESRF setting, the change of variable $(P_w^f)^{-1/2}\beta$ allows to work with variables decorrelated from each other – through the diagonalisation of the covariance matrix. In the RKHS-based case, $(P_w^f)^{-1/2} = K^{1/2}$, and the equivalent change of variable procedure constitutes an embedding in the RKHS.

3.3 Link between ESRF and EnF on a RKHS family

The ESRF can be viewed as a specific instance of RKHS-EnF, given an appropriately selected initial kernel. To establish the connection between these two techniques, we derive in this section the ESRF equations from the RKHS EnF equations by selecting a suitable initial kernel.

472 Let us consider a RKHS EnF which assimilates observations directly at the time they 473 are available, as a consequence $V = V_t$ and $\tilde{R} = R_t$ for a fixed time instant t. As $k = K\beta$, 474 by the definition of V, we have

$$\boldsymbol{V}_t \boldsymbol{K} \boldsymbol{\beta} = \mathbb{H} \boldsymbol{X}_t \;, \tag{47}$$

and it follows that $V_t K \beta^f = \mathbb{H} X_t^f$. However, as seen previously β^f is equal to the identity matrix. This represents the forecast members' weights expressed in their own basis, therefore $V_t K = \mathbb{H} X_t^f$. For simplicity, the time index t is omitted in the subsequent discussion since all observations are assumed to be taken simultaneously. Assuming the kernel is chosen such that $(K)^{-1} = \frac{1}{p-1} (I - 1/p) I (I - 1/p)$, the square root of K is given by $B^f = \frac{1}{\sqrt{p-1}} (I - 1/p)$, and the RKHS-EnF gain reads

$$\boldsymbol{G}_{\text{RKHS}} = \frac{1}{p-1} \left(\boldsymbol{I} - \mathbb{1}/p \right) \boldsymbol{I} \left(\boldsymbol{I} - \mathbb{1}/p \right) \left(\mathbb{H} \boldsymbol{X}^{f} \right)^{T} \\ \left(\widetilde{\boldsymbol{R}} + \frac{1}{p-1} \left(\mathbb{H} \boldsymbol{X}^{f} \right) \left(\boldsymbol{I} - \mathbb{1}/p \right) \boldsymbol{I} \left(\boldsymbol{I} - \mathbb{1}/p \right) \left(\mathbb{H} \boldsymbol{X}^{f} \right)^{T} \right)^{-1} .$$
(48)

It is worth noting that the kernel K defined here has a Mercer spectrum $\mu_i = 1$ for $0 \le i \le p$, and can be interpreted as the zero-correlation limit of a Gaussian, tending towards a Dirac measure, which is not a RKHS reproducing kernel (as it is unbounded). As $\mathbf{X}^f (\mathbf{I} - \mathbb{1}/p) =$

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$$A^{f}$$
 and $(I - 1/p)^{2} = (I - 1/p)$, we have

$$\boldsymbol{G}_{\text{RKHS}} = \frac{1}{p-1} \left(\mathbb{H}\boldsymbol{A}^{f} \right)^{T} \left(\widetilde{\boldsymbol{R}} + \frac{1}{p-1} \left(\mathbb{H}\boldsymbol{A}^{f} \right) \left(\mathbb{H}\boldsymbol{A}^{f} \right)^{T} \right)^{-1} = \boldsymbol{G}_{w} .$$
(49)

488 Considering the matrix S_{RKHS} under the same assumption, we get:

$$\mathbf{S}_{\text{RKHS}} = \left(\mathbf{I} + \left(\mathbb{H} \mathbf{X}^{f} \frac{1}{\sqrt{p-1}} \left(\mathbf{I} - \mathbb{1}/p \right) \right)^{T} \widetilde{\mathbf{R}}^{-1} \left(\mathbb{H} \mathbf{X}^{f} \frac{1}{\sqrt{p-1}} \left(\mathbf{I} - \mathbb{1}/p \right) \right) \right)^{-1/2}, \quad (50)$$

⁴⁹⁰ which simplifies to:

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$$\boldsymbol{S}_{\text{RKHS}} = \left(\boldsymbol{I} + \frac{1}{p-1} \left(\mathbb{H} \boldsymbol{A}^{f} \right)^{T} \widetilde{\boldsymbol{R}}^{-1} \left(\mathbb{H} \boldsymbol{A}^{f} \right) \right)^{-1/2} = \boldsymbol{S} .$$
 (51)

⁴⁹² Consequently, performing the assimilation step each time an observation is available, using a ⁴⁹³ well-chosen kernel, provides an RKHS EnF gain and a square root matrix that are equivalent ⁴⁹⁴ to those of a classical square root filter. Furthermore, since under these conditions VK =⁴⁹⁵ $\mathbb{H}X$, the two formulas used to update the ensemble are equivalent and the two filters are ⁴⁹⁶ equivalent.

3.4 Ensemble update strategy

The embedding of the dynamics into the RKHS family ensures that the ensemble 498 weights remain invariant over time. This invariance provides the reconstruction of the an-499 alyzed ensemble members at any time of the assimilation window, as a linear combination 500 of the forecast ensemble. The weights are calculated with respect to the initial kernel, and 501 consequently, in the method proposed herein, it is natural to construct the analyzed mem-502 bers at the beginning of the assimilation window. Due to the non-linear dynamics of the 503 system and the finite number of ensemble members, the analyzed members reconstructed at 504 the beginning of the assimilation window and subsequently propagated forward in time may 505

differ at the end of the assimilation window from the analyzed members. It is noteworthy 506 that, according to the theoretical framework, if an infinite number of ensemble members 507 were available, for an invertible dynamics the reconstruction at the beginning of the as-508 similation window, when transported by the dynamical system, would be identical to the 509 reconstruction at the end of the assimilation window. However, obviously due to the finite 510 size of the ensemble members and due to noninversible numerical dissipation process, this 511 is only approximately true. The ability of a direct reconstruction of the initial condition 512 at the beginning of the assimilation window is used to forecast updated trajectories of the 513 ensemble members. This approach referred to as the tiled RKHS ensemble filter is described 514 in Algorithm 1. 515

Remark 4 The tiled RKHS EnF is consistent with the theoretical framework, as all estimations are re-expressed at the initial time t_0 through the isometry between the evolving RKHSs. This procedure has also the advantage of introducing a spin-up phase during the assimilation window, where spurious numerical artefacts induced by the finite size linearcombination reconstructions are filtered-out by the dynamical system, and are ultimately replaced by physically relevant features.

