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# Spatial and Seasonal Distributions of Frontal Activity over the Continental Shelf in the Bay of Biscay Focus on Density Fronts in Winter

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**Ifremer** 



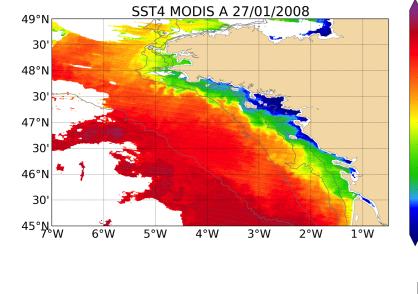




#### **AIMS**

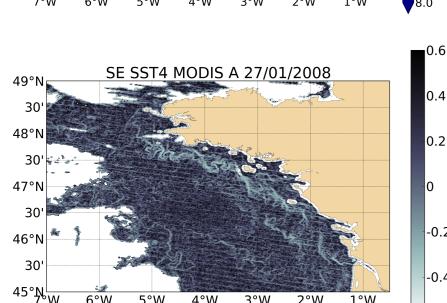
- Describing spatial and seasonal distributions of frontal activity over the continental shelf in the Bay of Biscay,
- Investigating the wintertime mid-shelf fronts in vicinity of the river plumes,
- Exploring the physical dynamics and the existence of baroclinic instabilities in such fronts from a scale decomposition of the vertical buoyancy flux.

#### FRONTAL ACTIVITY - OBSERVED VIA SATELLITE



**Product:** MODIS Aqua & Terra, Level-2, nighttime sea surface temperature (SST)

**Resolution:**  $\sim 1$  km, daily **Period:** 2003 - 2013



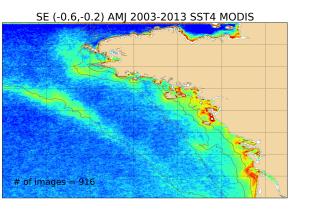
Singularity Exponent (SE) A diagnostic for front detection: A measure of the degree of regularity or irregularity of a function [1].

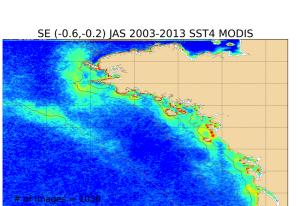
 $-SE \rightarrow stronger frontal activity$  $+SE \rightarrow$  weak frontal activity Frontal pixel:  $-0.2 \ge SE \ge -0.6$ 

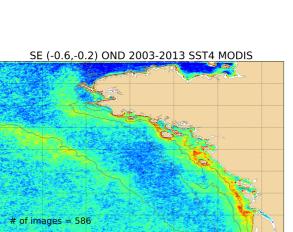
Figure 1: MODIS SST and SE field on 27/01/2008.

