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A low dispersion finite volume/element method for nonlinear wave propagation, breaking and runup on unstructured meshes

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Abstract

We study a hybrid approach combining a FV and FE method to solve a fully nonlinear and weakly-dispersive depth averaged wave propagation model. The FV method is used to solve the underlying hyperbolic shallow water system, while a standard P^1 finite element method is used to solve the elliptic system associated to the dispersive correction. We study the impact of several numerical aspects: the impact of the reconstruction used in the hyperbolic phase; the representation of the FV data in the FE method used in the elliptic phase and their impact on the theoretical accuracy of the method; the well-posedness of the overall method. For the first element we proposed a systematic implementation of an iterative reconstruction providing on arbitrary meshes up to third order solutions, full second order first derivatives, as well as a consistent approximation of the second derivatives. These properties are exploited to improve the assembly of the elliptic solver, showing dramatic improvement of the finale accuracy, if the FV representation is correctly accounted for. Concerning the elliptic step, the original problem is usually better suited for an approximation in $H(\text{div})$ spaces. However, it has been shown that perturbed problems involving similar operators with a small Laplace perturbation are well behaved in H^1 . We show, based on both heuristic and strong numerical evidence, that numerical dissipation plays a major role in stabilizing the coupled method, and not only providing convergent results, but also providing the expected convergence rates. Finally, the full mode, coupling a wave breaking closure previously developed by the authors, is thoroughly tested on standard benchmarks using unstructured grids with sizes comparable or coarser than those usually proposed in literature.

Keywords: Green-Naghdi equations, hybrid scheme, Finite Volumes, Finite Elements, high order, wave breaking

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Contents

1	Introduction	3
2	The fully-nonlinear/weakly-dispersive model	4
3	Solution strategy and geometrical notation	7
4	Hyperbolic step: third order FV scheme and derivatives recovery via successive corrections	8
4.1	Polynomial expansion and derivative reconstruction via successive corrections	10
4.1.1	First derivatives	10
4.1.2	Second derivatives and second order corrected gradients	11
4.1.3	Capturing of non-smooth solutions and limiting	13
4.2	Numerical verification for smooth and non-smooth flows	13
5	Finite element solver for dispersive effects	15
6	Finite element/volume coupling: consistency and well-posedness considerations	16
6.1	Consistency: using FV data in the FE solver and vice-versa	17
6.2	A comment on well-posedness	19
7	Dispersion error: time continuous spectral analysis	22
8	Boundary conditions, wave generation, and wave breaking closure	24
8.1	Wave breaking detection and closure	24
8.2	Wave generation and boundary conditions	25
9	Numerical validation and benchmarking	26
9.1	Grid convergence for the solitary wave	26
9.2	Circular shoal	26
9.3	Elliptic shoal	28
9.4	Solitary interacting with a cylinder	31
9.5	Solitary wave breaking on a 3D reef	32
9.6	The seaside experiment	36

10 Conclusions	37
Appendix A	38
Appendix B	39

1. Introduction

Accurate simulations of water wave's propagation and non-linear wave transformations is of fundamental importance to marine and coastal engineering. Over the last decades, significant efforts in the development of depth averaged models have been made in order to provide the means of accurately predicting near-shore wave processes.

One of the most applied depth averaged models is the Non-linear Shallow Water Equations (NSWE). They are applied to describe long wave hydrodynamics when the vertical acceleration of water particles can be neglected assuming the flow to be nearly horizontal. They are able to model important aspects of wave propagation phenomena, the general characteristics of the run-up process, and the wave breaking with broken waves represented as shocks. However they are not appropriate for deeper waters and shoaling since they neglect all the dispersive effects that play a very important role. Taking dispersive effects in to account is of critical importance if we want to study the nearshore wave propagation and transformation. The main tool for performing studies including also dispersion have been pioneered by Boussinesq [11], who derived a system of equations under the assumption that non-linearity and dispersion are weak and in the same order of magnitude. Peregrine [56] was the first to derive a Boussinesq-type (BT) system of equations with topography terms. During the 1990s researchers focused on improving the dispersive properties of the original model of Peregrine and push the range of validity of the equations towards deeper waters, leading mainly to BT models restricted to situations with weakly non-linear interactions. Some famous models among them are [51, 45, 6]. However in many practical applications the effects of the non-linearity are too large to be treated using weakly non-linear BT models. Green and Naghdi [30] derived a fully non-linear weakly dispersive model (GN model) which gained a lot of attention the last two decades. The range of validity of this last model requires only the dispersion parameter to be small, but it does not impose any restriction to the non-linearity. However, linear dispersion properties of the GN model are the same as those of Peregrine. In [16] and [10] an enhanced model with an improved dispersion relationship is proposed. This model is used in our work.

For the numerical discretization of the GN equations in 1DH all the common techniques like Finite Elements (FE), Finite Volumes(FV) and Finite Differences (FD) have been used, see for example [29, 16, 10, 50] and references therein. In less studied two horizontal dimensions (2D) cases, and only in Cartesian meshes, the numerical techniques that have been used are again FD [22, 73, 75], combinations of FV and FD methods [58, 66, 41] and Discontinuous Galerkin (DG) methods [42]. Up to the authors knowledge, the only works on general unstructured meshes are [24], where a new form of the GN equations is solved using a DG method, and the very recent work of [46], where a combined Hybridizable Discontinuous Galerkin and Runge-Kutta DG formulation is used for the same set of equations. In this work we extended the idea presented by the authors in [29] and [27]. We solve the enhanced GN equations, first introduced in [40], using a flexible combination of FV and the standard C^0 Galerkin FE method on unstructured meshes.

The manuscript is organized as follows. In section 2 we present the model equations while in section 3 we describe the solution strategy which we follow in this work. Section 4 presents a higher order solver for the hyperbolic part of the equations, i.e the shallow water equations and the next section shows how to incorporate the elliptic part in to the system using the FE technique. A discussion on the well-posedness of the coupled method is following while section 7 presents a time-continues dispersion error analysis. Section 8 is devoted to implementation details and finally in section 9 the numerical results justify and verify our choices. The paper is concluded by a discussion and outlook on future work.

2. The fully-nonlinear/weakly-dispersive model

In this work we use the enhanced GN (eGN) system of equations in the form proposed in [9]. This formulation provides and order $O(\mu)$ depth averaged approximation of the nonlinear wave equations, being μ the dispersion parameter defined as $\mu = h_0^2/\lambda^2$, where h_0 is the reference water depth and λ the wavelength. The two dimensional form of the system can be written in the following form:

$$\begin{aligned} h_t + \nabla \cdot (h\mathbf{u}) &= 0 \\ (I + \alpha T) \left(\mathbf{q}_t + \nabla \cdot \left(\frac{\mathbf{q} \otimes \mathbf{q}}{h} \right) + gh\nabla\eta \right) + \frac{g}{\alpha} h\nabla\eta + hQ(\mathbf{u}) &= 0 \end{aligned} \tag{1}$$

where the operators $T(\cdot)$ and $Q(\cdot)$ are:

$$T(\cdot) = -\frac{1}{3}\nabla\left(h^3\nabla\cdot\frac{(\cdot)}{h}\right) - \frac{h^2}{2}\left(\nabla\cdot\frac{(\cdot)}{h}\right)\nabla b + \frac{1}{2}\nabla\left(h^2\nabla b\cdot\frac{(\cdot)}{h}\right) + h\left(\nabla b\cdot\frac{(\cdot)}{h}\right)\nabla b, \quad (2)$$

$$\begin{aligned} Q(\cdot) &= \frac{2}{3h}\nabla\left(h^3\left(\nabla(\cdot)_1\cdot\nabla^\perp(\cdot)_2 + (\nabla\cdot(\cdot))^2\right)\right) + h^2\left(\nabla(\cdot)_1\cdot\nabla^\perp(\cdot)_2 + (\nabla\cdot(\cdot))^2\right)\nabla b + \\ &+ \frac{1}{2h}\nabla\left(h^2\left((\cdot)\cdot((\cdot)\cdot\nabla)\nabla b\right)\right) + ((\cdot)\cdot((\cdot)\cdot\nabla)\nabla b)\nabla b. \end{aligned} \quad (3)$$

We denote $h(\mathbf{x}, t) = h_0 + \eta(\mathbf{x}, t) - b(\mathbf{x})$ the total water depth, where $\eta(\mathbf{x}, t)$ the free surface elevation with respect to the water rest state h_0 , $b(\mathbf{x})$ the topography variation and $\mathbf{u}(\mathbf{x}, t) = (u, v)$ the flow velocity as shown in figure 1. $(\cdot)_1$ and $(\cdot)_2$ indicates respectively the first and second component of the vector (\cdot) and ∇^\perp states for the normal gradient operator.

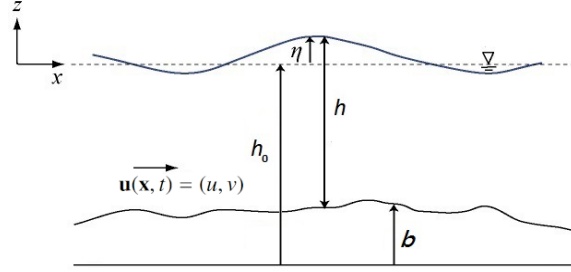


Figure 1: Description of the free surface flow problem and main notation.

The operator $T(\cdot)$ plays a key role, as its inversion is necessary to be able to obtain evolution equations for the physical variables. $T(\cdot)$ can be written in compact form involving two operators $S_1(\cdot)$, $S_2(\cdot)$ and their adjoints $S_1^*(\cdot)$, $S_2^*(\cdot)$, as:

$$T(\cdot) = S_1^*\left(hS_1\left(\frac{(\cdot)}{h}\right)\right) + S_2^*\left(hS_2\left(\frac{(\cdot)}{h}\right)\right) \quad (4)$$

where

$$S_1(\cdot) = \frac{h}{\sqrt{3}}\nabla\cdot(\cdot) - \frac{\sqrt{3}}{2}\nabla b\cdot(\cdot), \quad S_2 = \frac{1}{2}\nabla b\cdot(\cdot). \quad (5)$$

Note that this formulation is essential to show the coercivity of the operator $(I + \alpha T)$, see [29] and referenced therein for further details. Note also that when $\alpha = 1$ we retrieve the original GN equations.

The linear dispersion and shoaling properties of the fully nonlinear GN and eGN models are comparable to those of the weakly nonlinear models of Peregrine and Madsen and Sorensen, extensively described in

[27]. On the other hand, the nonlinear shoaling properties of a weakly/strongly nonlinear model is not easy to be examined analytically. As discussed in [27], one way to test the nonlinear shoaling properties of a model is by performing the test of Grilli et al. [31]. The test consists of a solitary wave with relative amplitude $\alpha/h_0 = 0.2m$, propagating on a water depth of $0.44m$ and shoaling on to a constant slope of $1 : 35$. Ten wave gauges have been placed along the flume to measure the free surface elevation. All of them are placed before the breaking point with the last one being the closest to the breaking point. Figure 2 compares the experimental wave's envelope with the result performed by four weakly non-linear models: Peregrine [56], Abbott [1], Madsen and Sorensen [45] and the MSP system. The last one is a modified system of Peregrine's equations written in a wave amplitude-velocity form, see [27, 28] for further details. In this work, we performed the same test on the fully nonlinear GN and eGN models and we added the computed results on the figure. Our result has been obtained using the discretization method presented in this work and it is a grid convergent solution, such that the plotted curve can be seen as genuine representations of the behaviour of the model. We can observe that as soon as the nonlinear effects dominate (this happens close the breaking region) the phenomenon is better reproduced by the fully nonlinear models and even better by the eGN equations used in this work.

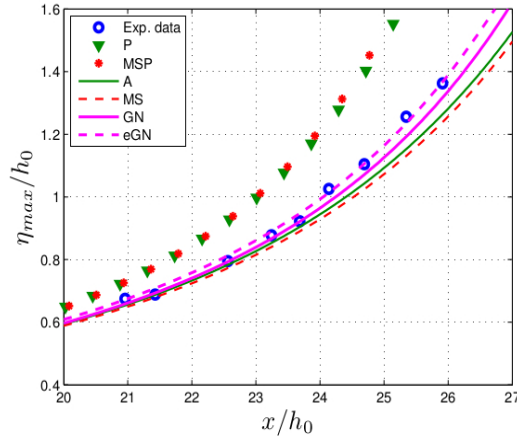


Figure 2: Wave envelope on the 9 gauges along the domain. Comparison between different weakly nonlinear models, GN, eGN and the experimental data.

3. Solution strategy and geometrical notation

To numerically solve (1), we rewrite the system of two dimensional enhanced GN equations as:

$$h_t + \nabla \cdot \mathbf{q} = 0, \quad (6)$$

$$\mathbf{q}_t + \nabla \cdot \left(\frac{\mathbf{q} \otimes \mathbf{q}}{h} \right) + gh \nabla \eta = \mathbf{\Phi} \quad (7)$$

$$(I + \alpha T) \mathbf{\Psi} = W - R, \quad (8)$$

$$\mathbf{\Phi} = \mathbf{\Psi} + \frac{gh}{\alpha} \nabla \eta \quad (9)$$

by splitting the original system in its elliptic and hyperbolic parts, through the definition of the new variable $\mathbf{\Phi} = [\phi_x, \phi_y]^T$. $\mathbf{\Phi}$ accounts for the dispersive effects and has the role of a non-hydrostatic pressure gradient in the Shallow water equations. We define

$$\begin{aligned} W &= -\frac{gh}{\alpha} \nabla \eta \\ R &= hQ\left(\frac{\mathbf{q}}{h}\right) \end{aligned} \quad (10)$$

where the operators $T(\cdot)$ and $Q(\cdot)$ are (2) and (3) respectively. In this work we solve (6)-(9) using a hybrid Finite Element (FE)- Finite volume (FV) scheme where the elliptic part of the system is discretized by means of the continuous Galerkin FE method. The hyperbolic part of the system is discretized by the two dimensional formulation of the finite volume scheme inspired by the works [37, 38]. We refer to the work [37], which has been proven to be a robust scheme, capable of simulating wave transformations providing accurate results in complex scenarios and over two dimensional unstructured triangular meshes.

So we consider a triangulation of the spatial domain which we denote by Ω_h , with the roman h denoting the largest element diameter. In the approach developed here, we will both make use of elements defined by each of the non-overlapping triangles of the mesh, as well as of a *median-dual* partition in order to generate non-overlapping nodal control volumes. Let us denote by K the generic triangular element, and by K_i the set of elements sharing node i . We then denote by C_i the median dual cell obtained by joining the gravity centers of the triangles in K_i with the midpoints of the edges meeting in i . Simple geometry shows that $|C_i| = \sum_{K \in K_i} \frac{|K|}{3}$. We also define D_i as the set of nodes connected to i . For any $j \in D_i$, the shared portion of boundary of C_i and C_j is named ∂C_{ij} , and it is composed by the union of two segments connecting the barycenters of the two triangles sharing the edge ij with the edge midpoint (see figure 3). The boundary of the median dual cell of i can thus be defined as: $\partial C_i = \sum_{j \in K_i} \partial C_{ij}$. Moreover, we define \mathbf{r}_{ij} the vector

connecting nodes i and j . Note finally, that the intersection of C_i intersects each element $K \in K_i$ can be split into two half cells associated to the two edges stemming from i . The half cell containing node j is denoted by C_{ij}^K , and we set $C_{ij} = \bigcup_{K \in K_i \cap K_j} C_{ij}^K$ so that $C_i = \bigcup_{j \in D_i} C_{ij}$.

4. Hyperbolic step: third order FV scheme and derivatives recovery via successive corrections

For simplicity we rewrite the system of conservation laws (6)-(7) as

$$\mathbf{U}_t + \nabla \cdot \mathbf{F}(\mathbf{U}) = \mathbf{S}_b + \Phi_i. \quad (11)$$

with $\mathbf{U} = [h, \mathbf{q}]^T$, $\mathbf{F} = [\mathbf{q}, \mathbf{q} \otimes \mathbf{q}/h + gh^2 \mathbf{I}_2]^T$, with \mathbf{I}_2 the rank 2 identity matrix, and with $\mathbf{S}_b = -[0, gh\nabla b]^T$. The FV integration over each computational cell C_i leads to the semi-discrete form of the scheme as:

$$\frac{\partial \bar{\mathbf{U}}_i}{\partial t} + \frac{1}{|C_i|} \sum_{j \in D_i} \int_{\partial C_{ij}} \hat{\mathbf{F}} \cdot \mathbf{n} = \frac{1}{|C_i|} \sum_{j \in D_i} \int_{C_{ij}} \mathbf{S}_b + \bar{\Phi}_i, \quad (12)$$

where $\bar{\mathbf{U}}_i$ is the volume averaged value of \mathbf{U} over C_i , \mathbf{n} is the unitary outward vector normal to ∂C_i , and with $\bar{\Phi}_i = \int_{C_i} \Phi$ evaluated using numerical quadrature (cf. section §5). In the above expression, $\hat{\mathbf{F}}$ is the numerical flux defined here using the approximate Riemann solver of [63]. The method used here is relatively standard and we will not provide much details. It is based on a well balanced formulation of the integrals of the fluxes and of the bathymetry source, as well as a robust modification of the reconstruction and numerical flux to cope with the wet/dry transition. We refer the interested reader to e.g. [8, 14, 15], and to [37, 39] for some details on our implementation.