Remark 5 The resulting tiled RKHS ensemble filter shares similarities with Bayesian smoo-522 thers and 4D-Var methods, with the key distinction that it avoids multiple forward-backward 523 integrations of the dynamical system, as we infer directly the initial condition of the sys-524 tem (at the beginning of the assimilation window). In the proposed method, the model is 525 integrated forward only once to propagate the analyzed ensemble to the next assimilation 526 window. An intermediate method, which can be referred to as sequential RKHS EnF, can 527 also be devised by reconstructing the ensemble at the end of the assimilation window, akin 528 to classical ensemble filters. Although this approach does not fully leverage the theoretical 529 framework's potential, it maintains the same computational cost as a classical sequential 530 filter. 531

Algorithm 1 Tiled RKHS EnF

Require: Ensemble X_0 , initial time t_0 , end time t_{end} , time step dt, assimilation interval $\begin{array}{l} T_a. \\ X_t^f \leftarrow X_0 \end{array}$ $t \leftarrow t_0$ $t_a \leftarrow T_a$ the assimilation time horizon. Save X_t for future reconstruction. Construct the kernel matrix K. Initialize VK and Y as empty. while $t < t_{end}$ do $t \leftarrow t + dt$ Propagates the ensemble members in time to update X_t^f . Append to VK the observations of the ensemble members. Append to Y the reference observations. if It is time to assimilate i.e. t is a multiple of t_a then $\beta^a \leftarrow \text{assimilate}(VK, Y, K)$ \triangleright Equations (45) and (46). Empty VK and Y. Apply the weights β^a to $X_{t-t_a}^f$ saved at $t - t_a$, to obtain $X_{t-t_a}^a$. Propagates $X_{t-t_a}^a$ on the time interval $[t - t_a, t]$, to obtain X_t^a . $X_t^f \leftarrow X_t^a$ Save X_t^f for future reconstruction. Update the kernel matrix K. end if end while

532 4 Experimental framework

In the following we first describe the experimental setup we built to assess the data assimilation we propose.

4.1 Dynamical model

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The data assimilation methods previously presented are compared using a North At-536 lantic Ocean model, covering the region between 9°N and 48°N in latitude and -98°E and 537 -3.84°E in longitude with realistic coasts. This area represents a domain of 4329 km by 8658 538 km. The ocean bottom is assumed to be flat, and all boundaries are considered closed, even 539 those at the equator and in the north, which are normally open boundaries. Due to the flat 540 bottom, the Gulf Stream develops close to the North American coast. To achieve a well-541 positioned Gulf Stream, areas with a bathymetry shallower than -100 meters are treated 542 as land. A realistic wind forcing is considered. A multi-layer quasi-geostrophic model im-543 plemented as the projection of a rotating shallow-water (Thiry et al., 2024) is considered. 544 The rotating shallow water is simulated with the model proposed by Roullet and Gaillard (2022), while the QG projection uses the elliptic solver of Thiry et al. (2023). Three layers 546 of heights H = (400, 1100, 2600) are considered. The evolution of three prognostic variables 547 are simulated: u, v the two components of the horizontal velocity vector, and h the height 548 anomaly in each layers. These variables are propagated over time according to the following 549 equations: 550

$$\begin{cases} \partial_t \boldsymbol{u} = (\boldsymbol{\omega} + \boldsymbol{f}) \, \boldsymbol{v} - \partial_x \, (\boldsymbol{p} + \boldsymbol{k}) \\ \partial_t \boldsymbol{v} = - \, (\boldsymbol{\omega} + \boldsymbol{f}) \, \boldsymbol{u} - \partial_y \, (\boldsymbol{p} + \boldsymbol{k}) \\ \partial_t \boldsymbol{h} = -\boldsymbol{H} \, (\partial_x \boldsymbol{u} + \partial_y \boldsymbol{v}) - \partial_x \, (\boldsymbol{u} \boldsymbol{h}) - \partial_y \, (\boldsymbol{v} \boldsymbol{h}) \end{cases} , \tag{52}$$

with $\boldsymbol{\omega} = \partial_x \boldsymbol{v} - \partial_y \boldsymbol{u}$ the relative vorticity, \boldsymbol{p} the pressure vector, $\boldsymbol{k} = (\boldsymbol{u}^2 + \boldsymbol{v}^2)/2$ the kinetic energy and $\boldsymbol{f}_i = f_0 + \beta y$ for $0 \le i \le 3$ the Coriolis parameter under a beta plane assumption, with $f_0 = 9.375 \times 10^{-5} \,\mathrm{s}^{-1}$ the mean Coriolis parameter, and $\beta = 1.754 \times 10^{-11} \,\mathrm{m}^{-1} \mathrm{s}^{-1}$ the Coriolis parameter gradient. The pressure values in layer $1 \le i \le 3$ are define such that $p_i = \rho_1 \sum_{j=1}^i \left(g'_j \sum_{k=j}^3 h_k\right)$ with $g_1 = g = 9.81 \,\mathrm{m} \cdot \mathrm{s}^{-2}$ and $g'_i = g\left(\rho_i - \rho_{i-1}\right)/\rho_1$ for i = 2, 3. ρ is the density in the different layers.

The QG equations are then obtained by projecting this shallow water system using a projector involving an elliptic operator. The full details of this procedure as well as the numerical codes can be found in Thiry et al. (2024). The simulation is performed on a $512 \times 1024 \times 3$ grid ($n_x = 512$, $n_y = 1024$, $n_l = 3$), resulting in a mesh size of approximately 8.46×8.46 kilometers square. The time step is set to 2000 seconds.

563 4.2 Ensemble generation

The initial ensemble X_0 is generated using a noise based on the spectral decomposition 564 of random local fluctuations (Bauer et al., 2020; Brecht et al., 2021). This technique is 565 applied to perturb the velocity field. Given a velocity field \boldsymbol{u} defined on a $n = n_l \times n_x \times n_y$ 566 grid, a local spatial window $W_{i,j}$ of size $n_w \times n_w$ (n_w being odd) in the horizontal direction 567 is considered around each grid point $x_{i,j}$, with $1 < i < n_x$ and $1 < j < n_y$. Within 568 each window, n_o samples are randomly picked, where $n_o > 1$. The sampled values within the window $W_{i,j}$ are denoted as $\boldsymbol{w}_k^{(i,j)}$, where $1 \leq k \leq n_o$, and these values represent n_l -569 570 dimensional vectors associated with the number of vertical layers. These values are centered 571 by subtracting their mean as follows: 572

$$\left(\boldsymbol{w}_{k}^{(i,j)}\right)' = \boldsymbol{w}_{k}^{(i,j)} - \overline{\boldsymbol{w}}^{(i,j)} \quad \text{with} \quad \overline{\boldsymbol{w}}^{(i,j)} = \frac{1}{n_{o}} \sum_{\ell=1}^{n_{o}} \boldsymbol{w}_{\ell}^{(i,j)}.$$
(53)

Hence a matrix $\mathbf{Z} \in \mathbb{R}^{n \times n_o}(\mathbb{R})$ is constructed by organizing the centered values $\left(\boldsymbol{w}_k^{(i,j)}\right)'$ for each grid point in the $n_l \times n_x \times n_y$ grid. A singular value decomposition (SVD) is performed



Figure 2: Observation masks for the periods of 33 minutes and 20 seconds (left) and 41 hours and 40 minutes (right), both from January 1, 2022.