# Front occurence frequency:

# of times frontal pixel  $\times$  100 # of times cloud free







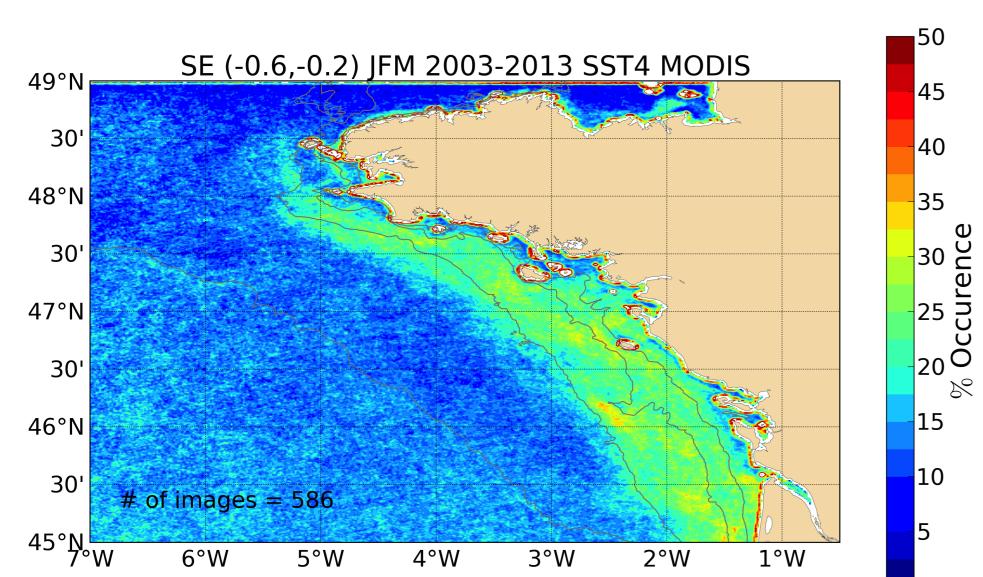


Figure 2: Front occurrence frequency in spring, summer, autumn, and winter.

# Seasonal & spatial variabilities of frontal activity

- Previously well-known tidal fronts and internal tidal wave induced shelf break front during thermally stratified months.