To reach high-order spatial accuracy, we have to reconstruct each component of the physical variables and bed topography. Following the classical strategy by [72] (cf also [66, 37] and [29]), to reduce the introduction of spurious numerical dispersion we avoid second order approximations for the hyperbolic terms, and look into the design of a fully third order method. This is achieved in this paper by means of a successive correction method which iteratively improves derivatives computed by means of the standard Green-Gauss formula. This allows to construct k-exact polynomials with all local operations, requiring only the exchange of information between adjacent cells. In particular for a third order method we need a quadratic polynomial requiring the knowledge of gradient and Hessian of the variables in the dual cell. The standard Green-Gauss formula is unfortunately not well suited for general unstructured meshes on which it provides first order derivatives which are at most consistent (1st order accurate), and second derivatives which may event be inconsistent. There are many methods to overcome this: from the classical least square

method used in k-exact method by Barth in [4, 2, 3], to more recent ones used in [69, 70, 76] and [17]. The basic limitation most of the above methods is the computational cost, related to the need of solving a more or less large linear system, and the complexity of the implementation, related to the need of assembling and using an enlarged stencil.

We follow here the method first proposed by [12, 35, 34] and more recently in [57]. In the references the authors constructed a generalized hybridization of Green-Gauss and Least square methods, called quasi-Green method, which results in a first-order accurate gradient on unstructured meshes. A successive correction method allows the construction of consistent gradient and Hessian on unstructured meshes. The idea of the corrections is to impose exact consistency with the monomials of appropriate degree. A thorough discussion and the general derivation of the method can be found in [57, 48] to which we refer for details. All the above works are using cell centered methods. In our work we have extended the approach to node centered finite volumes. Very recently (and independently on this work) [65] and [64] also provided a similar re-formulation for the linear advection equation and of the incompressible Euler equations.

In our work, we develop a node centered successive correction method for the hyperbolic nonlinear shallow water system, and appropriately combine it with a slope limiter to handle bores and hydraulic jumps. To our knowledge this is the first time that this approach is used for a hyperbolic system and combined with a limiter. We recall hereafter the basic steps to obtain a third order reconstruction. Most of the formulas allowing the implementation are provided in an appendix.

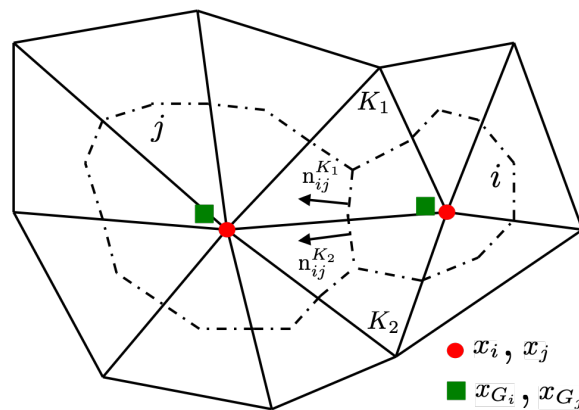


Figure 3: Notation and volume area used in the successive correction method

4.1. Polynomial expansion and derivative reconstruction via successive corrections

The reconstruction problem consists in defining a piece-wise polynomial of degree k that approximates $f(\mathbf{x})$ to the $(k + 1)^{\text{th}}$ order of accuracy. Our aim is to calculate approximations of the solution to the faces of the cells. To do this, we use high order polynomials obtained by Taylor expansions. Let us introduce the vector and tensor moments

$$\begin{aligned}\delta_{G_i}^{(1)}(\mathbf{x}) &= (\mathbf{x} - \mathbf{x}_{G_i}), \\ \delta_{G_i}^{(2)}(\mathbf{x}) &= \delta_{G_i}^{(1)} \otimes \delta_{G_i}^{(1)} = (\mathbf{x} - \mathbf{x}_{G_i}) \otimes (\mathbf{x} - \mathbf{x}_{G_i})\end{aligned}\tag{13}$$

For a third order scheme ($k = 2$) a conservative approximation is of the form [5, 52]

$$f_i(\mathbf{x}) = \bar{f}|_{G_i} + \mathbf{D}_f^{(1)}|_{G_i} \cdot (\mathbf{x} - \mathbf{x}_{G_i}) + \frac{1}{2} \mathbf{D}_f^{(2)}|_{G_i} : (\delta_{G_i}^{(2)}(\mathbf{x}) - M_i^{(2)}),\tag{14}$$

where where $\mathbf{D}_f^k|_{G_i}$ represents the order k spatial derivative of f (gradient, Hessian, etc) at the gravity center \mathbf{x}_{G_i} , and where the $A : B$ operator denotes the element by element lumped matrix product

$$A : B = \sum_{i,j} A_{ij} B_{ij}.$$

The matrix M_i contains the geometric moments:

$$M_i^{(2)} = \int_{C_i} \delta_{G_i}^{(2)}$$

Note that these geometric moments are mesh dependent quantities that can be pre-computed via numerical quadrature and stored (cf. Appendix A for mode details). The cell polynomials thus defined guarantee the conservation property

$$\frac{1}{|C_i|} \int_{C_i} f_i(\mathbf{x}) dS = \bar{f}|_{G_i}.$$

The crucial step is to computation of the spatial derivatives with the desired accuracy.

4.1.1. First derivatives

To get the first derivative at the gravity center of C_i , we apply the quasi-Green gradient approximation. This consists in a Green-Gauss reconstruction with a correction restoring the consistency of the operator on general meshes [48]. We end up with an operator approximating the gradient to first order accuracy on general meshes as (cf again figure 3 for the notation):

$$\mathbf{D}_f^{(1,\sigma 1)}|_{G_i} = M_1^{-1} \sum_{j \in D_i} \left[w_{ij}^{K_2} \bar{f}_i + (1 - w_{ij}^{K_1}) \bar{f}_j \right] \mathbf{n}_{ij}^{K_1}.\tag{15}$$

where the superscript $o1$ denotes that the approximation is first order accurate, and with the weights w_{ij} computed based on the relative distance of the cell center's \mathbf{x}_i to its face:

$$w_{ij}^{K_1} = \frac{\mathbf{s}_{ij}^{K_1} \cdot \mathbf{n}_{ij}^{K_1}}{\mathbf{r}_{ij}^{K_1} \cdot \mathbf{n}_{ij}^{K_1}}, \quad \mathbf{s}_{ij}^{K_1} = \frac{1}{2}\mathbf{x}_G^{K_1} + \frac{1}{4}(\mathbf{x}_i + \mathbf{x}_j) \quad (16)$$

and similarly for $w_{ij}^{K_2}$. The 2×2 matrix $(M_1)_i$ corrects the gradient to ensure its consistency. It is computed by imposing that for $f = x$ we have $\mathbf{D}_x^{(1,o1)}|_{G_i} = (1, 0)$, and similarly for $f = y$ we have $\mathbf{D}_y^{(1,o1)}|_{G_i} = (0, 1)$. This leads to:

$$M_{1i} = \left[\sum_{j \in D_i} w_{ij}^{K_1} (\mathbf{x}_{G_j} - \mathbf{x}_{G_i}) \otimes \mathbf{n}_{ij}^{K_1} + w_{ij}^{K_2} (\mathbf{x}_{G_j} - \mathbf{x}_{G_i}) \otimes \mathbf{n}_{ij}^{K_2} \right]^T \quad (17)$$

As the geometrical moments, this correction matrix is mesh dependent, but can be pre-computed and stored before the simulations. The first order gradient $\mathbf{D}_f^{(1,o1)}|_{G_i}$ allows to construct polynomials with second order of accuracy at most.

4.1.2. Second derivatives and second order corrected gradients

Once consistent first derivatives are available in all cells, we can proceed to a second iteration which will provide consistent second derivatives and improved gradients. As previously stated, a first order approximation of the second derivatives is enough to guarantee third order of accuracy for the overall polynomial (14). Unfortunately, unless the mesh presents special symmetries, applying (15) twice results in an inconsistent approximation of $\mathbf{D}_f^{(2)}|_{G_i}$. We denote hereafter this approximation as

$$\mathbf{D}_f^{(2,o0)}|_{G_i} = \mathbf{D}^{(1,o1)}(\mathbf{D}_f^{(1,o1)})|_{G_i}$$

The idea is to correct this quantity as done for the gradient

$$\mathbf{D}_f^{(2,o1)}|_{G_i} = M_{2i}^{-1} \mathbf{D}_f^{(2,o0)}|_{G_i} = M_{2i}^{-1} \mathbf{D}^{(1,o1)}(\mathbf{D}_f^{(1,o1)})|_{G_i} \quad (18)$$

As for the gradient, the correction matrix M_{2i} can be computed component by component by requiring the approximation to be consistent when applied to $\mathbf{x} \otimes \mathbf{x}$, so that for example $(\mathbf{D}_{x^2}^{(2,o1)}|_{G_i})_{11} = 2$, $(\mathbf{D}_{y^2}^{(2,o1)}|_{G_i})_{22} = 2$, $(\mathbf{D}_{yx}^{(2,o1)}|_{G_i})_{12} = 1$, etc. It can be shown [33] that M_{2i} can be obtained by a double application of the first derivative of first order derivative to $(\mathbf{x} - \mathbf{x}_{G_i}) \otimes (\mathbf{x} - \mathbf{x}_{G_i})$. For brevity we omit here the expressions obtained, which are reported in detail in appendix A.

The computation of a first order accurate second order derivative is not enough to achieve third order accuracy in the reconstruction. We also have to correct the approximation of the first derivative for it to be

at least second order. To obtain a correction strategy, we can compare the Taylor series development of the exact gradient with the one obtained using the available reconstructed derivatives :

$$\begin{aligned}\nabla f^{\text{exact}} &= \nabla f|_{G_i} + \nabla(\nabla f)|_{G_i} \cdot (\mathbf{x} - \mathbf{x}_{G_i}) + O(h^2) \\ &= \mathbf{D}_x^{(1),(o1)}|_{G_i} + O(h) + \mathbf{D}_f^{(2),(o1)}|_{G_i} \cdot (\mathbf{x} - \mathbf{x}_{G_i}) + O(h^2)\end{aligned}\tag{19}$$

The first order remainder on the second line, is due to the poor accuracy of the available gradient. For the gradient to be second order the second line should provide an exact answer for $f = (\mathbf{x} - \mathbf{x}_{G_i}) \otimes (\mathbf{x} - \mathbf{x}_{G_i})$. This is precisely the strategy suggested in [48, 57] to correct the gradient. So in practice we set

$$\mathbf{D}^{(1,o2)}|_{G_i} = \mathbf{D}^{(1,fo)}|_{G_i} + M_{l_i}^{o2} \mathbf{D}^{(2,o1)}|_{G_i}, \tag{20}$$

where $M_{l_i}^{o2}$ is obtained by requiring the errors in the second line of (19) to vanish when $f = (\mathbf{x} - \mathbf{x}_{G_i}) \otimes (\mathbf{x} - \mathbf{x}_{G_i})$. For our 2D case $M_{l_i}^{o2}$ is a 2×3 matrix of the form (full expressions available in appendix A):

$$M_1^{o2} = - \begin{bmatrix} (\alpha_i)_x - x_{G_i} & (\beta_i)_x & (\gamma_i)_x - y_{G_i} \\ (\alpha_i)_y & (\beta_i)_y - y_{G_i} & (\gamma_i)_y - x_{G_i} \end{bmatrix}.$$

Note that all of the above matrices are only involved in local operations (involving nearest neighbours), they can all be pre-computed and stored during a pre-processing step, and then used to update the gradients by simple matrix-vector multiplications. There is no need of solving multiple linear systems. These are the main advantages of this method.

Unfortunately, the correction matrices seem to have no theoretical property guaranteeing their invertibility. However in all the cases that we examined here and in the references using the same approach, no problem was ever observed even in quite irregular meshes. Another issue is how to preserve the accuracy near boundaries. In this paper we have been only concerned with two conditions: symmetry or periodic conditions along straight lines. In both cases we have used ghost cells. For the symmetry/slip wall conditions we have defined the ghost values on a locally mirrored mesh on which scalar quantities (depth and bathymetry) have been copied, and vectors rotated by 180° wrt normal. In this framework, third order of accuracy can be obtained easily only for straight boundaries for which two layers of elements are mirrored in order to have enough stencil to compute the successive correction derivatives. Periodic conditions are imposed simply by extending the connectivity of the mesh to include the correspondence of the periodic boundaries so that all the geometrical quantities, as well as the residuals account for periodicity.

4.1.3. Capturing of non-smooth solutions and limiting

In order to prevent oscillations from developing in the numerical solution we use the slope limiter proposed by Michalak and Ollivier-Gooch in [49], for higher order MUSCL numerical schemes on unstructured meshes using a cell centered fv scheme for the Euler equations. Following the spirit of the above work we write the limited form of the higher order reconstruction in the middle point M of an edge connecting the nodes i and j , as

$$f_i(\mathbf{x}_M) = \bar{f}|_{G_i} + Lim_M \left(\mathbf{D}_f^{(1)}|_{G_i} \cdot (\mathbf{x}_M - \mathbf{x}_{G_i}) + \frac{1}{2} \mathbf{D}_f^{(2)}|_{G_i} : \left(\delta_{G_i}^{(2)}(\mathbf{x}_M) - M_i^{(2)} \right) \right). \quad (21)$$

The design of the slope limiter requires three steps. First we have to find the minimum $(\delta f)_{min}^i$ and the maximum values $(\delta f)_{max}^i$ of the difference $\bar{f}|_{G_j} - \bar{f}|_{G_i}$ in the stencil formed by the cell i and all the direct neighbors j . Then we compute the unlimited reconstructed value $f(\mathbf{x})_M$ and finally we compute the maximum allowable value for Lim_i as:

$$Lim_M = \begin{cases} g((\delta f)_{min}^i) & \text{if } f(\mathbf{x})_M - \bar{f}_i > 0, \\ g((\delta f)_{max}^i) & \text{if } f(\mathbf{x})_M - \bar{f}_i < 0, \\ 1 & \text{if } f(\mathbf{x})_M - \bar{f}_i = 0 \end{cases} \quad (22)$$

where $g(x) = \frac{x^2 + 2x}{x^2 + x + 2}$.

4.2. Numerical verification for smooth and non-smooth flows

We provide here a quick verification of the hyperbolic step. First we consider the smooth travelling vortex solution proposed by [60], and widely used in literature to measure the accuracy of discretizations for the shallow water equations. Please refer to [60] for the precise definition of the test. We compute the solution at finite time as prescribed in the reference on a series of regular and irregular unstructured grids. The reference size of the coarsest mesh is $h = 0.107573$, which is reduced to half at each refinement step. The convergence of the depth error is reported on figure 4. Following the discussion in section 4 we performed the test in both structured and unstructured meshes confirming that the Green-Gauss reconstruction on unstructured meshes, spoils the convergence since is not able to produce consisted gradients. The picture confirms that the nominal accuracy is measured in practice in third order case when the derivatives are recovered via the successive correction approach. For the second order case the Green-Gauss is consisted with the gradient so the order of convergence is 2 interdependently of the mesh. The results are omitted for brevity.

We then consider the Monai valley benchmark [44], a classical test inspired by a flume experiment reproducing a scaled down version of the 1993 the Hokkaido-Nansei-Oki tsunami impact on the Monai

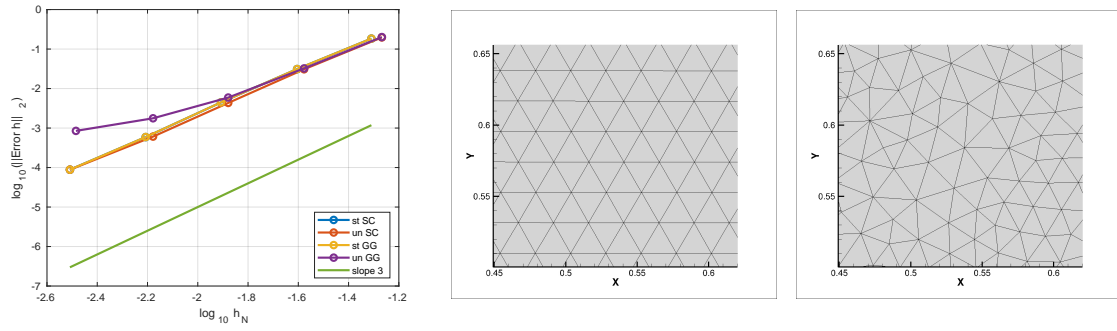


Figure 4: Error decay for the third order scheme. Representative structure and unstructure meshes.

valley. The test involves bore formation, propagation, and reflection, as well as runup. Following [59], we have run the experiment on an unstructured grid adapted to the bathymetry variations, and we show the wave patterns obtained at time $16.5s$ (see figure 5) with the second and third order scheme, as well as the water height times series in two of the gauges of the experiment (figure 6).

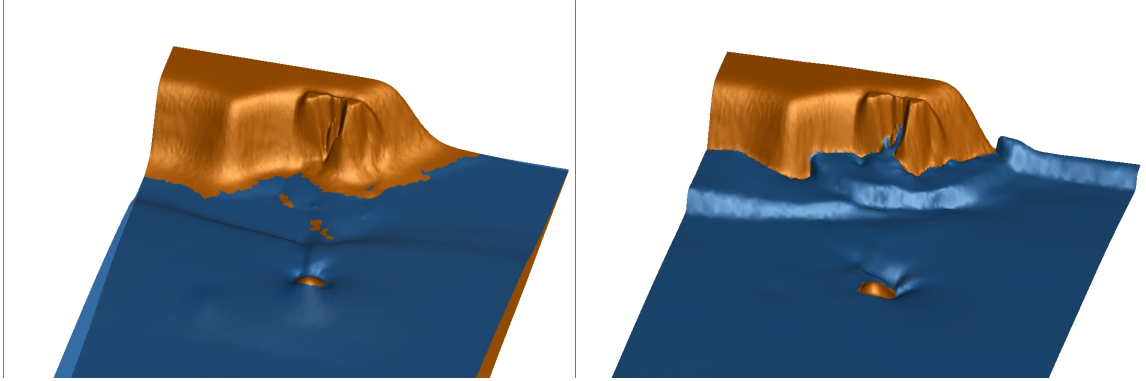


Figure 5: Monai valley: 3d view at time $t=14.5$ and $t=16.5s$ using the third order scheme.