 $Z = U\Sigma V^T$, allowing to define perturbation modes U_i from the columns of U and the diagonal matrix of singular values $\Sigma_{jj} = \lambda_j$. The perturbation, $\{X_0^{'(i)}, i = 1..., p\}$ are generated from the p eigenpairs of the empirical covariance matrix:

$$\mathbf{X}^{'(i)} = \sum_{j=1}^{n_o} \lambda_j \mathbf{U}_j \, \eta_j^{(i)}, \tag{54}$$

where $\eta \in \mathbb{R}^{n_o}$ is a multivariate random vector composed of n_o standard Gaussian variables and $\eta^{(i)}$ denotes the *i* th realization. The initial ensemble members are obtained by adding these random fields to the initial velocity component. In the experiments, 76 members have been generated using this method, with $n_w = 5$ and $n_o = 21$. One member is retained as the reference, while the remaining 75 members constitute the training set. The reference member is not included in the training ensemble.

4.3 Observations simulation

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An Observing System Simulation Experiment (OSSE) was conducted using sea surface 587 height (SSH) as the observed variable. The SSH corresponds to the sum of the layer height 588 anomalies h along the vertical. Synthetic observations were generated from the reference 589 simulation by applying spatial masks derived from realistic data provided by the Copernicus 590 along-track sea surface height database (DOI https://doi.org/10.48670/moi-00147). The 591 initial observation date was set to January 1, 2022. Figure 2 illustrates the observation 592 masks for two periods: one over 33 minutes and 20 seconds, corresponding to a single 593 simulation time step, and another over 41 hours and 40 minutes. It can be shown that over 594 41h40min, the altimeters swaths constitute a reasonable coarse grid observational coverage 595 of the basin. Usual OI methods, provide a dense field of observation, by interpolating these 596 observation spatially assuming they are all acquired at the same time. In the technique we 597 propose, the observations are treated as a time-series and are not interpolated. 598

Four distinct experiments were carried out. In the first two, observations were per-599 turbed with additive independent white Gaussian noise. The noise standard deviations were 600 set to 1% and 10% of the root mean square (RMS) of the reference member's SSH at the 601 initial time, corresponding to standard deviations of $\sigma_o = 0.18$ cm and $\sigma_o = 1.8$ cm, respec-602 tively. Considering spatially uncorrelated white Gaussian noise is a strong idealization of 603 real-world data noise. In operational setups, the noise is correlated, as raw sea surface height 604 measurements are corrected to remove barotropic tidal effects and the predictible internal 605 tide waves, compensate for satellite jitters, and account for atmospheric turbulence. 606

⁶⁰⁷ Due to this, two additional experiments with spatially correlated noise were conducted. ⁶⁰⁸ The noise was generated by smoothing a field of white Gaussian noise using a Gaussian

filter with a standard deviation of 250 km and a support length of 300 km. This approach 609 produced noise patterns with spatial correlations and, notably, the resulting noise was non-610 Gaussian due to the relatively limited support area of the filter. The initial variance of 611 the white Gaussian noise was calibrated so that the mean covariance square root of the 612 resulting spatially correlated noise was approximately of 1% and 10% of the SSH RMS 613 (0.18 cm and 1.8 cm respectively). Although this noise is more realistic than white Gaussian 614 noise, it is still obviously idealized. In all experiments, the observation error covariance 615 matrix R used in the filters was assumed to be diagonal, with all observations having the 616 same error variance σ_{α}^2 . As a result, the filter was not informed of the correlations between 617 observations, posing a challenge for the data assimilation method to remain robust despite 618 this misinformation. 619

4.4 Filters Setting

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4.4.1 Assimilation frequency

In the context of the RKHS EnF, the assimilation frequency can be adjusted to opti-622 mize the reconstruction quality. In the one hand, if the assimilation window is too short, the 623 observation time-series might be too short to exploit all the potential benefit of the RKHS 624 EnF. In the limit case where observations are sequentially assimilated at each time step 625 and the kernel is defined as described in section 3.3, the RKHS EnF and ESRF filters are 626 equivalent. However, an excessively long assimilation window may violate the invertibility 627 assumption of the dynamical system, leading to poorer outcomes for the ensemble linear 628 combination of trajectories over that period. Conversely, for the tiled version of the filter, a 629 longer window can enhance the development of small-scale feature variability and assist in 630 eliminating non-physical structures within the dynamical system. 631

In the following, for the untiled sequential version of the RKHS EnF, the assimilation 632 window length is set to 8 hours and 20 minutes when the observations are perturbed by 633 white Gaussian noise, and to 2 hours, 46 minutes, and 40 seconds when perturbed by spatially correlated noise. These correspond to analysis updates every 15 and 5 time steps 635 of the dynamics, respectively. The presence of spatially correlated noise degrades the filter's 636 performance, necessitating more frequent corrections to the ensemble. For the tiled version, 637 the assimilation window is set to 41 hours and 40 minutes, which corresponds to 75 time 638 steps of the dynamics, regardless of the type of noise affecting the observations. This 639 suggests a robustness in the tiled EnF's configuration, making it less sensitive to the noise 640 characteristics. The ESRF is updated every time step (*i.e.* every 33 minutes and 20 seconds). 641

Remark 6 It is important to note that a tiled version of the ESRF is theoretically not directly feasible. The possibility of using this strategy with the RKHS EnF arises from the stationary weights associated with the linear combination of trajectories and the time series of observations. Obtaining the assimilated state at the beginning of the assimilation windows requires either a forward-backward Bayesian smoother or the application of 4D VAR techniques.