# Winter maximum over the shelf:

- From mid-autumn to mid-spring, along a band between 30 100 m isobaths,
- ▶ Density fronts at northward propagating river plumes confined along the coast,
- Sustained by the downwelling-favorable winds and surface cooling,
- $\triangleright$  Average cross-front gradient of 2  $^{\circ}$ C km $^{-1}$ ,
- ► Locally increased at the offshore edges of islands.

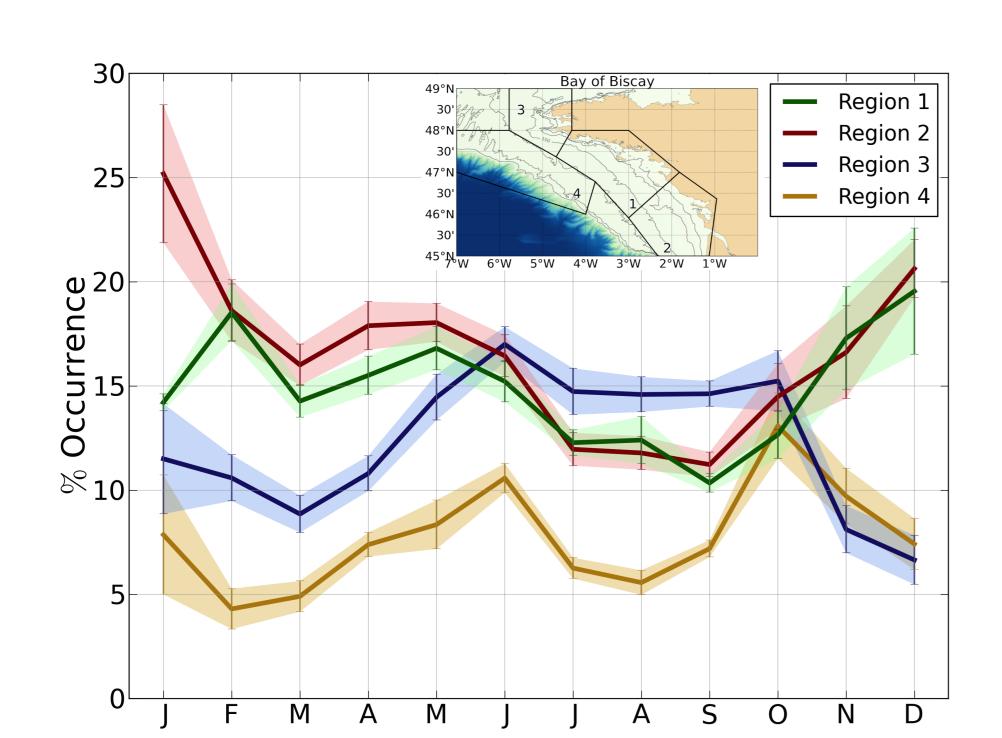
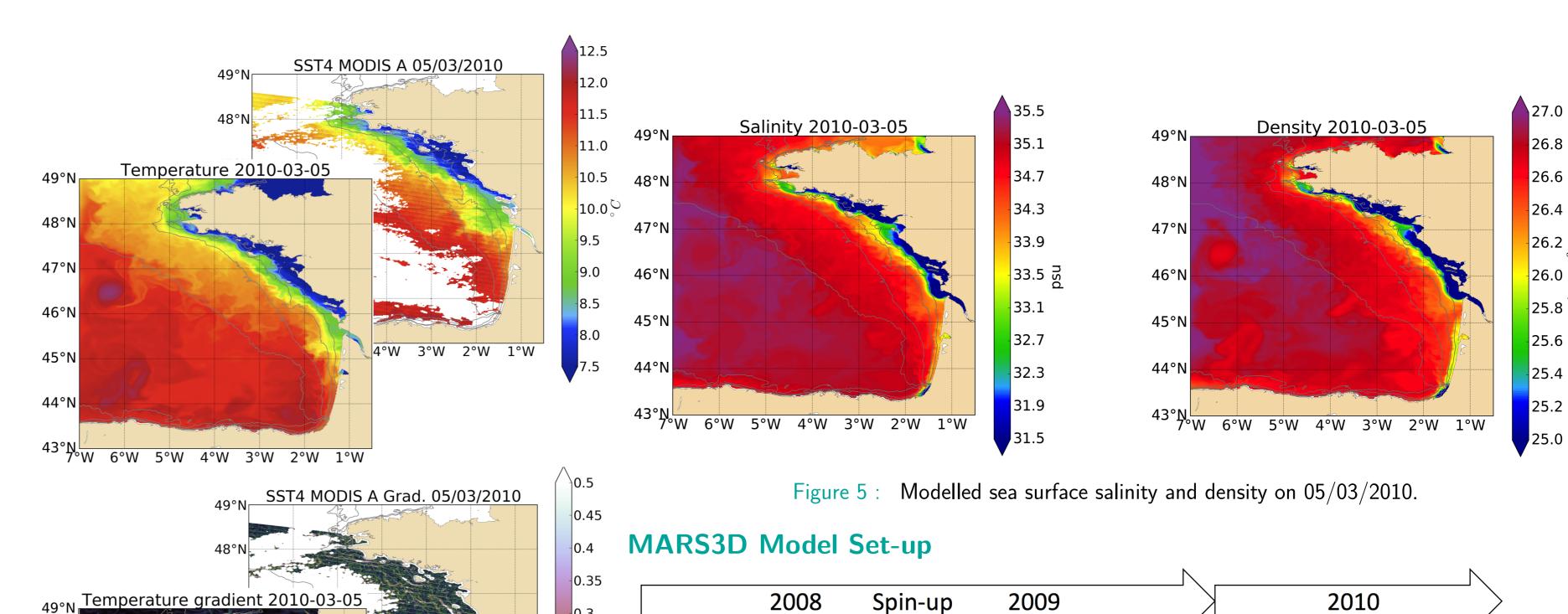