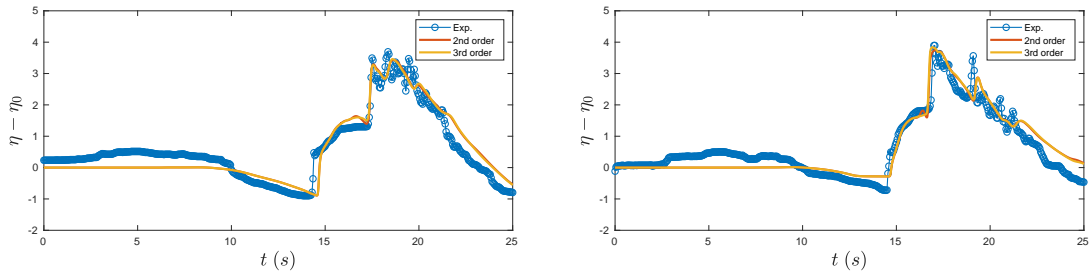


Figure 6: Monai valley: Free surface elevation at gauges 5 and 7.

5. Finite element solver for dispersive effects

Following the classical strategy [72], we now couple the non-dispersive hyperbolic solver to a second order solver for the physical weakly-dispersive effects. To this end, we propose to compute the auxiliary variable Ψ in (8) by means of a standard nodal P^1 (continuous) finite element (FE) method. Note that there is a notable change now in the meaning of the data associated to a mesh node i . The FV method evolves median dual cell averages which are in general different from the values of the variables at the nodes, used in the FE method. Similarly, we need to transfer from one representation to the other the derivatives of the depth and of the velocity appearing both in the coefficients in the operator T (equation (2) and equations (4)-(5)), and the right hand side where derivatives appear both in the definition of W and of R (cf. (10) and (3)). We will get back to this point in the next subsection.

The definitions of the operator T , $S_1(\cdot)$ and $S_2(\cdot)$, from (4) and (5) respectively, lead to the following variational form of the elliptic equation (8)

$$\int_{\Omega} \nu \cdot \Psi + \alpha \int_{\Omega} S_1(\nu) h S_1\left(\frac{\Psi}{h}\right) + \alpha \int_{\Omega} S_2(\nu) h S_2\left(\frac{\Psi}{h}\right) = \text{RHS} + \text{BCs} , \quad (23)$$

or equivalently using (5)

$$\int_{\Omega} \left\{ \frac{1}{3} \left(h \nabla \cdot \nu - \frac{3}{2} \nabla b \cdot \nu \right) \left(h^2 \nabla \cdot \left(\frac{\Psi}{h} \right) - \frac{3}{2} \nabla b \cdot \Psi \right) + \frac{1}{4} (\nabla b \cdot \nu) (\nabla b \cdot \Psi) \right\} = \text{RHS} + \text{BCs} , \quad (24)$$

with RHS a variational approximation of the term $W - R$, and BCs the boundary condition terms which we briefly discuss below. The value of Φ , required in the hyperbolic step, is recovered nodally from (9).

To obtain a fully discrete approximation of (24) we now consider the finite element approximation

$$\Psi_h = \sum_{K \in \Omega_h} \sum_{j \in K} \varphi_j \Psi_j \quad (25)$$

where $\text{span}\{\varphi_j\}_{j \in \Omega_h}$ is the classical continuous P^1 finite element space. We similarly introduce discrete approximations h_h , b_h , η_h , and \mathbf{u}_h for the elevations and velocity, as well as elemental discrete approximations of their first and second derivatives. Some options to provide these definitions and the solution we propose are discussed in the next subsection. The fully discrete variational form is expressed in terms of the array of the nodal values $\{\Psi_j\}_{j \in \Omega_h}$, which by abuse of notation we also label Ψ .

$$(\mathbb{M} + \alpha \mathbb{T}(h_h, b_h)) \Psi = \mathbb{W}(h_h, b_h) - \mathbb{R}(h_h, b_h, \mathbf{u}_h) \quad (26)$$

where the matrices on the left hand side are sparse 2×2 block matrices. In particular, \mathbb{M} is the mass matrix with entries

$$[\mathbb{M}]_{ij}^{mn} = \delta_{mn} \sum_{K \in K_i \cap K_j} \int_K \varphi_i \varphi_j, \quad (27)$$

while the entries of $\mathbb{T}(h_h, b_h)$ are evaluated using the relation $h^2 \nabla \cdot (\Psi/h) = h \nabla \cdot \Psi - \Psi \cdot \nabla h$ as

$$[\mathbb{T}(h_h, b_h)]_{ij}^{mn} = \sum_{K \in K_i \cap K_j} \int_K \left\{ \frac{1}{3} \left(h_h \partial_{X_m} \varphi_i - \frac{3}{2} \varphi_i (\partial_{X_m} b)_h \right) \left(h_h \partial_{X_n} \varphi_j - \frac{3}{2} \varphi_j (\partial_{X_n} b)_h - \varphi_j (\partial_{X_n} h)_h \right) + \frac{1}{4} \varphi_i (\partial_{X_m} b)_h \varphi_j (\partial_{X_n} b)_h \right\} \quad (28)$$

Note that in the above expression the mn indices run over the spatial components of the unknown, while ij run over the mesh nodes.

Finally the right hand side terms are defined as

$$\mathbb{W}(h_h, b_h) = -\frac{g}{\alpha} \sum_{K \in K_i} \int_K \varphi_i h_h (\nabla \eta)_h \quad (29)$$

and

$$\begin{aligned} \mathbb{R}(h_h, b_h, \mathbf{u}_h) &= \sum_{K \in K_i} \mathbb{R}^K \\ \mathbb{R}^K &= -\frac{2}{3} \int_{\Omega_h} \nabla \varphi_i h_h^3 \left((\nabla u)_h \cdot (\nabla^\perp v)_h + (\nabla \cdot \mathbf{u})_h^2 \right) + \int_{\Omega_h} \varphi_i h_h^2 \left((\nabla u)_h \cdot (\nabla^\perp v)_h + (\nabla \cdot \mathbf{u})_h^2 \right) (\nabla b)_h \\ &\quad - \frac{1}{2} \int_{\Omega_h} \nabla \varphi_i h_h^2 \left((\mathbf{D}_b^{(2)})_h : (\mathbf{u}_h \otimes \mathbf{u}_h) \right) + \int_{\Omega_h} \varphi_i h_h \left((\mathbf{D}_b^{(2)})_h : (\mathbf{u}_h \otimes \mathbf{u}_h) \right) (\nabla b)_h. \end{aligned} \quad (30)$$

having used the notation of section §4.1.1 for the Hessian of the bathymetry $\mathbf{D}_b^{(2)}$, and for its term by term product with the tensorized velocity.

Once the local polynomials representing h , b , \mathbf{u} and their derivatives are defined over the element, all the above formulas can be evaluated by means of a sufficiently accurate quadrature formula. In practice we have used here a 6 points symmetric formula exact for polynomials of degree 4 taken from [23]. This definition is the objective of the next section.

6. Finite element/volume coupling: consistency and well-posedness considerations

This section provides some additional constraints on some of the numerical choices possible with the method proposed. These are justified by means of some theoretical (albeit heuristic) arguments, as well as

by strong numerical evidence. We consider first the issue of ensuring a compatible data representation in the two phases of the computation. This will give an indication on how to use FV data in the FE solver (and vice-versa). We then provide a few comments on the well-posedness of the overall procedure which show the importance of using dissipative numerical fluxes.

6.1. Consistency: using FV data in the FE solver and vice-versa

We start by recalling that the two approaches being used to solve the equations are based on different representation of the data. The FV scheme evolves the solution averages \bar{U}_i over the dual cells C_i , and local polynomials within the cells are reconstructed using essentially all the neighbouring information. The FE method uses a collocated nodal representation, and within each element the polynomial variation is obtained by interpolating the data available at the nodes.

Although both methods used unknowns associated to nodes, their meaning is substantially different. More importantly, the approximation of the derivatives has an impact on the accuracy of the right hand sides of the elliptic problem. For this reason we have chosen here to proceed as follows:

- when passing the FV solution to the elliptic solver we sample the reconstructed polynomials and their derivatives at the nodes:

$$\begin{aligned} f_i &= \bar{f}|_{G_i} + \mathbf{D}_f^{(1)}|_{G_i} \cdot (\mathbf{x}_i - \mathbf{x}_{G_i}) + \frac{1}{2} \mathbf{D}_f^{(2)}|_{G_i} : (\delta_{G_i}^{(2)}(\mathbf{x}_i) - M_i^{(2)}). \\ (\nabla f)_i &= \mathbf{D}_f^{(1,o2)}|_{G_i} + \mathbf{D}_f^{(2,o1)}|_{G_i} \cdot (\mathbf{x}_i - \mathbf{x}_{G_i}) \\ (\mathbf{D}_f^{(2)})_i &= \mathbf{D}_f^{(2,o1)}|_{G_i} \end{aligned} \tag{31}$$

We then use these nodal values as a basis for a linear finite element approximation, so within any element $K \in \Omega_h$ we set:

$$f_h = \sum_{j \in K} \varphi_j f_j, \quad (\nabla f)_h = \sum_{j \in K} \varphi_j (\nabla f)_j, \quad (\mathbf{D}_f^{(2)})_h = \sum_{j \in K} \varphi_j (\mathbf{D}_f^{(2)})_j.$$

This, combined with the successive corrections method, allows to ensure on general meshes the second order of accuracy of all the first derivative terms (of h , b , and \mathbf{u}) appearing in (29), (30), and (28), and at least first order for the second derivatives of the bathymetry in (30). For configurations with high curvature in the topography, this is not enough and this value should be improved. A possible solution in relative simple cases, as those considered here, is to use the point-wise analytical value.

- The nodal finite element values Ψ_i are used to compute the post-processed average non-hydrostatic term

$$\overline{\Phi}_i := \int_{C_i} \Psi_h + \frac{g}{\alpha} \int_{C_i} h_h (\nabla \eta)_h$$

this formula are evaluated by splitting the integral in local contributions over the quadrangular shapes $C_i \cap K$, then further splitting the quadrangles in triangles by joining the node i to the gravity center of K , and then by using numerical quadrature on each sub-triangle.

Note that these choices have a direct impact on the theoretical accuracy attainable by the method. To see this, let us write an estimate on the local truncation error, for simplicity in the case of flat bathymetry.

We start by recasting (1) in dimensionless form. Using the standard fully nonlinear scaling leading to the Green-Naghdi system (see e.g. [40, 41] as well as [21] section §5) one easily shows that (6)-(9) can be written in dimensionless form as

$$\begin{aligned} \partial_t h + \nabla \cdot \mathbf{q} &= 0 \\ \partial_t \mathbf{q} + \nabla \cdot \left(\frac{\mathbf{q} \otimes \mathbf{q}}{h} \right) + h \nabla h &= \mu \left(\Psi + \frac{h}{\alpha} \nabla h \right) \\ (I + \mu \alpha T) \Psi &= -\frac{h}{\alpha} \nabla h + h Q(\mathbf{u}) \end{aligned} \quad (32)$$

where we recall that $\mu = h_0^2 / \lambda^2$ is the ratio of the reference depth on reference wavelength, and measures the shallowness and importance of dispersion. More importantly, as already recalled in the introduction, the above model is an approximation of the full non-linear free surface potential equations within an asymptotic error of $\mathcal{O}(\mu^2)$ [40]. We now consider an exact smooth solution, and combine (32) with (11) to write the following local error:

$$\begin{aligned} TE_i &= \frac{1}{|C_i|} \int_{C_i} (\mathbf{U}_i^{\text{ex}}(\mathbf{x}) - \mathbf{U}^{\text{ex}}(\mathbf{x})) + \frac{1}{|C_i|} \int_{C_i} \sum_{j \in D_i \cap C_{ij}} \int (\hat{\mathbf{F}}(\mathbf{U}_i^{\text{ex}}(\mathbf{x}), \mathbf{U}_j^{\text{ex}}(\mathbf{x})) - \mathbf{F}(\mathbf{U}^{\text{ex}}(\mathbf{x}))) \\ &\quad + \frac{\mu}{|C_i|} \int_{C_i} \left(\Psi_h^{\text{ex}} + \frac{h_h^{\text{ex}}}{\alpha} \nabla h_h^{\text{ex}} - \Psi^{\text{ex}} - \frac{h^{\text{ex}}}{\alpha} \nabla h^{\text{ex}} \right) \end{aligned} \quad (33)$$

having denoted by $\mathbf{U}_i^{\text{ex}}(\mathbf{x})$ the reconstructed polynomial obtained starting from the averages of a smooth exact solution $\mathbf{U}^{\text{ex}}(\mathbf{x})$, and similarly by h_h^{ex} the finite element approximation of exact nodal data. We can now proceed to a term by term estimation of the right hand side. The first one is by construction zero for quadratic polynomials, giving a rest of order $\mathcal{O}(h^3)$, and a similar result is easily proven for the second term (see e.g. [71] §2). The last one needs an estimate on the error of the solution of the elliptic step. Standard finite

element error estimates for elliptic equations (see e.g. [18, 25, 13]) rely firstly on a consistency estimates involving two main components: an estimate of the interpolation error for the solution, an estimate of the residual error, related to the approximation of the right hand side of the problem. For linear finite elements, the approximation error in L^2 norm is of an order $O(h^2)$. Concerning the right and side, and inspection of (30) reveals that the limiting factor, for constant bathymetry, is the accuracy in the approximation of the derivatives of the velocity. This indicates that, *provided that the gradient approximation is second order accurate on general meshes the consistency of the scheme is of order $O(\mu h^2)$, which is within the modelling error as soon as $h = O(\mu)$* . This is essentially the same rationale behind the method of [72], extended to unstructured grids.

To confirm numerically the impact of these choices, we consider a travelling solitary wave which is an exact solution of the GN equations for $\alpha = 1$. Depth and velocity are known analytically and given by:

$$h(x, y, t) = h_0 + \alpha_0 \text{sech}^2(k(x - c * t)) , \quad u(x, y, t) = c \left(1 - \frac{h_0}{h(x, y)} \right) \quad (34)$$

with h_0 the still water depth, α_0 the wave's amplitude, and with $k = \sqrt{3\alpha_0/4h_0^2(h_0 + \alpha_0)}$, and $c = \sqrt{g(h_0 + \alpha_0)}$. Although this is essentially a 1d solution, we have run it on 2D unstructured triangulations (rightmost picture on figure 4) to perform a grid convergence. We compare the results obtained by using the correctly sampled values of the solution and of its derivatives at the nodes, as discussed above (orange curve), against the result (blue curve) obtained by passing the nodal average as it is, and using it to construct the finite element approximation, including the elemental derivatives computed on each element as $(\nabla u)_h = \sum_{j \in K} \nabla \varphi_j u_j$. The result shows the importance of accounting for the meaning of the data in the FE/FV coupling.

6.2. A comment on well-posedness

The choice of the numerical fluxes plays a fundamental for the robustness of the hyperbolic step in presence of irregular solutions for which the use of dissipative/upwind fluxes is necessary (cf. sections §4.1.3 and §4.2). When considering the propagation of smooth dispersive waves one may think that non-dissipative fluxes could be more appropriate. It turns out that for the method propose here this is not case, and numerical dissipation plays a major role also in the propagation region. To show this we will consider the simplified setting of the linearized dimensionless equations without bathymetry, which can be written as

$$\begin{aligned} (I - \alpha\mu(\nabla\nabla\cdot))\Psi &= -\frac{\nabla\eta}{\alpha} \\ \partial_t\eta + \nabla \cdot \mathbf{u} &= 0 \\ \partial_t\mathbf{u} + \nabla\eta &= \mu\Psi + \mu\frac{\nabla\eta}{\alpha} \end{aligned} \quad (35)$$

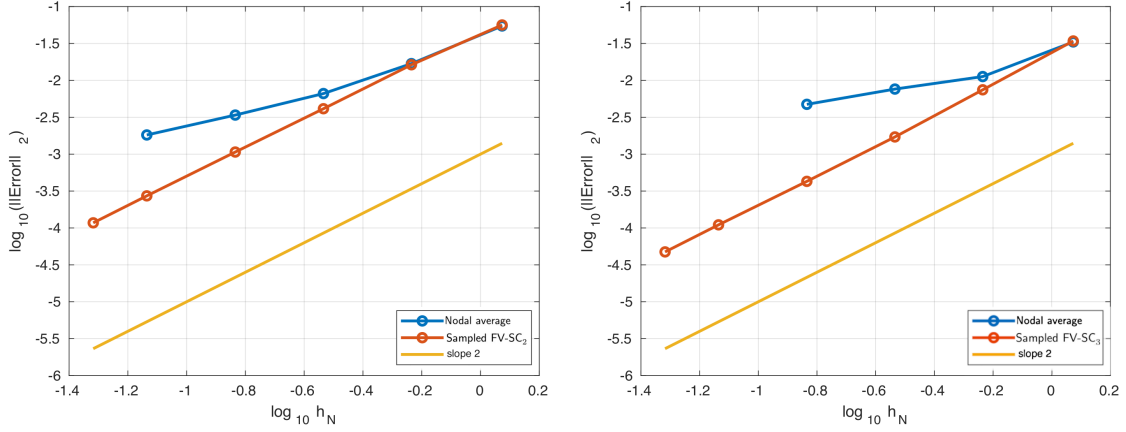


Figure 7: Error decay for the solitary wave. Left: hyperbolic step with second order successive correction. Right: hyperbolic step with third order successive correction. Orange: consistently sampled FV data used in the FE step. Blue: nodal averages in the FE step.