648 4.4.2 Kernel selection

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⁶⁴⁹ In the experiments, a centered Gaussian kernel, defined by the following expression, ⁶⁵⁰ was employed:

$$\forall \boldsymbol{X}, \boldsymbol{Y} \in \Omega_{0}, \ k_{g}\left(\boldsymbol{X}, \boldsymbol{Y}\right) = \exp\left(-\frac{\left\|\boldsymbol{u}_{\boldsymbol{X}} - \boldsymbol{u}_{\boldsymbol{Y}}\right\|_{\Omega_{x}}^{2} + \left\|\boldsymbol{v}_{\boldsymbol{X}} - \boldsymbol{v}_{\boldsymbol{Y}}\right\|_{\Omega_{x}}^{2}}{l_{G}^{2}}\right), \tag{55}$$

where $(\boldsymbol{u}_{\boldsymbol{X}}, \boldsymbol{v}_{\boldsymbol{X}})$ and $(\boldsymbol{u}_{\boldsymbol{Y}}, \boldsymbol{v}_{\boldsymbol{Y}})$ denote the velocity vectors associated with states \boldsymbol{X} and \boldsymbol{Y} , respectively, and $\|\boldsymbol{\bullet}\|_{\Omega_x}$ denoting the L^2 norm over the domain Ω_x . The kernel matrix associated with this kernel is denoted \boldsymbol{K}_G . The length scale parameter l_G is chosen to maintain a constant ratio between the smallest and largest Mercer eigenvalues of \boldsymbol{K}_G which is set to 10^{-2} . The Gaussian kernel matrix is then normalized by multiplying it by p-1and then centered by pre- and post-multiplying it by the projection operator (I - 1/p):

$$\boldsymbol{K} = \alpha(p-1) \left(\boldsymbol{I} - \mathbb{1}/p \right) \boldsymbol{K}_{G} \left(\boldsymbol{I} - \mathbb{1}/p \right) , \qquad (56)$$

(57)

(61)

where α is a constant. This centering procedure is simply equivalent to a change of variable 659 toward the anomaly of the coefficients $\beta \mapsto \beta'$. Consequently, this is equivalent to work with 660 a full rank kernel matrix (*i.e.* the kernel and the associated RKHS is well defined), which 661 operates on the weight anomalies. Specifically, α is set to 0.75 for uncorrelated Gaussian 662 noise, 2 for correlated (non-Gaussian) low noise, and 5 for correlated (non-Gaussian) high 663 noise. As a rule of thumb, for correlated noise, α is set to a value greater than 1. This 664 adjustment compensates for the correlations not accounted for in the R matrix. The higher 665 the correlation between observations, the more α should be increased. 666

⁶⁶⁷ Otherwise, since the rank of (I - 1/p) is p - 1, the kernel matrix K is not invertible. ⁶⁶⁸ Thus, its Moore-Penrose pseudo-inverse must be used. Given that (I - 1/p) is a projection ⁶⁶⁹ matrix, it equals its own pseudo-inverse, thus the Moore-Penrose inverse of K is given by:

$$\boldsymbol{K}^{\dagger} = \frac{1}{\alpha \cdot (p-1)} \left(\boldsymbol{I} - \mathbb{1}/p \right) \boldsymbol{K}_{G}^{-1} \left(\boldsymbol{I} - \mathbb{1}/p \right).$$

The square root of the kernel matrix pseudo-inverse is typically computed through SVD decomposition. The kernel pseudo-inverse and its square root are used to compute the gain (38) and the covariance square root (42). The centering of the kernel ensures consistency with classical square root filtering methods. The kernel used in section 3.3 to demonstrate the connection between the ESRF and the RKHS EnF is a specific case of the kernel presented here, with l_G tending towards 0.

4.5 Diagnostics

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To assess the reconstruction quality of each methods, we consider the reconstructed relative vorticity fields. Relative vorticity is a less smooth variable compared to velocity or water height, making it a more effective discriminator between the different reconstruction methods. Three metrics are considered to evaluate the effectiveness of the methods. In this subsection the vorticity of the *i* th ensemble member to be evaluated is denoted as $\omega^{(i)}$ and the reference is denoted as ω_{ref} .

The root mean square error (RMSE) is a widely used metric for quantifying the discrepancy between the ensemble members and the reference state. It is defined as:

$$e_{\text{RMSE}} = \sqrt{\frac{1}{p} \sum_{i=1}^{p} \left\| \boldsymbol{\omega}^{(i)} - \boldsymbol{\omega}_{\text{ref}} \right\|_{\Omega_{x}}^{2}}.$$
(58)

It penalizes the deviation of each ensemble member from the reference, providing a measure of the overall accuracy of the ensemble. To improve interpretability, the RMSE can be expressed as a percentage by normalizing it with respect to the norm of the reference state, as follows:

$$e_{\text{NRMSE}} = \frac{e_{\text{RMSE}}}{\|\boldsymbol{\omega}_{\text{ref}}\|_{\Omega_x}}.$$
(59)

This normalized RMSE offers a more intuitive understanding of the ensemble's mean performance. It is important to note that this measure reflects the mean accuracy of the individual ensemble members, rather than the accuracy of the ensemble mean, which is captured by the bias. The bias is defined as the distance between the reference and the ensemble mean estimation:

$$\boldsymbol{\omega}_{\mathrm{bias}} = \overline{\boldsymbol{\omega}} - \boldsymbol{\omega}_{\mathrm{ref}} \ .$$
 (60)

⁶⁹⁸ By computing the norm of the bias vorticity field over the domain Ω_x , a scalar value can ⁶⁹⁹ be obtained to represent the bias error across the entire spatial domain:

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$$e_{ ext{bias}} = \left\| oldsymbol{\omega}_{ ext{bias}}
ight\|_{\Omega_{T}} \; .$$

This indicator represent the accuracy of the ensemble, the lower it is the more accurate is the estimation. The standard deviation is expressed as

$$\sigma_e = \sqrt{\frac{1}{p} \sum_{i=1}^{p} \left\| \boldsymbol{\omega}^{(i)} - \overline{\boldsymbol{\omega}} \right\|_{\Omega_x}^2} \,. \tag{62}$$

The standard deviation represents the dispersion of the ensemble. A small standard deviation indicates that the ensemble members are close to each others. The three metrics presented here are interconnected, considering the following inequality for $1 \le i \le p$,

$$\left\|\boldsymbol{\omega}^{(i)} - \boldsymbol{\omega}_{\mathrm{ref}}\right\|_{\Omega_{x}}^{2} \leq \left\|\boldsymbol{\omega}^{(i)} - \overline{\boldsymbol{\omega}}\right\|_{\Omega_{x}}^{2} + \left\|\overline{\boldsymbol{\omega}} - \boldsymbol{\omega}_{\mathrm{ref}}\right\|_{\Omega_{x}}^{2}, \tag{63}$$

⁷⁰⁸ which gives with an empirical mean

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$$e_{\rm RMSE}^2 \le \sigma_e^2 + e_{\rm bias}^2 \,. \tag{64}$$

A low standard deviation may indicate a good result if the bias is also low, meaning that all ensemble members are close to the reference, resulting in a low RMSE. However, if the ensemble spread is too small, there is a risk that the reference may fall outside the ensemble's range, making it difficult for the method to adjust the ensemble members to track the reference in the future. This can lead to divergence over time. Therefore, a larger standard deviation within the ensemble can contribute to more robust predictions.