Figure 3: Monthly averaged front occurrence frequency in each of the regions defined. Error bars represent  $\sigma/n^{1/2}$  centered around the average,  $\sigma =$  standard deviation, and n =number of independent images.

#### FRONTAL ACTIVITY - NUMERICAL MODELLING



43°N 6°W 5°W 4°W 3°W 2°W 1°W Figure 4: Modelled & observed SST and SST gradient on 05/03/2010.

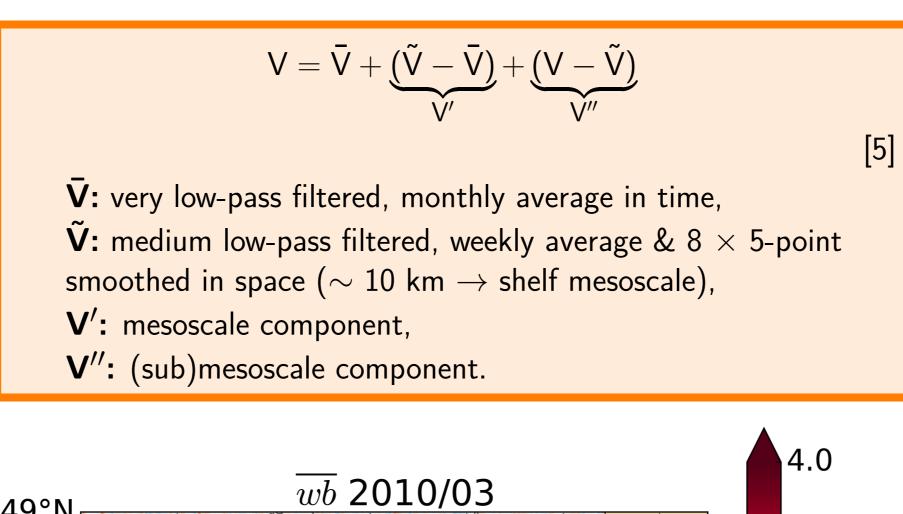
- ▶ Configuration adapted for the Bay of biscAy and the English CHannel (BACH) [2],
- $\Delta x = \Delta y = 1$  km, 40  $\sigma$ -layers,
- ▶ Atmospheric forcings: ARPEGE HR Meteorological model output,  $0.1^{\circ}$  [3],
- ▶ Rivers: CDOCO (Data Center for French Coastal Operational Oceanography) dataset, daily,
- ▶ Initial & open boundary cond.: ORCA12-MJM88 global simulation,  $1/12^{\circ}$  [4],
- ▶ Runs on Bull super-computer OCCIGEN in CINES (https://www.cines.fr).