Despite the scheme having been derived and coded for the above form, for the discussion of this section it is more appropriate to start from the following equivalent formulation

$$\begin{aligned} \partial_t \eta &= -\nabla \cdot \mathbf{u} \\ (I - \alpha\mu(\nabla\nabla\cdot))\partial_t \mathbf{u} &= -(I - (1 - \alpha)\mu(\nabla\nabla\cdot))\nabla\eta \end{aligned} \quad (36)$$

The operator to be inverted to evolve the velocity \mathbf{u} (as well as to pre-compute Ψ) is a grad-div operator quite common in the modelling of electromagnetic waves. The important aspect of this type of equation is that it is naturally formulated in the functional space of vectors $H(\text{div})$, as its variational form involves the scalar product

$$(\mathbf{v}, \mathbf{u})_\Omega := \int_\Omega \mathbf{v} \cdot \mathbf{u} + \alpha\mu \int_\Omega \nabla \cdot \mathbf{v} \cdot \nabla \cdot \mathbf{u}$$

which readily generates the equivalent squared $H(\text{div})$ norm $\mathbf{u}^2 + \alpha\mu(\nabla \cdot \mathbf{u})^2$. It is well known that, despite the symmetry of the bilinear form induced by the equation, H^1 finite elements as the one used here to solve the elliptic problem are not well posed as prone to spurious modes related to the rotational of the solution. For electromagnetic waves this is a long time known fact [19, 20, 32]. Divergence conforming elements allow of course to side-step this issue. To use H^1 elements, which are easier to implement, and better suited to be coupled with a hyperbolic solver something needs to be done. Usually, this is achieved by introducing, at the PDE level or in the scheme, a regularising operator that stabilizes the incomplete norm generated by the variational formulation.

In our method this is essentially the case, although we don't modify explicitly the elliptic solver. The stabilization arises in our case from the use of dissipative numerical fluxes. More particularly, consider the regularized time discrete system

$$\begin{aligned} \frac{\eta^{n+1} - \eta^n}{\Delta t} - \nabla \cdot (\epsilon_h^\eta \nabla \eta)^{n+\delta} &= -\nabla \cdot \mathbf{u}^n \\ (I - \alpha\mu(\nabla\nabla\cdot)) \frac{\mathbf{u}^{n+1} - \mathbf{u}^n}{\Delta t} - \nabla \cdot (\epsilon_h^\mathbf{u} \nabla \mathbf{u})^{n+\delta} &= -(I - (1 - \alpha)\mu(\nabla\nabla\cdot)) \nabla \eta^n \end{aligned} \quad (37)$$

We have explicitly added artificial diffusion terms, which are related to the use of upwind numerical fluxes. In particular, the coefficients ϵ_h^η and $\epsilon_h^\mathbf{u}$ depend on the numerical flux. For the dimensionless linearized problem under consideration, these can be both approximated by $\epsilon_h^\eta = \epsilon_h^\mathbf{u} \approx \mathbb{C} h$ for some mesh dependent constant \mathbb{C} which we assume for simplicity to be equal to 1. If we take $\delta = 1$, the method becomes semi-implicit, with a fully implicit approximation of the numerical dissipation, as e.g. in the CPR (Centered Potential Regularization) method discussed in [54, 53]. Assuming that (in dimensionless form) $\Delta t = \nu h$ with νm in this case the second equation can be cast as

$$(I - \nu h^2(\nabla \cdot \nabla) - \alpha\mu(\nabla\nabla\cdot)) \mathbf{u}^{n+1} = -(I - \alpha\mu(\nabla\nabla\cdot)) \mathbf{u}^n + \nu h (I - (1 - \alpha)\mu(\nabla\nabla\cdot)) \nabla \eta^n \quad (38)$$

Differently from what is usually done in electromagnetics, the operator appearing now on the left hand side is a div-grad regularization naturally related to the use of upwind fluxes. This type of regularization has been proposed in the context of porous media flows and shown to overcome the weak stability of H^1 elements for the grad-div equation, see e.g. [47]. The reference in particular proves stability bounds independent on the regularization constant, which suggests that stability should be retained in our case also under mesh refinement.

For the case of a fully explicit method as the one used in our computations, we can't invoke directly the results of the last reference. In this case the velocity update reads

$$(I - \alpha\mu(\nabla\nabla\cdot)) \mathbf{u}^{n+1} = (I + \nu h^2(\nabla \cdot \nabla) - \alpha\mu(\nabla\nabla\cdot)) \mathbf{u}^n + \nu h (I - (1 - \alpha)\mu(\nabla\nabla\cdot)) \nabla \eta^n \quad (39)$$

We can formally pre-multiply both sides by the operator $(I - \nu h^2(\nabla \cdot \nabla))$ and use the relation $I - \alpha\mu(\nabla\nabla\cdot) = I + \alpha\mu(\nabla\nabla\cdot) + O(\mu^2)$ to argue that at the continuous level

$$\mathbf{u}^n = (I - \nu h^2(\nabla \cdot \nabla)) \mathbf{u}^{n+1} - \nu h (I + \alpha\mu(\nabla\nabla\cdot))(I - \nu h^2(\nabla \cdot \nabla))(I - (1 - \alpha)\mu(\nabla\nabla\cdot)) \nabla \eta^n + O(h^4, h^2\mu, \mu^2)$$

So we can estimate the effects of numerical, and argue that (39) is equivalent to

$$\begin{aligned}
(I - \nu h^2(\nabla \cdot \nabla) - \alpha \mu(\nabla \nabla \cdot))\mathbf{u}^{n+1} = & (I - \alpha \mu \nabla \nabla \cdot)\mathbf{u}^n + \nu h(I - (1 - \alpha)\mu(\nabla \nabla \cdot))\nabla \eta^n \\
& - \nu^2 h^3(\nabla \cdot \nabla)(I + \alpha \mu(\nabla \nabla \cdot))(I - \nu h^2(\nabla \cdot \nabla))(I - (1 - \alpha)\mu \nabla \nabla \cdot)\nabla \eta^n \quad (40) \\
& + O(h^4, h^2\mu, \mu^2)
\end{aligned}$$

So within a modelling and numerical error of $O(h^4, h^2\mu, \mu^2)$ the fully explicit scheme is in principle able to stabilize the grad-div approximation.

We propose some numerical evidence to confirm this fact by studying again the grid convergence of the solitary wave solution (34). We perform the following experiment. On one hand we solve the nonlinear shallow water equations forced with $\Phi = \Phi^{\text{exact}}$ obtained by replacing (34) in the momentum equation. This corresponds to imposing the solitary wave as a manufactured solution (cf [62]). On the other we solve the full system including the dispersive terms for which (34) is an exact solution if $\alpha = 1$. In both cases, we perform a grid convergence with centered numerical fluxes, as well as with the full upwind flux.

We perform the computations up to time $t = 0.1$ on 7 unstructured meshes starting with a triangulation with 328 triangles, and then halving the mesh size. The results are reported in figure 8 for the second order (left), and third order (right) polynomial reconstruction in the hyperbolic fluxes. For the shallow water equations all configurations converge. The centered fluxes fail to provide third order of accuracy, at least on the meshes considered, however full second order is observed. For the GN system, the error obtained with centered fluxes quickly stalls, and error blow-up is observed on the last meshes for the final time considered. Full second order of convergence is restored with the upwind fluxes, confirming our heuristics.

7. Dispersion error: time continuous spectral analysis

To further characterize the scheme proposed, we briefly discuss the results of a time continuous spectral analysis in the 1D case. We recall that the exact enhanced Green-Naghdi equations have a phase defined by:

$$\omega_{gn}^2 = gh_0 k^2 \frac{1 + \frac{\alpha-1}{3} k^2 h_0^2}{1 + \frac{\alpha}{3} k^2 h_0^2}, \quad (41)$$

having denoted by k the wavenumber. To characterise the dispersion error, we need to replace a Fourier mode into our scheme. So we assume that for a wavenumber k , the solution has the form $\mathbf{U} = \mathbf{U}_0 e^{\nu_h t + j k x}$, with j the imaginary unit, and where $\nu_h = \xi_h + j\omega_h$ and ξ_h, ω_h represent the discrete amplification rate and

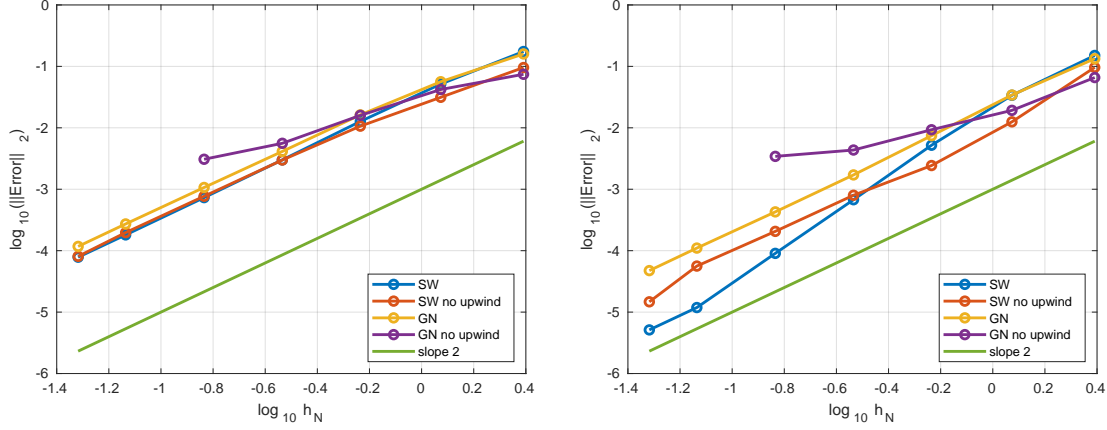


Figure 8: Solitary wave. Left: second order successive correction. Right: third order successive correction. Blue: shallow water with upwind fluxes and manufactured solution. Orange: shallow water with centered fluxes and manufactured solution. Yellow: Green-Naghdi with upwind fluxes. Magenta: Green-Naghdi with ventered fluxes.

phase. As done to couple the FV and FE method, one has to be careful here not to confuse cell averages with nodal values. We can indeed find that

$$\bar{U}_i = \int_{x_{i-1/2}}^{x_{i+1/2}} U_0 e^{v t + j k x} = \frac{U_i}{j \mu_h} (e^{j \frac{\mu_h}{2}} - e^{-j \frac{\mu_h}{2}}) = \frac{U_i}{j \mu_h} 2 \sin\left(\frac{\mu_h}{2}\right) \quad (42)$$

with $U_i = U_0 e^{v_h t + j k x_i}$, and $\mu_h = k h$. Using this relation consistently, and replacing in the one-dimensional version of the scheme we obtain

$$\begin{aligned} (j v_h C + j B) \eta_i + h_0 A u_i &= 0 \\ g A \eta_i + (j v_h C + j B) u_i &= j \mu_h \frac{1}{h_0} M^{FV} D \eta_i. \end{aligned} \quad (43)$$

The right hand side of the second equation is the inverse of the Fourier symbol of the finite element discretization of the elliptic equation reducing in 1D to a Poisson equation (cf. e.g. [29]). The coefficients A , B , C , D are reported in Appendix B for completeness. The impact of the definition of the matrix M^{FV} arising depending on the quadrature of Φ is also provided. Only the best results are left here for clarity. These are obtained when all matrices are evaluated exactly. Straightforward computations show that

$$\omega_{\Delta_x}^2 = (g h A^2 + \mu M^{FV} A D h) / C^2. \quad (44)$$

Figure 9 plots the relative dispersion errors (rde) $|\omega_s - \omega_{gn}| / \omega_{gn}$ with respect to the dispersion relation of the eGN equations (41). For comparison, we also plot the dispersion errors of sake of reference we are also

plotting the relative dispersion errors that comes when using 1. a second and fourth order finite difference schemes (see [29, 61] for the full expressions). The pictures show that dispersion error of the scheme presented in this paper is smaller or close to the one of FD4 depending the number of nodes per wavelength. For completeness we also plot on figure 10 the error $|\omega_s - \omega_{\text{airy}}|/\omega_{\text{airy}}$ with respect to the exact dispersion relation $\omega_{\text{airy}}^2 = ghk^2 \tanh$). We observe a compensation between discretization error and modelling error so that the scheme has a lower error w.r.t the exact dispersion relation than it has wrt the Green-Naghdi one. This is related to the interaction of the phase advance/lag error of the scheme and the form of the phase of the Boussinesq model. This could be in the future a design criterion for low order (second and third) schemes.

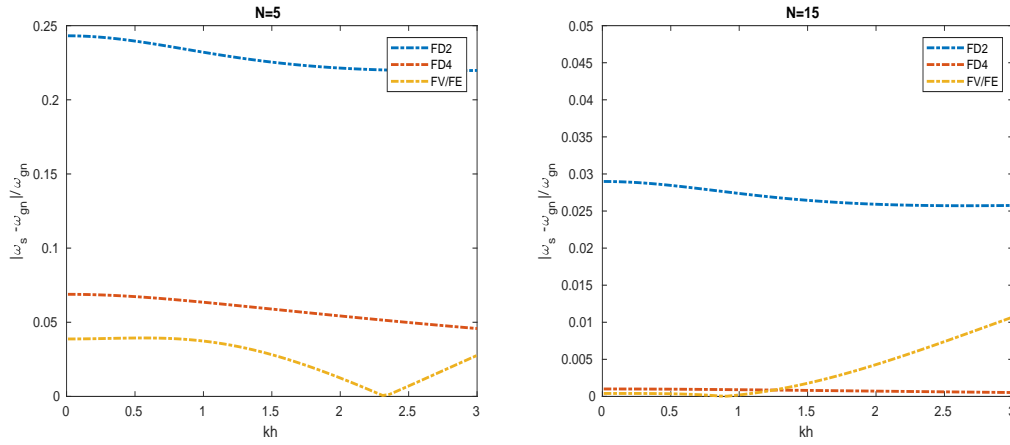


Figure 9: Dispersion error with respect to ω_{gn} for nodes per wave length $N = 5$ and $N = 15$

8. Boundary conditions, wave generation, and wave breaking closure

8.1. Wave breaking detection and closure

A hybrid strategy for wave breaking treatment is implemented in the scheme. We first estimate the location of breaking waves using explicit criteria, then we apply the NLSW equations to solve the flow in breaking regions and the GN ones elsewhere. Following the work of [38], we use the combination of the two above phase-resolving criteria for the triggering mechanism:

- the surface variation criterion: $|\eta_t| \geq \gamma \sqrt{gh}$ with $\gamma \in [0.35, 0.65]$
- the local slope angle criterion: $\|\nabla\eta\| \geq \tan \phi_c$ with ϕ_c the critical angle value.

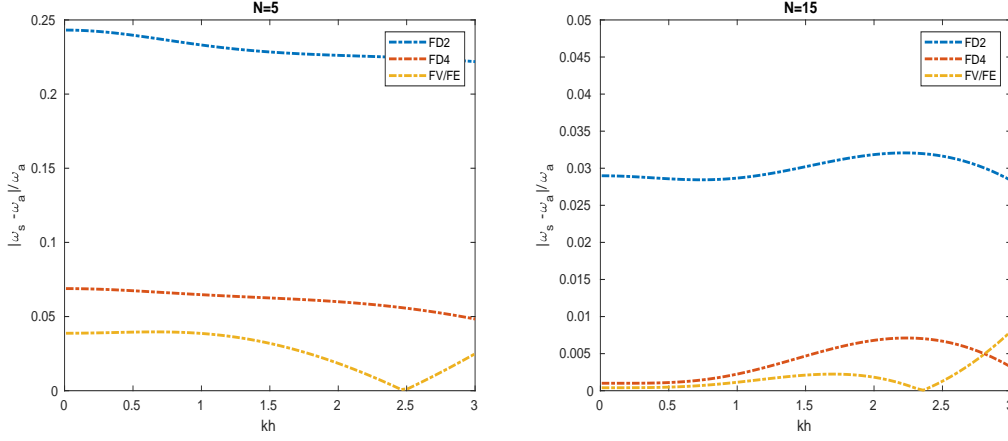


Figure 10: Dispersion error with respect to ω_{airy} for nodes per wave length $N = 5$ and $N = 15$

The values of γ and ϕ_c are depending on the type of the breaker. The first criterion flags for breaking when η_t is positive, since breaking starts on the front face of the wave, while the second criterion, acting complementary to the first, is useful for the detection of hydraulic jumps. In this work the value of $\phi_c = 30^\circ$ is used. Moreover, the estimation of the Froude number of the wave is used to established when to switch of the breaking and to detect non-breaking bores. A practical implementation of the breaking mechanism can be found in [38, 29, 27].

8.2. Wave generation and boundary conditions

In this work, we have implemented periodic, fully reflective/wall and absorbing boundary conditions. For the walls for all boundary nodes we modify the elliptic solver to set the conditions

$$\begin{aligned}\Phi \cdot \mathbf{n} = 0 &\Rightarrow \Psi \cdot \mathbf{n} = -h \frac{\nabla \eta \cdot \mathbf{n}}{\alpha} \\ \partial_{\mathbf{n}} \Phi \cdot \tau = 0 &\Rightarrow \partial_{\mathbf{n}} (\Psi \cdot \tau + h \frac{\nabla \eta \cdot \tau}{\alpha}) = 0\end{aligned}$$

For straight walls the first condition is a consequence of $\mathbf{u} \cdot \mathbf{n} = 0$, while the second forces the rotational components of Φ to zero at the walls, which is consistent with the condition $\nabla \wedge \Phi = 0$ which can be easily proved for the continuous equation used in the bulk. In practice we modify both the RHS and matrix of the elliptic system to account for these conditions. In the hyperbolic solver we also make sure that $\mathbf{u} \cdot \mathbf{n} = 0$ by modifying the hyperbolic fluxes and source and removing the normal component.

Absorbing boundaries are also applied in order to dissipate completely the energy of the incoming waves, trying at the same time to eliminate any non-physical reflection. This kind of boundaries requires the definition of a sponge layer in which the surface elevation and the momentum are damped by multiplying their values by the coefficient:

$$m(\mathbf{x}) = \sqrt{1 - \left(\frac{\mathbf{x} - d(\mathbf{x})}{L_s}\right)^2}.$$

where L_s is the sponge layer width and $d(\mathbf{x})$ is the normal distance between the cell center with coordinates \mathbf{x} and the absorbing boundary. Typical values of the sponge width are related to the wave length λ of the incoming wave and usually are: $\lambda \leq L_s \leq 1.5\lambda$.