T16 Striking a balance between standard deviation and bias, known as the bias-variance T17 trade-off, is crucial. A standard deviation that is too low can lead to divergence, while T18 one that is too high may indicate that individual ensemble members poorly represent the T19 system. A wide spread can also impact the bias, as overly dispersed ensemble members may T20 produce an over-smoothed mean, failing to capture small-scale structures accurately.

721 5 Results

In this section, we compare the performance of the data assimilation methods presented 722 in this study over 100 days. Let us mention that the state space dimension is large, while 723 the number of ensemble members is limited, making this scenario particularly challenging 724 for ensemble methods. It is well known that in such cases, localization techniques are 725 recommended to prevent filter divergence. However, we will not consider them here, as the 726 focus is on the efficiency and robustness provided by the RKHS embedding, as well as its 727 capacity to handle a 4D VAR problem at low cost using the tiled simulation strategy. While 728 extending the proposed methods with localization techniques is possible, it lies beyond the 729 scope of this paper and presents a clear avenue for future research. 730

As an initial comparison, Figure 3 shows the NRMSE for both the classical ESRF and the RKHS EnF. The NRMSE is averaged over five different reference datasets, each with its own initial ensemble, generated as described in Section 4.2. The shaded areas around the mean NRMSE curves represent the standard deviation across these experiments. A wide shaded area indicates high sensitivity of the results to both the initial reference (from which the observations are extracted) and the initial ensemble, while a narrow band highlights low variability in performance, suggesting greater robustness.

In the experiments with 1% Gaussian observation noise, the classical ESRF shows a
 constant increase in error and significant variability, indicating a lack of robustness. The
 RKHS EnF performs similarly, though with slightly lower variability. In contrast, the tiled
 RKHS EnF demonstrates remarkable long-term stability, maintaining consistent performance along time.

For higher, decorrelated noise (10% of the reference height field RMS), all methods experienced an increase in error. However, the tiled version of the RKHS EnF outperformed the ESRF, achieving lower NRMSE with smaller standard deviation across different experiments, demonstrating greater resilience to high noise levels. The ESRF performs better
with uncorrelated noise compared to the RKHS EnF, but similarly to the lower noise level,
it exhibits slightly higher robustness in terms of a tighter standard deviation.

For 1% spatially correlated non-Gaussian noise, the tiled RKHS EnF significantly outperforms both ESRF and RKHS EnF, maintaining consistently low error and much smaller standard deviation. This highlights its strong robustness against misrepresentation of the observation covariance, which was assumed diagonal and hence representative of independent, uncorrelated spatial noise. ESRF and RKHS EnF perform similarly in terms of NRMSE average and standard deviation for this noise type.

With higher spatially correlated noise (10%), all methods see a more substantial in-755 crease in NRMSE. However, both the RKHS EnF and its tiled version significantly out-756 perform the ESRF, showing greater robustness (with a tighter standard deviation) and 757 maintaining a constant NRMSE at the noise level up to day 50. In contrast, the ESRF 758 shows increasing error from around day 15. This experiment reveals a clear separation in 759 the filters' performance. Theoretically, ESRF is not well-suited to nonlinear, non-Gaussian 760 systems, whereas the RKHS EnF appears better equipped to handle non-Gaussianity and 761 correlated noise. 762

As noted earlier, the performance of these filters could likely be improved with the use of localization techniques. For the RKHS EnF, localization could be achieved by localizing the observables and with the consequence of introducing localized stationary weights (i.e stationary spatial vectors of weights) instead of global weights.

These experiments demonstrate that the tiled version of the RKHS EnF offers clear 767 improvements. This approach consistently produces better results, showing less sensitivity 768 to initial conditions (*i.e.*, the initial ensemble) by showcasing a tighter standard deviation 769 for all noise levels. Additionally, the RKHS EnF is notably less sensitive to errors in the 770 noise covariance modelling. The tiled RKHS EnF benefits from the theoretical properties of 771 embedding the dynamical system in the dynamically evolving RKHS family, allowing it to 772 formulate a 4D-Var assimilation as a single-iteration filtering process, similar to sequential 773 Bayesian filters. 774

Figure 4 presents the analyzed ensemble mean of the upper-layer relative vorticity 775 after 100 days, using observations perturbed by the 10% non-Gaussian, spatially correlated 776 noise. This figure shows that all assimilation methods successfully reconstruct a vorticity 777 map close to the reference vorticity field. At naked eye the estimations are almost indistin-778 guishable from the reference. The enstrophy spectra from the upper layer, calculated within 779 the rectangular region bounded by -72.25°E to -42.28°E longitude and 14.64°N to 39.47°N 780 latitude, following the method described by Durran et al. (2017) are shown in Figure 5. 781 This figure reveals that the enstrophy spectra for the three methods are nearly identical 782 to the reference one, indicating accurate reproduction of the reference field energy across 783 all scales. Enstrophy spectra for the other layers lead to the same conclusion (not shown). 784 Importantly, the effective resolution of the reference data has not been compromised. This 785 suggests that employing ensemble methods, without resorting to optimal interpolations or 786 gridded data prior to assimilation, maintains the resolution integrity. 787

However, a clearer understanding of the differences between the methods is provided by the bias fields, which represent the discrepancies between the reference and analyzed fields. Figures 6, figure 7, and 8 illustrate the relative vorticity bias fields for the three assimilation methods across the upper, middle, and lower layers, respectively. Notably, the middle and lower layers are better reconstructed by all methods compared to the upper layer, likely due to the reduced variability and the fewer small-scale structures in these layers. To improve clarity, the color scale for the bias fields in these two layers has been refined. An analysis of the bias fields reveals similar trends across all three layers. The classical ESRF



Figure 3: Comparison of the Normalized Root Mean Square Error (NRMSE) for the ESRF (blue), RKHS EnF (red), and Tiled RKHS EnF (green). The first two rows display results for experiments with additive white Gaussian noise, while the last two rows present results for experiments with non-Gaussian, spatially correlated noise. Rows one and three correspond to experiments with 1% standard deviation noise, and rows two and four to experiments with 10% standard deviation noise.



Figure 4: Reference upper layer relative vorticity field (top row) after 100 days, analyzed fields for the ESRF (second row), RKHS EnF (third row), and tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 5: Upper layer enstrophy spectra after 100 days within the region defined by -72.25°E to -42.28°E longitude and 14.64°N to 39.47°N latitude, presented as a function of wavenumber. The reference spectra (black) is compared to the results from the ESRF (blue), RKHS EnF (orange), and tiled RKHS EnF (green), all corresponding to the experiment with 10% standard deviation spatially correlated noise.

exhibits significant bias, particularly along the Gulf Stream, indicating poorer performance in capturing the dynamics of this region. In contrast, the RKHS EnF demonstrates improved results, with lower bias compared to the ESRF. The tiled RKHS EnF outperforms both methods, producing substantially lower biases. This suggests that the tiled method assimilates the data more effectively, leading to more accurate reconstructions. The study of the bias fields aligns with the conclusions drawn from the RMSE analysis done in the previous paragraph. This consistency is expected, given the inherent connection between these diagnostic variables.