#### **SCALE DECOMPOSITION**

45°N

49°N

48°N



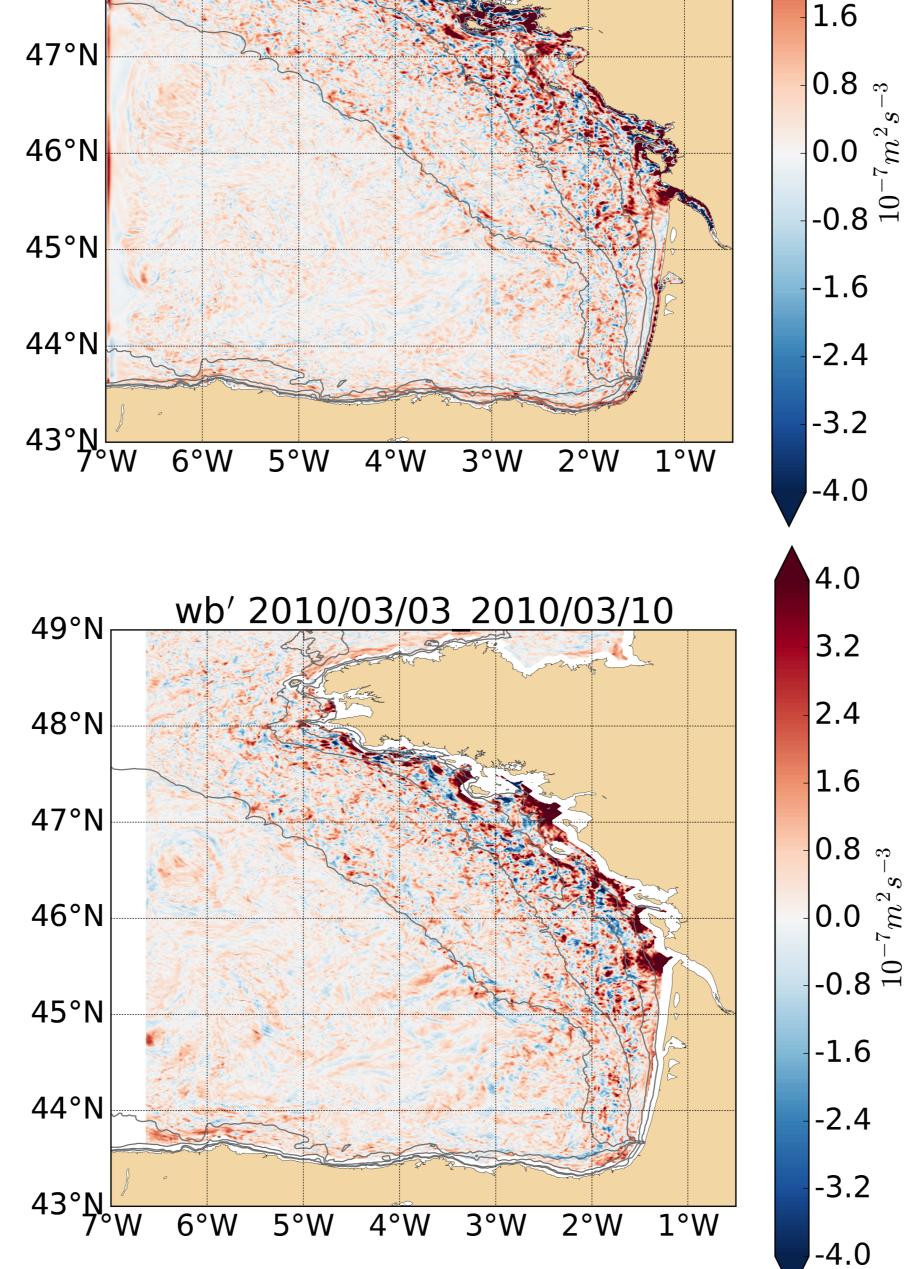


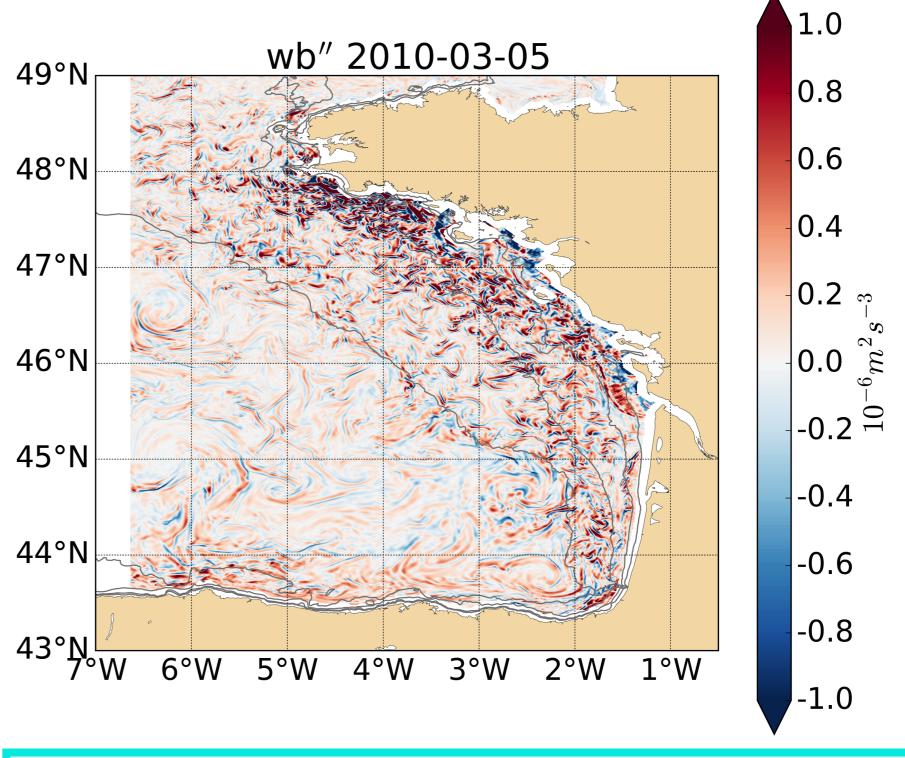
Figure 6: Modelled vertical buoyancy flux, its mean, meso-, and (sub)mesoscale components on 05/03/2010.

# wb 2010-03-05 2.4 48°N 1.6 47°N 0.8 46°N 0.0 -0.8 <sup>2</sup> 45°N -2.4 -3.2 -4.0

# wb: Vertical buoyancy flux

- Available potential to kinetic energy conversion at density fronts,

Baroclinic instabilities in vicinity of the river plumes



# **CONCLUSIONS:**

- ▶ Plume fronts, driven by the increased fresh water input, wind stress, and air-sea fluxes, dominate the frontal activity over the shelf in winter,
- $\triangleright$  (sub)mesoscale component, w"b", prevails in vicinity of these fronts,
- Instabilities occur at the plume fronts and at a succession of weaker fronts inside the plumes,
- $\triangleright$  Spatial variability of w"b" along the plumes: topography, tides, and atmospheric forcings are potential influences.

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