A large number of numerical tests demand the generation of monochromatic waves. One very common approach is to use an internal wave generator. This means, generating the waves in the computational domain avoiding complications that boundaries can create. In this work we, thus, make use of an internal wave generator, first described in [73]. It is obtained by adding a source term to the mass equation. More details on the topic can be found in [61].

9. Numerical validation and benchmarking

9.1. Grid convergence for the solitary wave

We return to the exact solitary wave to compare on figure 11 the errors obtained when using the second (blue lines) and third order (orange lines) reconstructions. Convergence plots are reported for the free surface, and for Φ . We observe that using the third order scheme in the hyperbolic step, thus passing from $O(h^2)$ to $O(\mu h^2)$, allows a more or less consistent reduction of the error by a factor between 3 and 5.

9.2. Circular shoal

Whalin [74] studied the focusing effect induced by a semi-circular shoal on wave trains of different periods. The wave tank used is of $6.096m$ wide and $25.6m$ long. A semi-circular shoal was placed at the middle portion of the tank leading the water depth to decrease from $0.4572m$ at the wave maker region to $0.1524m$ to the end of the tank. The topography function reads as:

$$z = \begin{cases} 0 & \text{if } 0 \leq x < 10.67 - G(y), \\ (10.67 - G(y) - x)/25 & \text{if } 10.67 - G(y) < x \leq 18.29 - G(y), \\ 0.3048 & 10.29 - G(y) \leq x \end{cases} \quad (45)$$

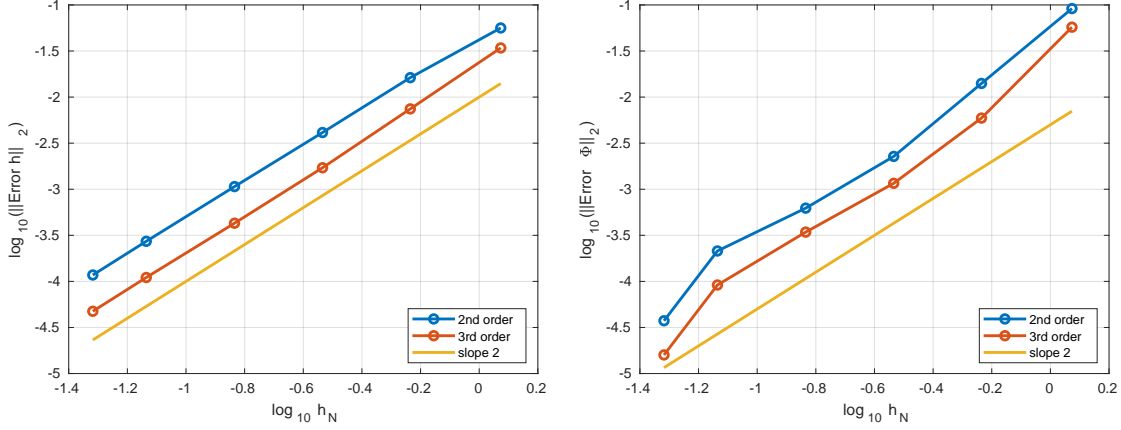


Figure 11: Error decay for second and third order schemes for h and Φ .

with $G(y) = \sqrt{y(6.096 - y)}$. This test case has been used by many authors in order to test the dispersive properties of their models, see for example [45, 6, 68, 26, 67, 37, 61]. The three test cases that have been reproduced here are:

- (a) $T = 1s, A = 0.039m, h_0/\lambda=0.306$ and $\epsilon = 0.085$
- (b) $T = 2s, A = 0.015m, h_0/\lambda=0.117$ and $\epsilon = 0.033$
- (c) $T = 3s, A = 0.0136m, h_0/\lambda=0.075$ and $\epsilon = 0.030$

where T is the wave period, A the wave height and λ the wave length. Wave gauges are placed along the center line to record the time series of the free surface elevation, which are analysed in the frequency domain using a Discrete Fourier Transform (DFT) to obtain the first three harmonic amplitudes. The computational domain used is $[-10, 36] \times [0, 6.096]m$. Periodic waves are generated using the internal generator placed at $x = 4m$ and sponge layers of $6m$ are placed at the left and right end of the domain. Reflective boundary conditions are imposed at the remaining boundaries. For the computation of the first case a triangular grid was used, consisting of triangles with side lengths $h_x = 0.05m$ and $h_y = 0.1m$ leading to a mesh of $N = 56211$ nodes. For the last two cases the grid consists of equilateral triangles, with $h_N = 0.01m$ leading to a mesh of $N = 28151$ nodes. The CFL value used was 0.5.

For all the test cases the incoming waves are linear in the deeper portion of the tank and they are steepening due to the wave shoaling. Wave energy gradually spreads out to higher harmonics, which increase in amplitude in the shoaling region. For both case (a) and (b) the agreement between the numerical and experimental data is quite satisfactoring for all the harmonics while in the last test case (c) the numerical results

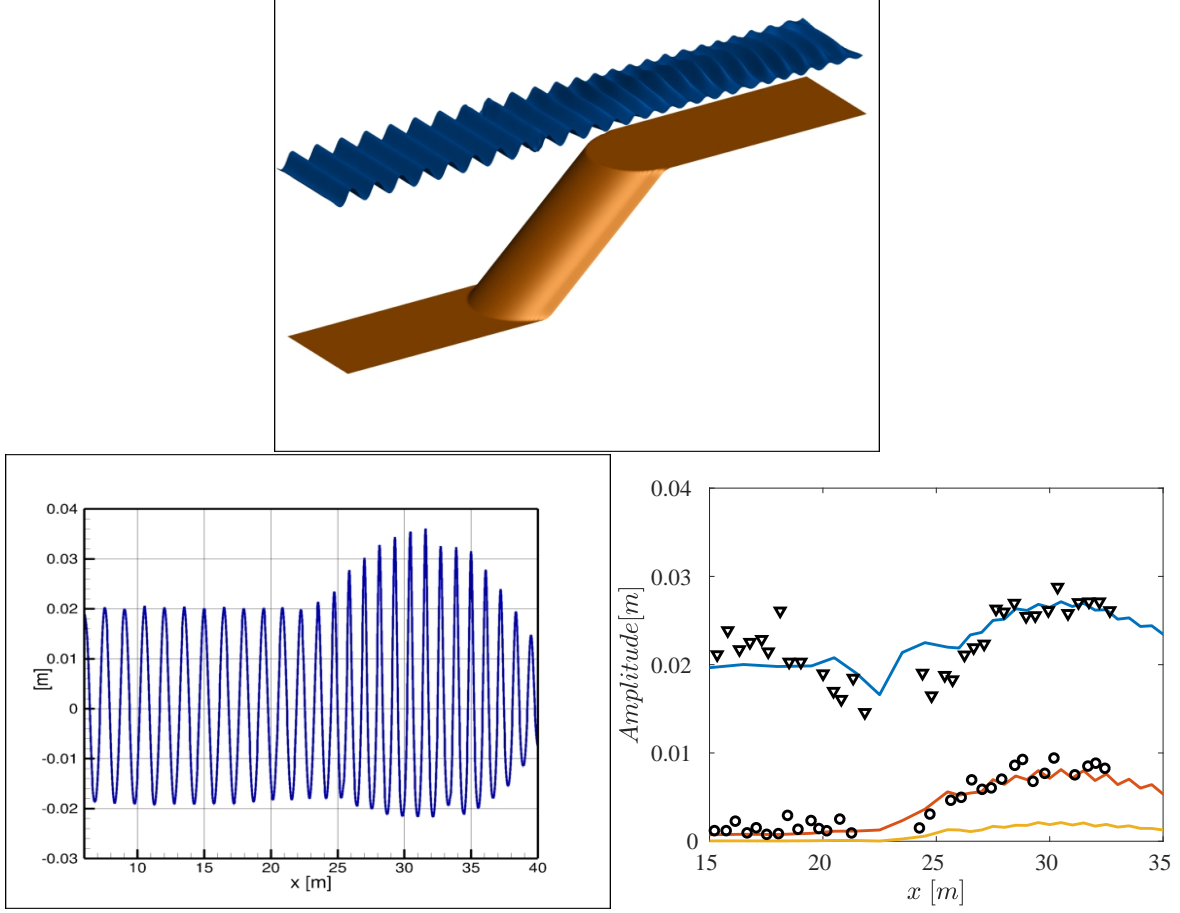


Figure 12: Wave diffraction over a semi-circular shoal. Case (a): Left: Computed free surface data along the centerline of the domain. Right: comparison of the DFT of the numerical data over the centerline with experimental data.

overestimates the first harmonic and underestimates the other two. This behavior has also been observed by other authors [45, 6, 67, 43] and the discrepancies are attributed to the shorter evolution distance of this test case or to the presence of free reflected waves.

9.3. Elliptic shoal

This test case studies the refraction and diffraction of a regular wave over a complex bathymetry and it is a reproduction of the experiment of [7]. It is mainly used to verify models based on mild-slope equations but also the extended Boussinesq-type equations. The numerical domain is 20m wide and 22m long, with $x \in [-10, 10]$ and $y \in [-17, 15]$. The bathymetry consists of an elliptic shoal placed on a ramp of constant

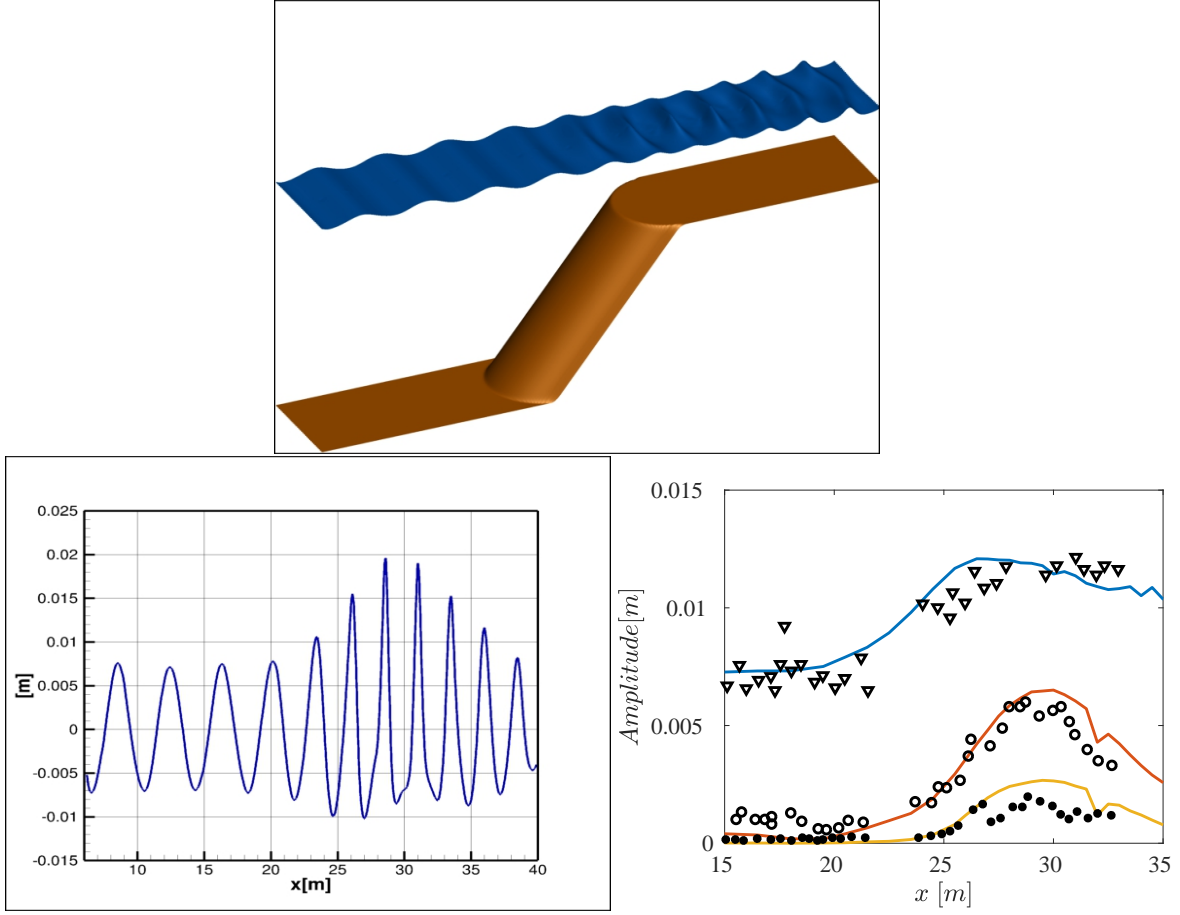


Figure 13: Wave diffraction over a semi-circular shoal. Case (b): Left: Computed free surface elevation at time $t = 40\text{sec}$. Right: comparison of the DFT of numerical data over the center-line with experimental data.

slope forming a 20° angle with the x axis. The maximum water depth is $h_0 = 0.45\text{m}$ at the wave maker's position which is placed at $y = -13\text{m}$. The bathymetry is described by the formula $z = z_0 + z_s$, where

$$z_0 = \begin{cases} (5.82 + y_r)/50 & \text{if } y_r \leq -5.82 \\ 0 & \text{otherwise,} \end{cases} \quad (46)$$

$$z_f = \begin{cases} -0.3 + \frac{1}{2} \sqrt{1 - \left(\frac{x_r}{5}\right)^2 - \left(\frac{4y_r}{15}\right)^2} & \text{if } \left(\frac{x_r}{4}\right)^2 + \left(\frac{y_r}{3}\right)^2 \leq 1 \\ 0 & \text{otherwise,} \end{cases} \quad (47)$$

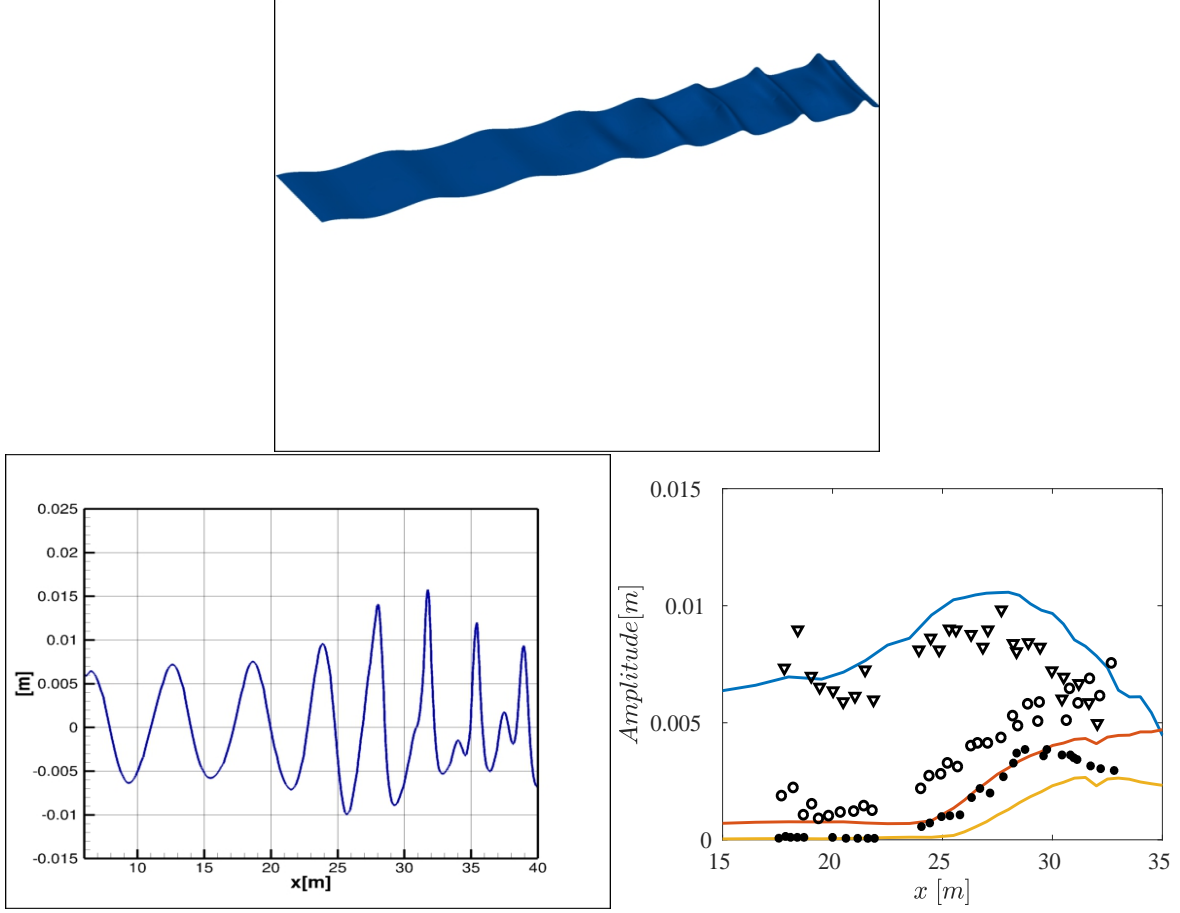


Figure 14: Wave diffraction over a semi-circular shoal. Case (b): Left: Computed free surface elevation at time $t = 40\text{sec}$. Right: comparison of the DFT of the numerical data over the center-line with experimental data.

with the transformed coordinates $x_r = x \cos(20^\circ) - y \sin(20^\circ)$ and $y_r = x \cos(20^\circ) + y \sin(20^\circ)$. The monochromatic wave's characteristics are: period $T = 1\text{s}$ and amplitude $a = 0.0232\text{m}$ corresponding to a non linearity degree $\epsilon = a/h = 0.3$. The normalized time average wave height was measured in eight different sections (see figure). Wall boundary conditions are imposed on the left and right boundaries, while sponge layers of 4m are placed on the bottom and top ends of the domain. In this test case an unstructured grid was used, which is refined in the region of the shoal. In particular the grid size in the y direction varies from $h_y \approx 0.1\text{m}$ on the top and bottom boundaries to $h_y = 0.05\text{m}$ in the region around the shoal. The simulation period is 50s , and the CLF used is 0.5 . A three dimensional view of the water elevation at the final time is reported

on figure 15. In order to compare the numerical results with the experimental data [7], time dependent data have been extracted, the last 25sec of the simulation, from the sections 1-8. The time series of the surface elevation are analysed using the zero-up crossing technique to isolate single waves and to compute the averaged wave height. The results are normalised by the incoming wave height $2a = 0.0464m$ and are reported on figures 16 and 17. The agreement on the numerical results and the experimental data are satisfactory and comparable to the results given by the literature ([61],[67], [68]).