To further quantify the errors, Figure 9 presents the enstrophy bias spectra for the 804 upper layer. The spectra indicate that most errors occur at wavenumbers greater than 10^{-2} , 805 corresponding to structures smaller than 600 km. The bias error energy of the RKHS EnF 806 is slightly lower than that of the ESRF, while the tiled RKHS EnF exhibits significantly 807 reduced error energy. Notably, the improvement is substantial across all scales, from large-808 scale features to the eddy-diffusion cut-off, with the greatest enhancement around 300 km. 809 This highlights the superior performance of the tiled method in capturing both large-scale 810 and small-scale features in the vorticity field. 811

To finally complement these analysis results, we present in Figure 10 the RMSE ob-812 tained for the upper layer. As observed, the ESRF exhibits the highest RMSE, while the 813 RKHS EnF shows a lower RMSE, and the Tiled RKHS EnF further improves upon it. The 814 other layers, which follow the same trend of lower RMSE, are not shown. The spread (stan-815 dard deviation) of the ensemble for the upper layer is also displayed in Figure 11. We observe 816 that the spread is attenuated for the ESRF, suggesting a collapse of the ensemble. Both 817 the RKHS EnF and Tiled RKHS EnF display some variation in spread. While the spread 818 is weak for the Tiled RKHS EnF, the bias is also minimal, indicating a less problematic 819 situation compared to the ESRF. 820

6 Conclusion and perspectives

This study introduces a novel approach for assimilating observations of a non-linear 822 dynamical system by using a family \mathcal{W} of reproducing kernel Hilbert spaces (RKHS). Each 823 RKHS in this family is associated with kernels k_t , which are defined by transporting an 824 initial kernel k_0 along the system's dynamics up to time t. This framework establishes an 825 isometry between the evolving RKHS, fully justifying the superposition principle, allow-826 ing new system trajectories to be constructed as linear combinations of known ensemble 827 members trajectories. Importantly, the isometry ensures these linear combinations remain 828 invariant over time. 829

An ensemble filter is developed within this RKHS family. The theoretical properties 830 of this framework naturally enable the assimilation of time series of observations, effectively 831 accumulating data at different time instants before analyzing the entire trajectory within 832 the assimilation window. This leads to more robust outcomes. By leveraging the frame-833 work's properties, the analyzed ensemble can be provided at the start of the assimilation 834 window and propagated forward. This tiled version of the filter has demonstrated superior 835 performance in reconstructing physical structures and improving accuracy in estimating the 836 reference state. 837

The methods presented in this study have been applied to a quasi-geostrophic model of the North Atlantic Ocean and compared to a classical square root filter (ESRF), which can be viewed as a limit of the sequential version of the RKHS-based ensemble filter.

Improvements to the RKHS EnF could be made by exploring methods to optimize the choice of kernels. One such approach, as proposed by Akian et al. (2022), involves an algorithm that adjusts the kernel eigenvalues to minimize accuracy loss when data are removed from a training set. Investigating alternative kernels beyond the Gaussian could also be beneficial. For example, as noted by Stein (1999), the Gaussian kernel may be too



Figure 6: Reference upper layer relative vorticity field (top row) after 100 days, bias field corresponding to it for the ESRF (second row), RKHS EnF (third row), tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 7: Reference middle layer realtive vorticity field (top row) after 100 days, bias field corresponding to it for the ESRF (second row), RKHS EnF (third row), tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 8: Reference lower layer realtive vorticity field (top row) after 100 days, bias field corresponding to it for the ESRF (second row), RKHS EnF (third row), tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 9: Upper layer enstrophy bias spectra after 100 days within the region bounded by -72.25°E to -42.28°E longitude and 14.64°N to 39.47°N latitude, plotted as a function of wavenumber. Results are shown for the ESRF (blue), EnF on an RKHS family (orange), and tiled EnF on an RKHS family (green), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 10: Upper layer relative vorticity RMSE field after 100 days for the ESRF (top row), RKHS EnF (second row), tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.



Figure 11: Upper layer relative vorticity standard deviation field after 100 days for the ESRF (top row), RKHS EnF (second row), tiled RKHS EnF (bottom row), all corresponding to the experiment with 10% standard deviation spatially correlated noise.

smooth, whereas the Matérn kernel offers greater flexibility in terms of smoothness while
 maintaining a reasonable number of parameters to configure.

One significant obstacle to improving state reconstructions in ensemble-based filtering 848 methods is the inherent limited number of ensemble members that can be computationally 849 afforded. Localization methods, as discussed by Sakov and Bertino (2010), help address this 850 limitation. Localization is introduced either by localizing the observation in the assimilation 851 process or by localizing the covariance matrix. This second approach seems more difficult 852 to implement with the RKHS EnF, which operates directly in the weight space. The first 853 approach, which result in local filter is more appealing and doable with the introduction of localized observable and considering spatial vector of stationary weights. In the same spirit 855 Bocquet (2016) introduces a localization strategy that reconstructs trajectories through 856 linear combinations, where the localization operator is propagated over time. This class of 857 methods could potentially be naturally adapted within the time-evolving RKHS framework 858 to introduce dynamical localization process. 859

In addition let us mention that, applying a tiled filter with local analysis could mitigate imbalances from localized approximate reconstructions.

Another perspective is that the framework allows direct access to the system's Lyapunov times (Dufée et al., 2024), which indicate when the predictability horizon has been exceeded. This information is valuable for determining the optimal moments for assimilation, facilitating the renewal of the ensemble and creating an adaptive assimilation step filter.

Additionally, the initial ensemble significantly impacts the method's performance, making it crucial to select a well-chosen starting set. Generative methods can be employed to establish an initial ensemble that best represents the reference simulation's initial conditions, thereby enhancing the filtering method's effectiveness.

In this study, the reference dynamics were modeled using the same model and resolution as the training set elements. However, modeling a reference that differs from the training set ensemble poses challenges. A different resolution can lead to variations in energy levels between the reference and training ensemble members; for instance, higher-resolution reference models generally exhibit more energy than their coarser counterparts. To address this, variability can be introduced into the training ensemble through methods such as inflation or by employing a stochastic model with calibrated noise (Dufée et al., 2022).