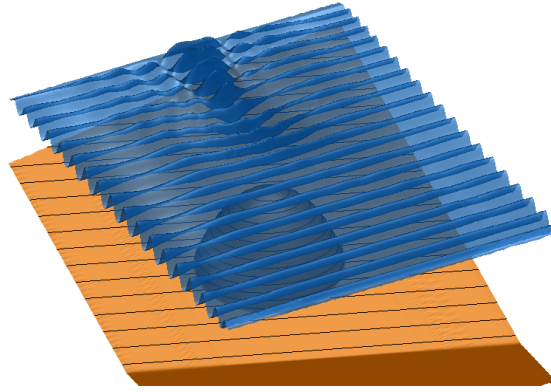


Figure 15: Wave diffraction over an elliptical shoal: 3D view of the free surface elevation at time $t = 50sec$.

9.4. Solitary interacting with a cylinder

In this test case we examine numerically the propagation and scattering of a solitary wave with a vertical cylinder. The laboratory experiment [22] investigates the interaction of the wave with the cylinder and it has been used by [36, 37] and references therein to validate their numerical models. The numerical domain used here is $x \times y \in [0, 14m] \times [0, 0.55m]$. A vertical cylinder of diameter $0.16m$ was placed at $(x_0, y_0) = (8.5, 0.275)$. The solitary has amplitude $A = 0.0375m$ and placed at $x = 4m$. The undisturbed water depth is $h_0 = 0.15m$ so the nonlinearity of the wave is $\epsilon = 0.25$. Six wave gauges were recording the free surface elevation located at: $wg1 = (8.4, 0.275)$, $wg2 = (8.5, 0.170)$, $wg3 = (8.5, 0.045)$, $wg4 = (8.6, 0.275)$, $wg5 = (8.975, 0.275)$, $wg6 = (9.375, 0.275)$. The mesh has.. nodes and is refined around the cylinder. Figure 18 presents the 3d view of the free surface elevation after 4 sec, when the solitary interacts with the cylinder. This interaction causes the generation of scattering waves that propagate downstream while the rest of the wave recovers to a solitary wave and propagates upstream. The first wave that interacts

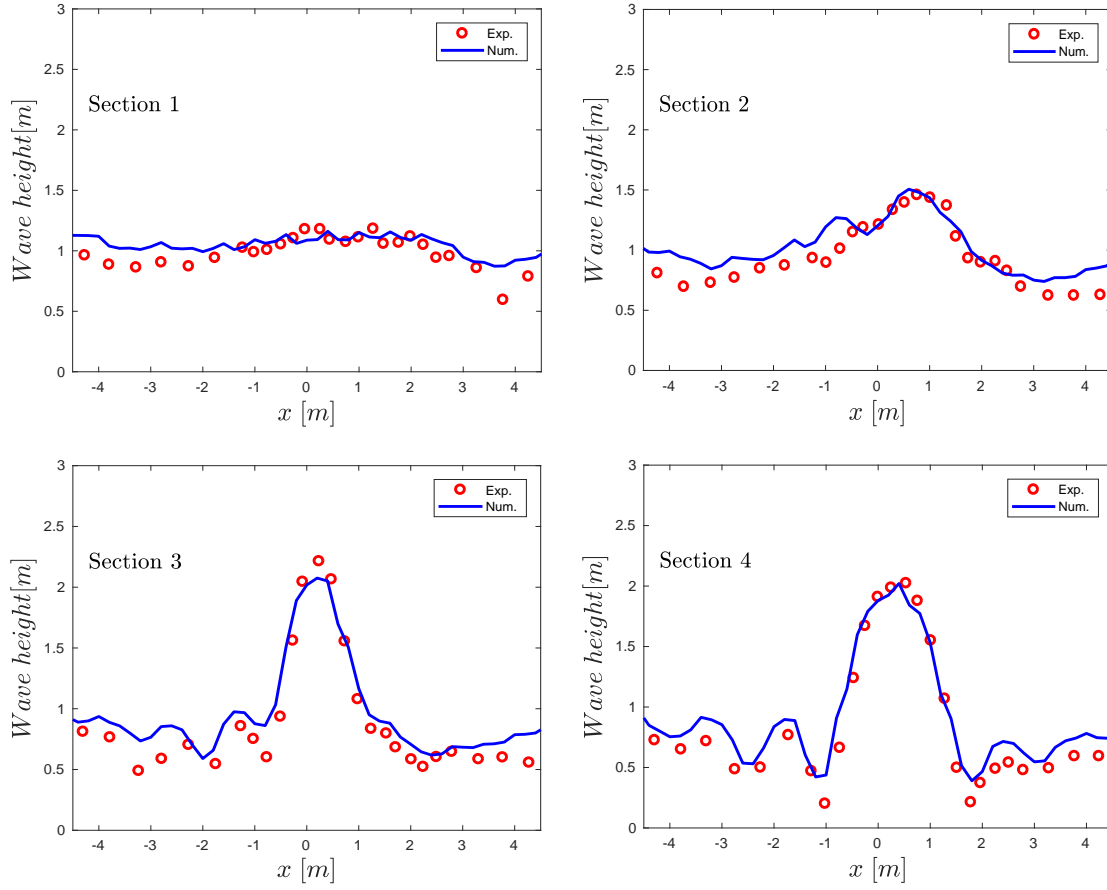


Figure 16: Wave diffraction over an elliptic shoal: Normalised wave height from sections 1 to 4.

with the cylinder and propagates upstream is computed quite accurately compared to the experimental data as seen in figures 19-21. The reflected waves, even though are better resolved compared to the ones that can be found in the bibliography [36, 37], presents some discrepancy compared to the experimental data. This may indicate that a fully dispersive model is needed for this case.

9.5. Solitary wave breaking on a 3D reef

Swigler et Lynett (2011) performed laboratory experiments at the O.H. Hinsdale Wave Research Laboratory of Oregon State University to study the specific phenomena which occurs when a tsunami like wave approaches the coast, such as shoaling, refraction, breaking and run-up. Many authors have used this case to validate their codes, since it is quite demanding and involves multiple physical phenomena. The computa-

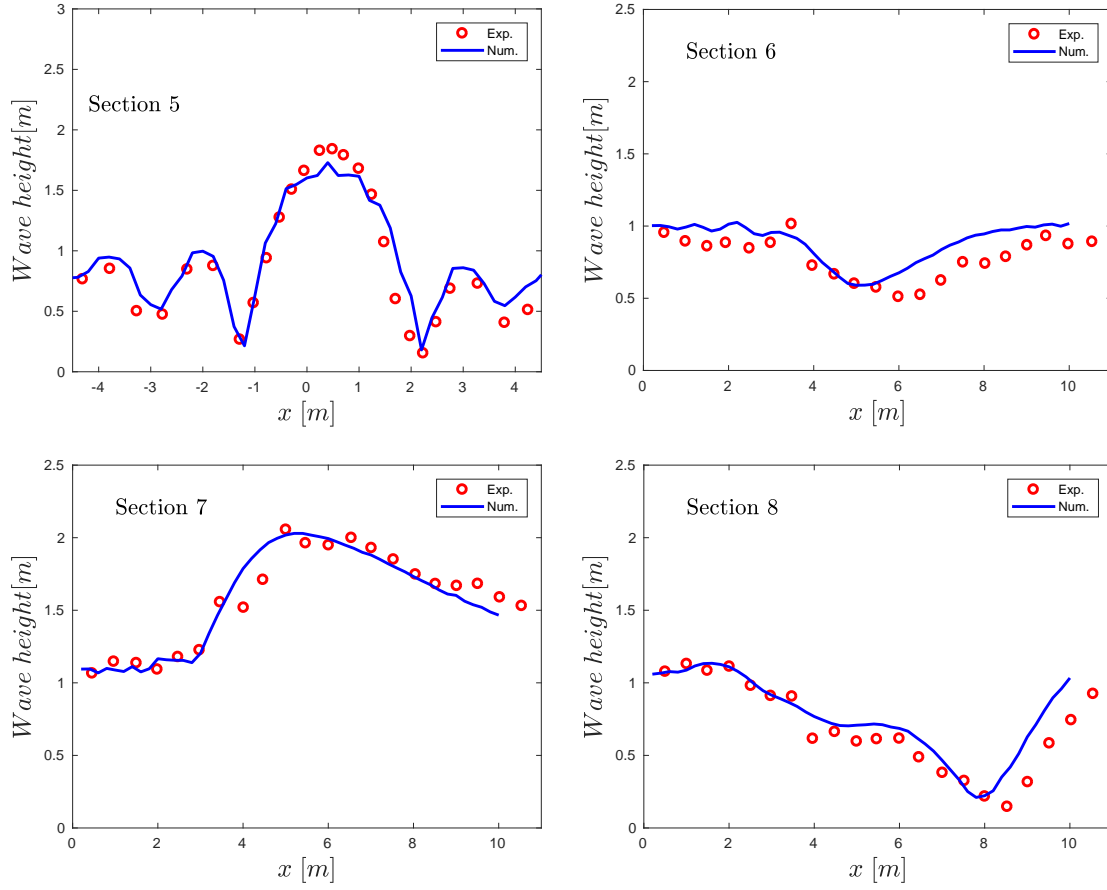


Figure 17: Wave diffraction over an elliptic shoal: Normalised wave height from sections 5 to 8.

tional domain is $48.8m$ long and $26.5m$ wide. Figure 22 shows the test set up along with the position of the wave gauges. The topography is determined from a laser scan and it consists of a slope of 1:30 connected with a triangular reef flat submerged between $0.75m$ and $0.9m$ below the still water level. The offshore shelf edge has an elevation of $0.71m$ with the apex located at $x=12.6m$. The steepest slope of the shelf is at the apex and becomes milder moving along the shelf edge toward the basin side walls. A concrete cone is also placed at the apex of the reef between $x=14m$ and $20m$. It has a diameter of $6m$ and a height of $0.45m$. The planar beach continues up to $x=31m$ and then becomes level until the back of the basin. Nine wave gauges were placed into the basin in order to measure the variation of the free surface elevation: gauges 1, 2, 3, 7 were located at $y = 0m$ and $x = 7.5, 13, 21, 25m$ gauges 4, 5, 6, 8 were located at $y = 5m$ and $x = 7.5, 13, 21, 25m$, while gauge 10 have been set at $y = 10m$ and $x = 25m$.

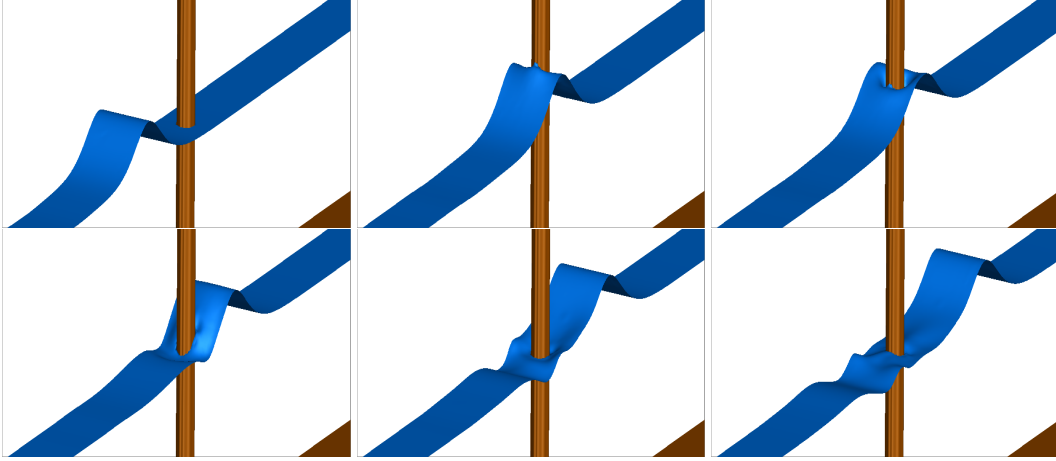


Figure 18: 3D view of the free surface elevation. Interaction of the solitary wave with the vertical cylinder.

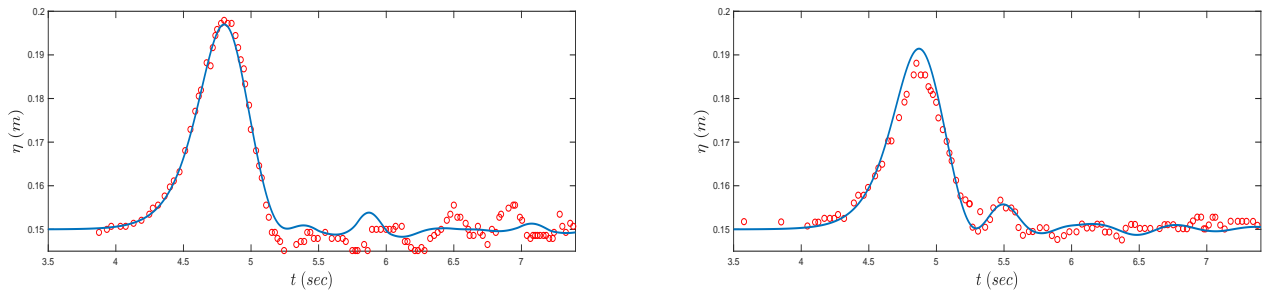


Figure 19: Free surface elevation at wave gauges 1 and 2

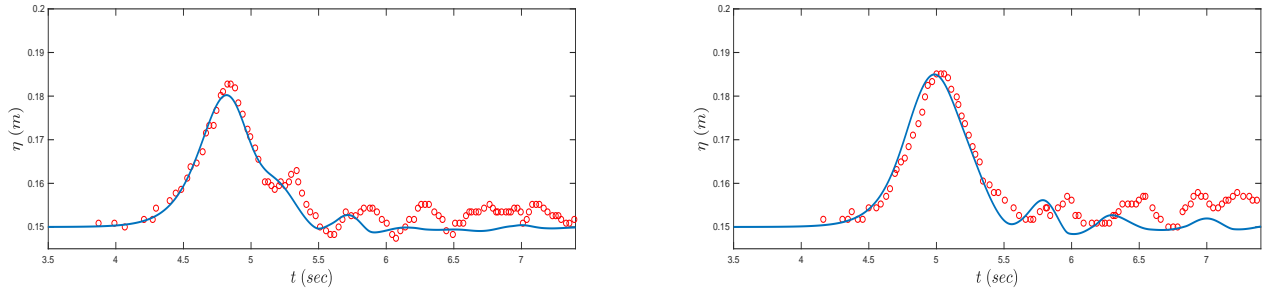


Figure 20: Free surface elevation at wave gauges 3 and 4

Compared to the experimental case, the computational domain has been extended from $x = 0m$ to $x = 5m$

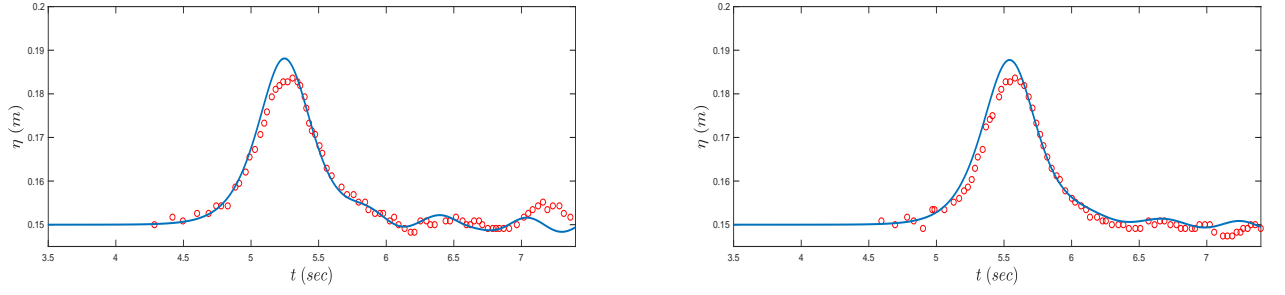


Figure 21: Free surface elevation at wave gauges 5 and 6

in order to be able to completely contain the initial solitary wave. It has been discretized by means of a non-uniform unstructured grid, adapted to the bed curvature, as shown in 22, and characterized by reference maximum and minimum size respectively: $\max(h_K) = 0.3m$ and $\min(h_K) = 0.125m$. A solitary wave of amplitude $A = 0.39m$, corresponding to $\epsilon = 0.5$, is initially placed in $x = 0$ and wall reflecting boundary conditions are imposed in each boundary of the domain. We used a Manning coefficient $N_m = 0.014$ for representing bed roughness. A CFL number of 0.5 was used, together with $\gamma = 0.6$ for the breaking detection criterion.

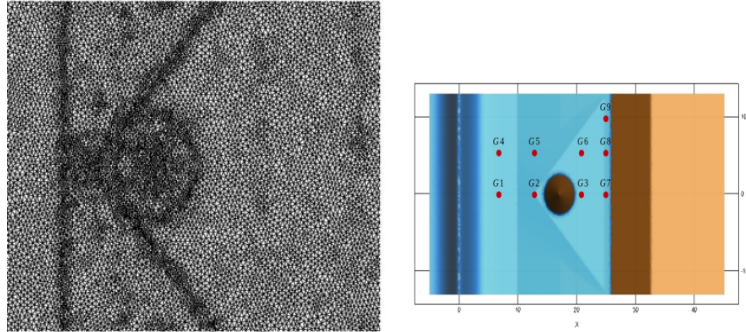


Figure 22: Solitary wave propagation over a three dimensional reef: Close up view of the adapted mesh used for the computation(right) and positioning of wave gauges (left)

Figure 23 shows the computed free water surface at different time instants. With white color we denote the time evolution of the breaking regions detected by the criteria of the breaking mechanism. As the solitary wave propagates towards the beach it shoals, increases its steepness and nonlinearity, up to reaching a

breaking point at $t = 5$ on the center line of the domain, when it reaches the apex of the triangular shelf. At $t = 6.5s$ the central part of the wave has completely overtopped the concrete cone, while on the two sides, the surge continues to shoal, diffracting around the base of the cone. By $t = 8.5s$, the refracted and diffracted waves collide on the lee side of the shelf. After $t = 9s$, the water starts to withdraw from the cone top and a bore-front forms, from the combined waves after the diffraction, and propagates on the shelf behind the cone and then onshore. After $t = 15s$, a new bore is created from the drawn-down of the water and collides with the refracted waves.

The next figure 24 plots the computed free surface time series on the wave gauges 1- 9 against the experimental data. The arrival of the first incoming wave is correctly captured in gauges 1 and 2, as it is for the refracted and diffracted waves at the lee side of the cone, as can be seen from gauge 3. The signal at the gauges located at the north side of the cone indicates that wave shoaling, breaking and propagation on the shelf is accurately predicted, together with the complex nonlinear interaction between diffracted and refracted waves.

9.6. The seaside experiment

This final test case examines the impact of a tsunami on an urban area. The laboratory benchmark took place in the Oregon State University and has been served as a blind benchmark test case for the NTHMP Mapping and Modeling Benchmarking Workshop: Tsunami Currents (http://coastal.usc.edu/currents_workshop/index.html). A physical and numerical comparison presented in [55]. It has also been used in the project TANDEM (<http://www-tandem.cea.fr>) as a benchmark test case. It involves a complex topography including a seawall and several buildings inspired of the real city of Oregon at 1:50 scale. It is a very demanding test case since it tests the ability of the model to simulate flows in complex and real topography. It involves wave propagation, shoaling wave breaking and flooding in an urban area involving wet/dry fronts. The rectangular basin was 48.8m long 26.5m wide, 2.1m deep and was equipped with a piston type wave maker with a maximum stroke of 2.1m and maximum velocity of 2 m/sec. Details in the physical experiment can be found in [55]. The data available involves a detailed topography, the forcing signal the positions as well as the time histories of water height and velocity in a large number of wave gauges in the streets of the city model. The numerical wave tank is $x \times y \in$ and an unstructured triangular mesh of $N = 172854$ is used with a min $h_{ref} = 0.03$. The mesh is refined in the region where the buildings are placed. The individual structures and buildings are approximated as steep-sided topography

and a manning coefficient of 0.01 is used. The CFL number is set to 0.3 and the the breaking parameters are $\gamma = 0.3$ and $\tan(\phi_c) = 0.37$.

The incoming wave enters from the left boundary and wall boundary conditions are imposed in the remaining boundaries. In the domain, 31 wave gauges measure the free surface elevation and the velocity . Figure 25 sows the measurement locations in the onshore region. Furthermore, four surface wave gauges were fixed in the basin at the following locations: $wg1 = (2.086, -0.515)$, $wg2 = (2.068, 4.065m)$, $wg3 = (18.618, 0.000m)$ and $wg4 = (18.618, 2.860m)$. Time series comparison of the free surface elevation between the experimental data and the numerical ones at Locations A, B, C, D are shown in figures 27, 28, 29, 30 respectively. In all figures we perform comparison for the GN equations (blue line) and the NSWE (black line). Figure 31 presents the comparison of the free surface elevation between the experimental and numerical data in the wave gauges WG1-WG4.

10. Conclusions

In this work we presented a hybrid numerical approach for the solution of the Green-Naghdi equations on unstructured meshes. We split the original system in a hyperbolic and an elliptic part. For the hyperbolic part, we used a a third order in space, node centered FV scheme. We achieve a third order reconstruction of the physical variables by means of a successive correction method which iteratively improves derivatives computed by means of the standard Green-Gauss formula. This approach guarantees third order accuracy even on unstructured meshes. In order to prevent oscillations on non smooth solutions we used a slope limiter [49] applied for a first time on a node centered scheme using the derivative reconstruction via the successive correction. We coupled the non-dispersive hyperbolic solver to a second order solver for the physical weakly dispersive effects. We used the standard P1 FE method. We ensured compatible data representation in the two phases of the computations since both methods evolve unknowns associated to the nodes, but with a totally different meaning.

We examined the impact of those data representation on the theoretical accuracy by writing an estimate of the local truncation error for constant bathymetry concluding that providing a second order accurate gradient on general meshes the consistency of the scheme is of order $O(\mu h^2)$. The importance of accounting for the meaning of the data in the FE/FV coupling has also been confirmed by performing a convergence test.

An other conclusion of this work is related to the choice of the numerical fluxes in the hyperbolic step. It turns out that we have to use dissipative/upwind fluxes in order to stabilize the incomplete norm generated

by the variational formulation, since we use H^1 finite elements, in the elliptic step. We provided numerical evidences to confirm this fact by studying again grid convergence.

Finally we showed that when using the third order scheme in the hyperbolic step, allows a consistent reduction of the error by a factor between 3 and 5.

Acknowledgments

The authors would like to thanks Dr Pierre Brenner for the useful discussion and exchange of notes on the topic of higher order reconstruction and the successive correction method.

Appendix A.

For a C_i a computational cell the geometric tensor of order k is given by

$$\frac{1}{|C_i|} \int \int_{C_i} (\mathbf{x} - \mathbf{x}_{G_i})^{\otimes k} dS$$

with \otimes the tensor product. This means for a third order scheme ($k = 2$) we will have to define before the beginning of our computations the geometric moments:

$$\begin{aligned} (XY)_i &= \frac{1}{|C_i|} \int_{C_i} (x - x_{G_i})(y - y_{G_i}) dS, \\ (X^2)_i &= \frac{1}{|C_i|} \int_{C_i} (x - x_{G_i})^2 dS, \\ (Y^2)_i &= \frac{1}{|C_i|} \int_{C_i} (y - y_{G_i})^2 dS. \end{aligned}$$

The construction of the M_2 matrix is based on the calculation of three specific mesh depended variables

$$\begin{aligned} \alpha_i &= \frac{x_{G_i}^2 + (X^2)_i}{2}, \\ \beta_i &= \frac{y_{G_i}^2 + (Y^2)_i}{2}, \\ \gamma_i &= x_{G_i}y_{G_i} + (XY)_i \end{aligned}$$

where G is the cell's center of gravity and X_i^2 , Y_i^2 , XY_i are the second order momentum of the cell C_i . Then, M_2^{-1} is a 3×3 matrix and is computed using the second derivatives of the moments. It is defined as:

$$M_2 = \begin{bmatrix} (\alpha_i)_{xx} & (\beta_i)_{xx} & (\gamma_i)_{xx} \\ (\alpha_i)_{yy} & (\beta_i)_{yy} & (\gamma_i)_{yy} \\ ((\alpha_i)_{xy} + (\alpha_i)_{yx})/2 & ((\beta_i)_{xy} + (\beta_i)_{yx})/2 & ((\gamma_i)_{xy} + (\gamma_i)_{yx})/2 \end{bmatrix}$$

Appendix B.

We report here the coefficients of system (43) presented in section 7 and involved in the dispersion properties of the scheme. From the discretization of the hyperbolic part using the FV scheme we easily obtain system (43) with :

$$A = \frac{k}{\mu} \left(\frac{81}{48} \cos\left(\frac{3\mu}{2}\right) - \frac{69}{48} \cos\left(\frac{\mu}{2}\right) - \frac{13}{48} \cos\left(\frac{5\mu}{2}\right) + \frac{1}{48} \cos\left(\frac{7\mu}{2}\right) \right) \quad (\text{B.1})$$

$$B = \frac{jck}{\mu} \left(\frac{21}{48} \sin\left(\frac{\mu}{2}\right) + \frac{61}{48} \sin\left(\frac{3\mu}{2}\right) - \frac{14}{48} \sin\left(\frac{5\mu}{2}\right) + \frac{1}{48} \sin\left(\frac{7\mu}{2}\right) \right) \quad (\text{B.2})$$

$$C = 2 \sin\left(\frac{\mu}{2}\right). \quad (\text{B.3})$$

From the solution of the elliptic problem we get:

$$D = j \frac{-gh^3}{3} T^G (M^G)^{-1} \left(M^G - \frac{ah^2}{3} S^G \right)^{-1} \text{ and} \quad (\text{B.4})$$

$$M^{FV} = \frac{1}{8} (6 + 2 \cos(\mu)) \quad (\text{B.5})$$

where

- The galerkin mass matrix: $M^G = \frac{1}{6} (4 + 2 \cos(\mu))$
- second order space derivatives: $S^G = \frac{k^2}{\mu^2} (2 \cos(\mu) - 2)$
- third order space derivatives: $T^G = \frac{k^3}{2\mu^3} (2 \sin(2\mu) - 4 \sin(\mu)).$

In a similar way to [29], we analyse in figure B.32 different discretization possibilities with or without mass lumping on M^{FV} and M^G matrices in figure B.32. Now the curves are obtained for two values $kh = 0.5$ and $kh = 2.5$, corresponding to a long and to a shorter wave (or shallow and deep waters respectively) and plotted against the inverse number of the number of nodes per wavelength. The optimum choice that minimizes the rde is not to lump any matrix, which was also the case for the 1DH scheme of [29]. This choice is thus employed in this work.

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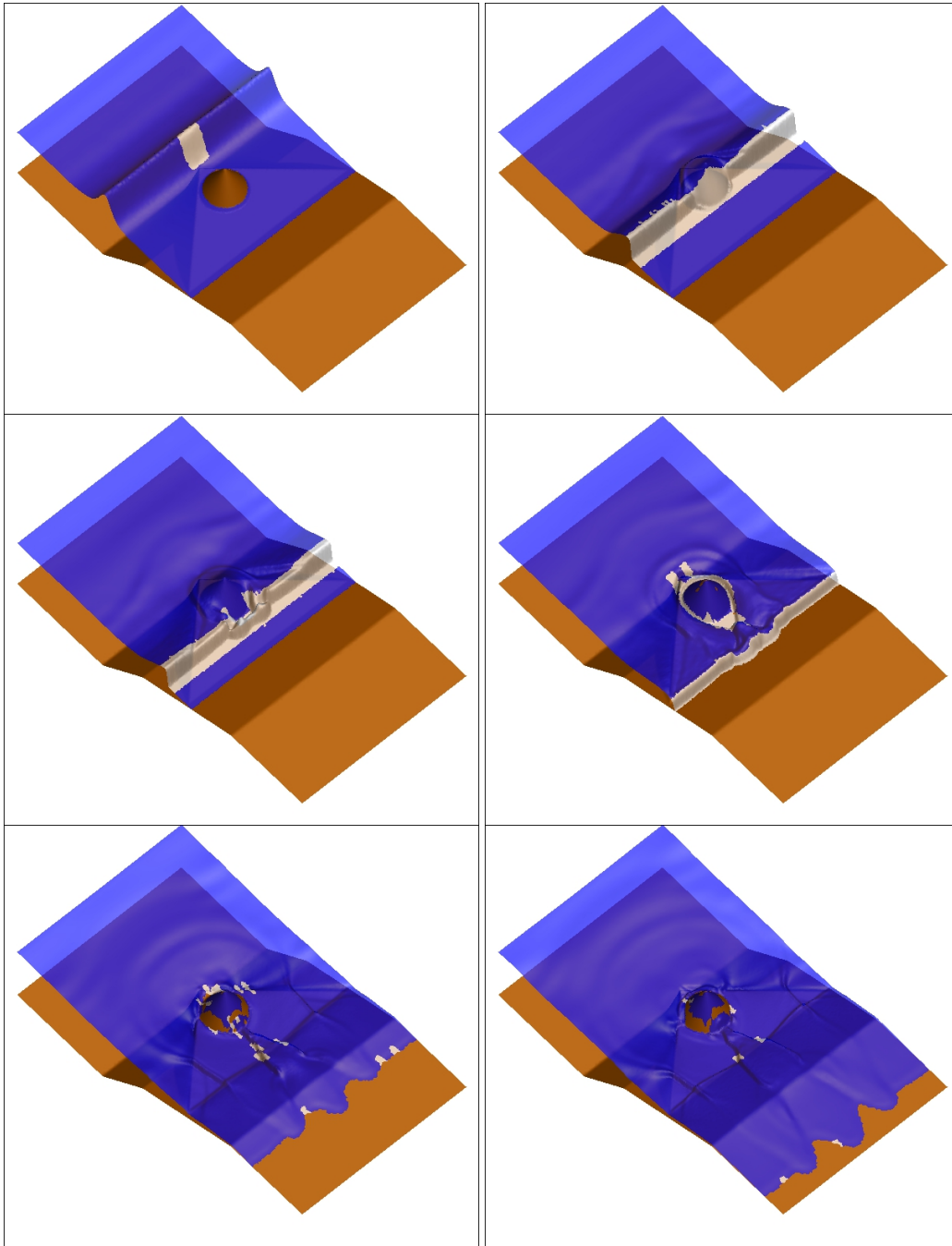


Figure 23: Solitary wave propagation over a three dimensional reef: evolution of free surface solution. The white area represents the region where wave breaking is detected and the NLSW equations are solved.

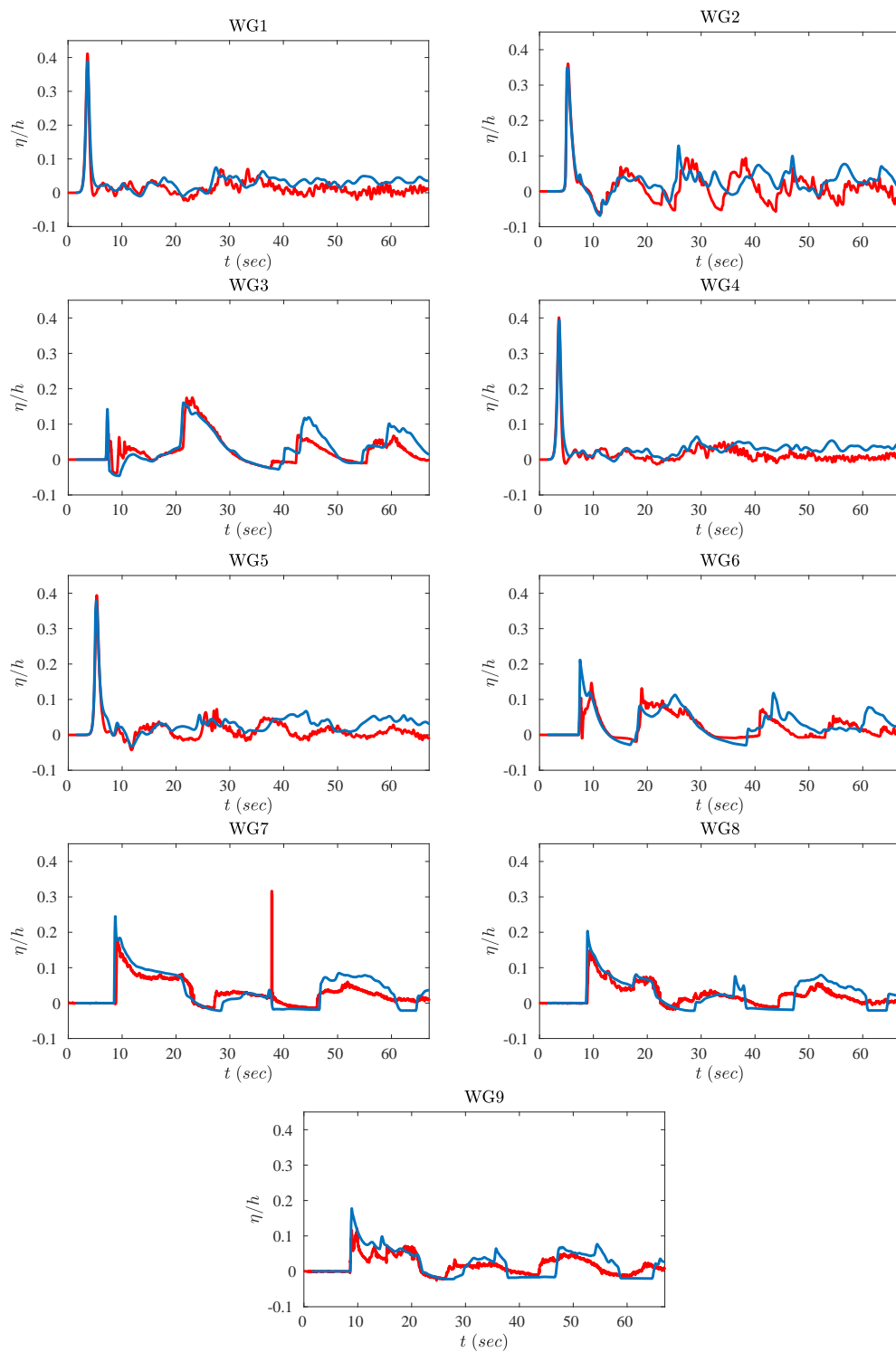


Figure 24: Solitary wave propagation over a three dimensional reef: computed time series of the free surface elevation on gauges positions