Moreover, to model the reference differently from the training set, the quasi-geostrophic 878 projection can be omitted in the reference model run. This approach would enable the ref-879 erence to capture fast waves, which would not be reconstructed by the training ensemble. 880 Incorporating multi-modal data could further enhance the assimilation process. For exam-881 ple, combining sea surface height data with sea surface temperature and integrating data 882 from the Surface Water Ocean Topography (SWOT) mission alongside traditional altimeter 883 data could improve the assimilation of real data, as discussed by (Le Guillou, Lahaye, et al., 884 2021).885

886 Open Research Section

The data and the code files necessary to reproduce the results presented in this paper are publicly available at: https://gitlab.inria.fr/mael.jaouen/RKHS_EnF/

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nicus Marine Service Information. 895

References 896

897

898

- data in gaussian process regression. with application to aerodynamics. Journal of Computational Physics, 470, 111595. 800 Aronszajn, N. (1950). Theory of reproducing kernels. Transactions of the American math-900 ematical society, 68(3), 337-404. 901 Auroux, D., & Blum, J. (2008). A nudging-based data assimilation method: the back and 902 forth nudging (bfn) algorithm. Nonlinear Processes in Geophysics, 15(2), 305–319. 903 Bauer, W., Chandramouli, P., Chapron, B., Li, L., & Mémin, E. (2020). Deciphering 904 the role of small-scale inhomogeneity on geophysical flow structuration: A stochastic 905 approach. Journal of Physical Oceanography, 50(4), 983 - 1003. 906 Bishop, A. N., & Del Moral, P. (2023). On the mathematical theory of ensemble (linear-907 gaussian) Kalman-bucy filtering. Mathematics of Control, Signals, and Systems, 908 35(4), 835-903.909 Bishop, C. H., Etherton, B. J., & Majumdar, S. J. (2001). Adaptive sampling with the en-910 semble transform Kalman filter. part I: Theoretical aspects. Monthly Weather Review, 911 129(3), 420 - 436.912 Bocquet, M. (2016). Localization and the iterative ensemble Kalman smoother. Quarterly 913 Journal of the Royal Meteorological Society, 142(695), 1075-1089. 914 Bocquet, M., & Sakov, P. (2014, July). An iterative ensemble Kalman smoother. Quarterly 915 Journal of the Royal Meteorological Society, 140(682), 1521-1535. 916 Brajard, J., Carrassi, A., Bocquet, M., & Bertino, L. (2020). Combining data assimilation 917 and machine learning to emulate a dynamical model from sparse and noisy observa-918 tions: A case study with the lorenz 96 model. Journal of Computational Science, 44, 919 101171. 920 Brecht, R., Li, L., Bauer, W., & Mémin, E. (2021). Rotating shallow water flow under 921 location uncertainty with a structure-preserving discretization. Journal of Advances 922 in Modeling Earth Systems, 13(12), e2021MS002492. 923 Carrassi, A., Bocquet, M., Bertino, L., & Evensen, G. (2018). Data assimilation in the geo-924 sciences: an overview of methods, issues, and perspectives. WIREs Climate Change, 925 9. 926 Cheng, S., Quilodrán-Casas, C., Ouala, S., Farchi, A., Liu, C., Tandeo, P., ... Arcucci, 927 R. (2023). Machine learning with data assimilation and uncertainty quantification 928 for dynamical systems: A review. IEEE/CAA Journal of Automatica Sinica, 10(6), 929 1361-1387. 930 Desroziers, G., Camino, J.-T., & Berre, L. (2014). 4DEnVar: link with 4D state formulation 931 of variational assimilation and different possible implementations. Quarterly Journal 932 of the Royal Meteorological Society, 140(684), 2097-2110. 933 Doucet, A., Freitas, N. D., & Gordon, N. (2001). Sequential Monte Carlo methods in 934 practice. Springer Verlag. 935 Dufée, B., Hug, B., Mémin, E., & Tissot, G. (2024). Ensemble forecasts in reproducing 936 kernel Hilbert space family. Physica D: Nonlinear Phenomena, 459, 134044. 937 Dufée, B., Mémin, E., & Crisan, D. (2022). Stochastic parametrization: An alternative to 938 939 inflation in ensemble Kalman filters. Quarterly Journal of the Royal Meteorological Society, 148(744), 1075-1091. 940 Durran, D., Weyn, J. A., & Menchaca, M. Q. (2017). Practical considerations for computing 941
- dimensional spectra from gridded data. Monthly Weather Review, 145(9), 3901 3910. 942
- Evensen, G. (2009). Data assimilation, the ensemble Kalman filter. Springer Berlin, Hei-943 delberg. 944