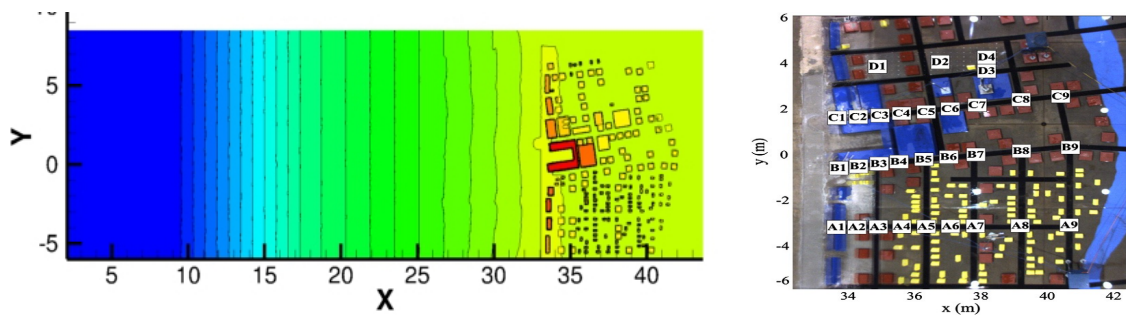


Figure 25: Left: Contour lines of the topography. Right: Measurement locations. Picture taken from [55]

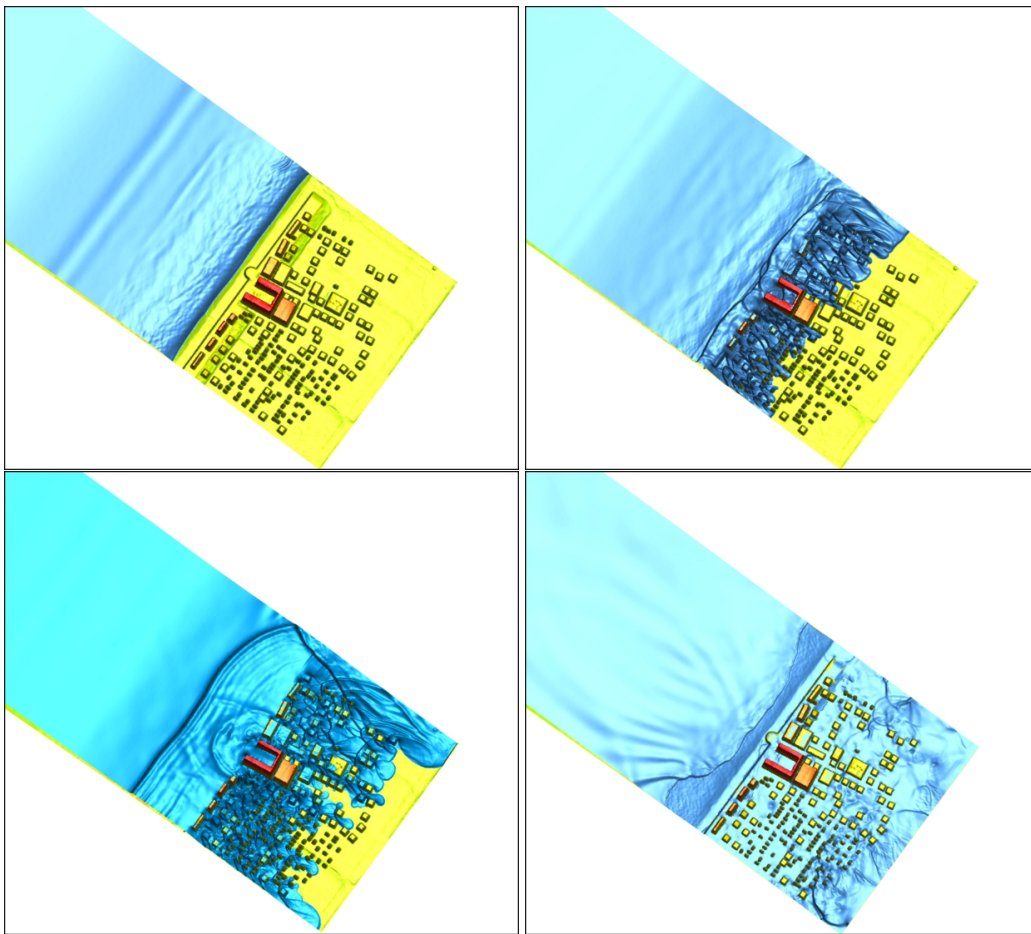


Figure 26: 3D view snapshots of the evolution of the wave after 14sec.

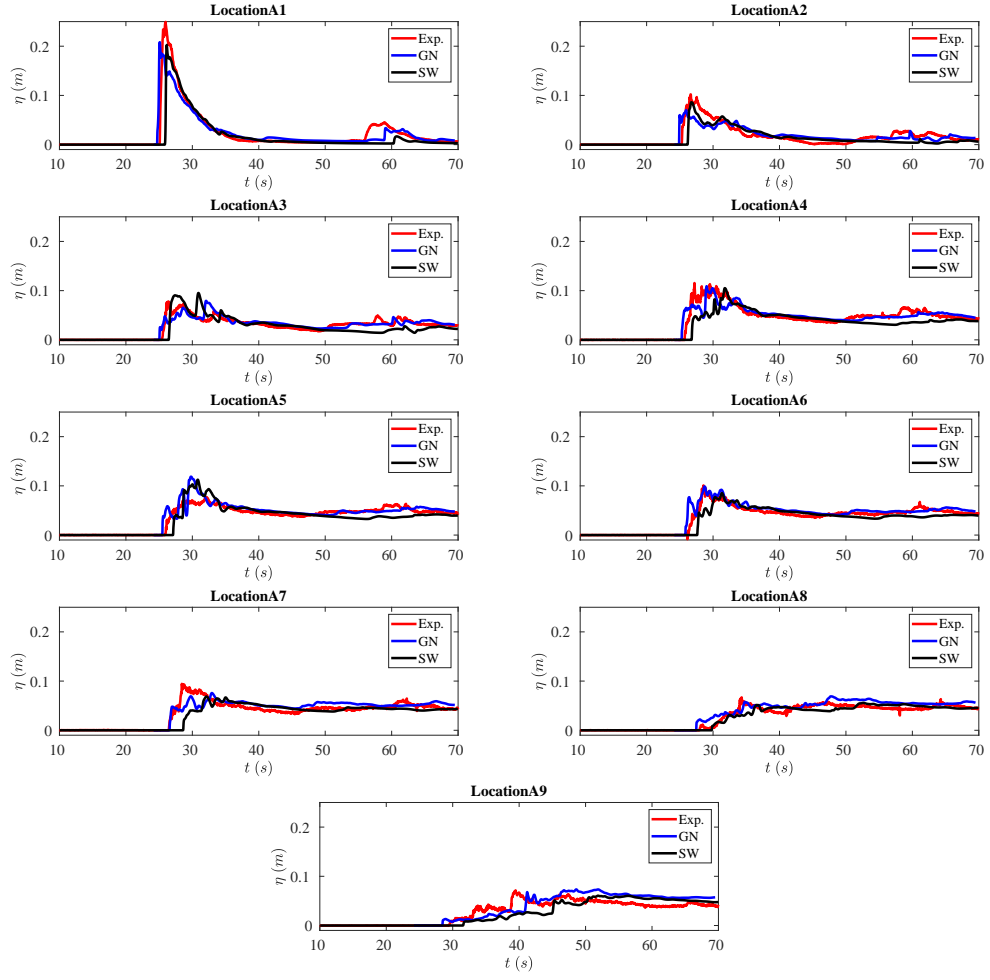


Figure 27: Free surface elevation measured in locations A

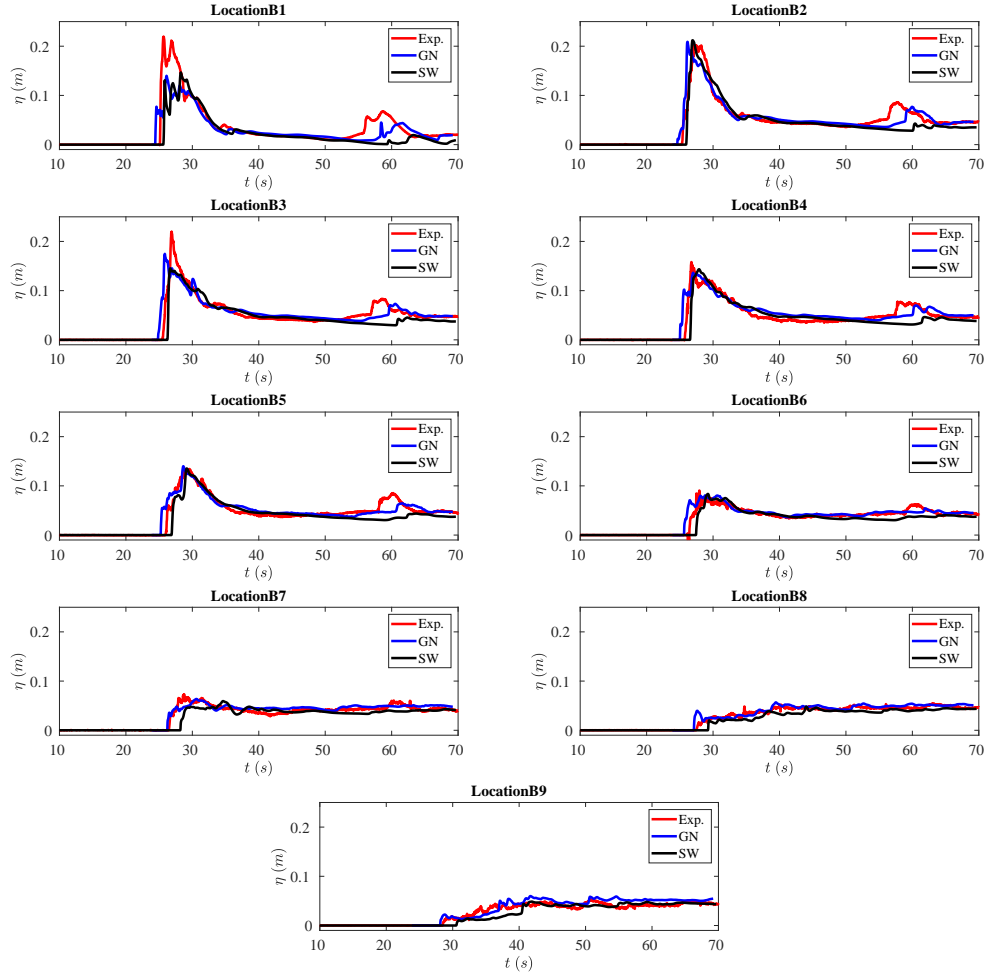


Figure 28: Free surface elevation measured in locations B

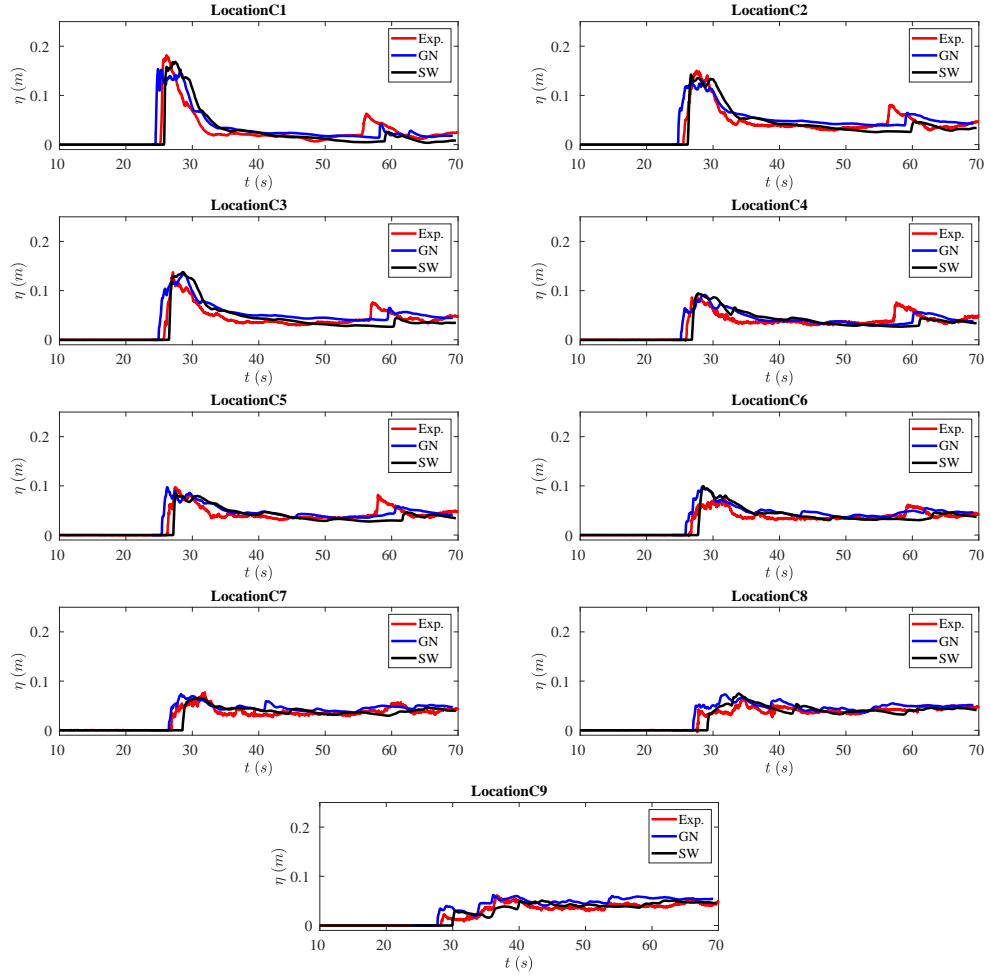


Figure 29: Free surface elevation measured in locations C

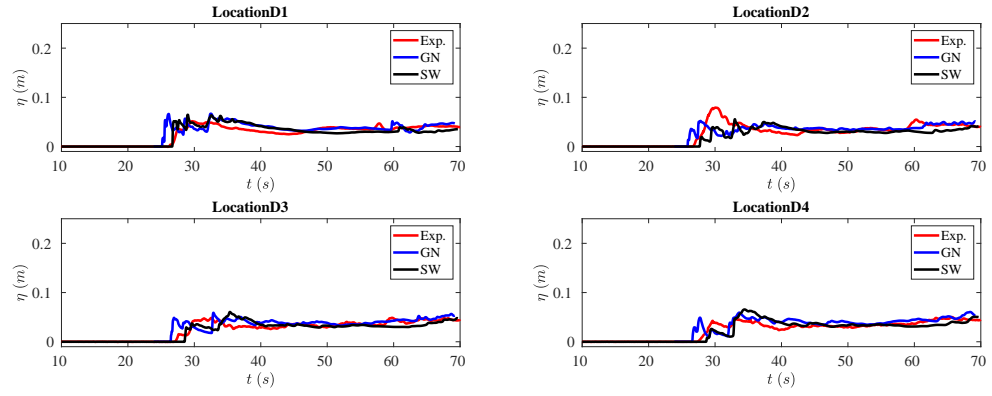


Figure 30: Free surface elevation measured in locations D

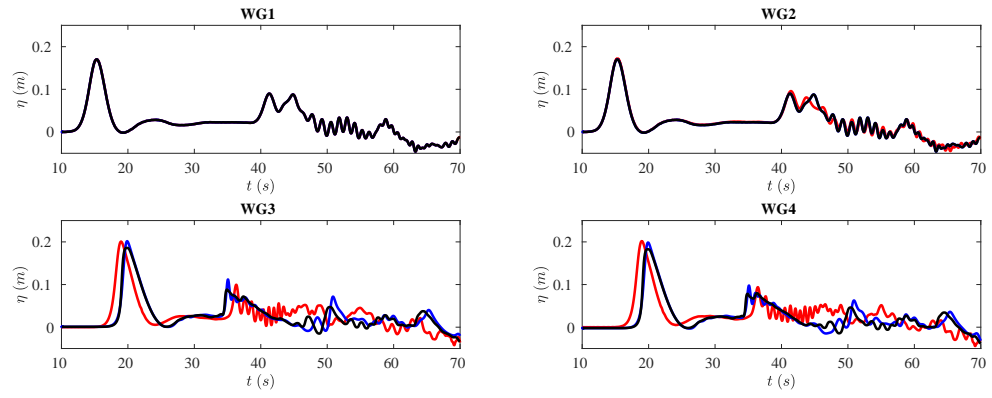


Figure 31: Free surface elevation measured in wave gauges 1-4

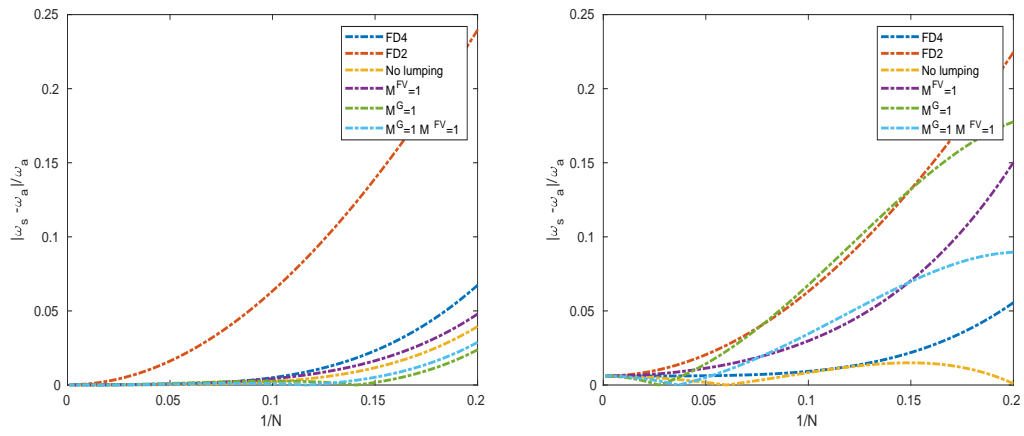


Figure B.32: Dispersion error for FV with respect to ω_{airy} . $kh = 0.5$ and 2.5 : impact of lumping strategy