Fablet, R., Chapron, B., Drumetz, L., Mémin, E., Pannekoucke, O., & Rousseau, F. (2021). 945 Learning variational data assimilation models and solvers. Journal of Advances in 946 Modeling Earth Systems, 13(10), e2021MS002572. (e2021MS002572 2021MS002572) 947 Fablet, R., Chapron, B., Le Sommer, J., & Sévellec, F. (2024). Inversion of sea surface 948 currents from satellite-derived sst-ssh synergies with 4DVarNets. Journal of Advances 949 in Modeling Earth Systems, 16(6), e2023MS003609. (e2023MS003609 2023MS003609) 950 Farchi, A., Laloyaux, P., Bonavita, M., & Bocquet, M. (2021). Using machine learning to 951 correct model error in data assimilation and forecast applications. Quarterly Journal 952 of the Royal Meteorological Society, 147(739), 3067-3084. 953 Friedman, J. H. (1987). Exploratory projection pursuit. Journal of the American Statistical 954 Association, 82(397), 249–266. 955 Gordon, N., Salmond, D., & Smith, A. (1993, April). Novel approach to nonlinear/non-956 gaussian bayesian state estimation. IEE Proceedings F (Radar and Signal Processing), 957 140, 107-113(6).958 Gottwald, G. A., & Reich, S. (2021). Supervised learning from noisy observations: Combin-959 ing machine-learning techniques with data assimilation. Physica D: Nonlinear Phe-960 nomena, 423, 132911. 961 Hoke, J. E., & Anthes, R. A. (1976). The initialization of numerical models by a dynamic-962 initialization technique. Monthly Weather Review, 104(12), 1551 - 1556. 963 Hunt, B. R., Kalnay, E., Kostelich, E., Ott, E., Patil, D., Sauer, T., ... Zimin, A. (2004). 964 Four-dimensional ensemble Kalman filtering. Tellus A: Dynamic Meteorology and 965 Oceanography, 56(4), 273-277.966 Hunt, B. R., Kostelich, E. J., & Szunyogh, I. (2007). Efficient data assimilation for spa-967 tiotemporal chaos: A local ensemble transform Kalman filter. Physica D: Nonlinear 968 Phenomena, 230(1), 112-126. (Data Assimilation) 969 Koivunen, A. C., & Kostinski, A. B. (1999). The feasibility of data whitening to improve 970 performance of weather radar. Journal of Applied Meteorology, 38(6), 741 - 749. 971 Le Gland, F., Monbet, V., & Tran, V. (2011). Large sample asymptotics for the ensemble 972 Kalman filter. In D. Crisan & B. Rozovskii (Eds.), Handbook on nonlinear filtering. 973 Oxford University Press. 974 Le Dimet, F.-X., & Talagrand, O. (1986). Variational algorithms for analysis and assimila-975 tion of meteorological observations: theoretical aspects. Tellus A, 38A(2), 97-110. 976 Leeuwenburgh, O., Evensen, G., & Bertino, L. (2005). The impact of ensemble filter 977 definition on the assimilation of temperature profiles in the tropical pacific. Quarterly 978 Journal of the Royal Meteorological Society, 131(613), 3291-3300. 979 Le Guillou, F., Lahaye, N., Ubelmann, C., Metref, S., Cosme, E., Ponte, A., ... Vidard, A. 980 (2021). Joint estimation of balanced motions and internal tides from future wide-swath 981 altimetry. Journal of Advances in Modeling Earth Systems, 13(12), e2021MS002613. 982 (e2021MS002613 2021MS002613) 983 Le Guillou, F., Metref, S., Cosme, E., Ubelmann, C., Ballarotta, M., Le Sommer, J., & 984 Verron, J. (2021). Mapping altimetry in the forthcoming swot era by back-and-forth 985 nudging a one-layer quasigeostrophic model. Journal of Atmospheric and Oceanic 986 Technology, 38(4), 697 - 710.987 Le Traon, P. Y., Faugère, Y., Hernandez, F., Dorandeu, J., Mertz, F., & Ablain, M. (2003). 988 Can we merge GEOSAT follow-on with TOPEX/Poseidon and ERS-2 for an improved 989 description of the ocean circulation? Journal of Atmospheric and Oceanic Technology, 990 20(6).991 Liu, C., Xiao, Q., & Wang, B. (2008). An ensemble-based four-dimensional variational 992 data assimilation scheme. part I: Technical formulation and preliminary test. Monthly 993 Weather Review, 136(9), 3363 - 3373. 994 Lorenc, A. C. (2003). The potential of the ensemble Kalman filter for NWP—a comparison 995 with 4D-Var. Quarterly Journal of the Royal Meteorological Society, 129(595), 3183-996 3203. 997 Luo, X. (2019, July). Ensemble-based kernel learning for a class of data assimilation 998 problems with imperfect forward simulators. PLOS ONE, 14(7), 1–40. 999

- Manucharyan, G. E., Siegelman, L., & Klein, P. (2021). A deep learning approach to
 spatiotemporal sea surface height interpolation and estimation of deep currents in
 geostrophic ocean turbulence. Journal of Advances in Modeling Earth Systems, 13(1),
 e2019MS001965. (e2019MS001965 2019MS001965)
- Mauran, S., Mouysset, S., Simon, E., & Bertino, L. (2023). A kernel extension of the ensemble transform Kalman filter. In J. Mikyška, C. de Mulatier, M. Paszynski,
 V. V. Krzhizhanovskaya, J. J. Dongarra, & P. M. Sloot (Eds.), Computational science - iccs 2023 (pp. 438–452). Cham: Springer Nature Switzerland.
- Ouala, S., Fablet, R., Herzet, C., Chapron, B., Pascual, A., Collard, F., & Gaultier, L.
 (2018). Neural network based Kalman filters for the spatio-temporal interpolation of satellite-derived sea surface temperature. *Remote Sensing*, 10(12).
- Reich, S., & Cotter, C. (2015). Probabilistic forecasting and bayesian data assimilation.
 Cambridge University Press.
 - Roullet, G., & Gaillard, T. (2022). A fast monotone discretization of the rotating shallow water equations. *Journal of Advances in Modeling Earth Systems*, 14(2), e2021MS002663. (e2021MS002663 2021MS002663)
 - Sakov, P., & Bertino, L. (2010). Relation between two common localisation methods for the enkf comput. *Computational Geosciences*, 15, 225-237.
 - Stein, M. L. (1999). Interpolation of spatial data. Springer New York.

1013

1014

1015

1016

1017

1018

1022

1023

1024

1025

1026

- 1019Taburet, G., Sanchez-Roman, A., Ballarotta, M., Pujol, M.-I., Legeais, J.-F., Fournier, F.,1020... Dibarboure, G. (2019). Duacs dt2018: 25 years of reprocessed sea level altimetry1021products. Ocean Science, 15(5), 1207–1224.
 - Thiry, L., Li, L., Mémin, E., & Roullet, G. (2024). A unified formulation of quasi-geostrophic and shallow water equations via projection. *Journal of Advances in Modeling Earth* Systems, 16(10), e2024MS004510.
 - Thiry, L., Li, L., Roullet, G., & Mémin, E. (2023). Mqgeometry-1.0: a multi-layer quasigeostrophic solver on non-rectangular geometries. *EGUsphere*, 2023, 1–25.
- Ubelmann, C., Dibarboure, G., Gaultier, L., Ponte, A., Ardhuin, F., Ballarotta, M., &
 Faugère, Y. (2021). Reconstructing ocean surface current combining altimetry and
 future spaceborne doppler data. *Journal of Geophysical Research: Oceans*, 126(3),
 e2020JC016560. (e2020JC016560 2020JC016560)
- Ubelmann, C., Klein, P., & Fu, L.-L. (2015). Dynamic interpolation of sea surface height
 and potential applications for future high-resolution altimetry mapping. Journal of
 Atmospheric and Oceanic Technology, 32(1).
- Van Leeuwen, P. J., Künsch, H. R., Nerger, L., Potthast, R., & Reich, S. (2019). Particle
 filters for high-dimensional geoscience applications: A review. *Quarterly Journal of the Royal Meteorological Society*, 145(723), 2335-2365.
- Wan, L., Bertino, L., & Zhu, J. (2010). Assimilating altimetry data into a hycom model of
 the pacific: Ensemble optimal interpolation versus ensemble Kalman filter. Journal of
 Atmospheric and Oceanic Technology, 27(4), 753 765.
- ¹⁰⁴⁰ Whitaker, J. S., & Hamill, T. M. (2002). Ensemble data assimilation without perturbed ¹⁰⁴¹ observations. *Monthly Weather Review*, 130(7), 1913 - 1924.
- Yang, Y., Robinson, C., Heitz, D., & Mémin, E. (2015). Enhanced ensemble-based 4DVar
 scheme for data assimilation. *Computers & Fluids*, 115, 201-210.
- Zhang, Z., Archibald, R., & Bao, F. (2022). A pde-based adaptive kernel method for solving
 optimal filtering problems. Journal of Machine Learning for Modeling and Computing,
 3(3).