

Application of optimal transport to data assimilation

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EXAMPLE OF OCEANOGRAPHIC DATA ASSIMILATION

Given a state \mathbf{x} whose evolution is simulated through a model \mathcal{M} , data assimilation is an inverse problem aiming at reconstructing an initial condition \mathbf{x}_0 of \mathbf{x} given only several partial observations $(\mathbf{y}_i^o)_i$ of \mathbf{x} . For example, the state $\mathbf{x}(t, x, y, z)$ of the ocean gathers :

- Temperature T(t, x, y, z)
- Velocities $\mathbf{u}(t, x, y, z)$
- Salinity S(t, x, y, z)
- Water height h(t, x, y)

The question is then: What is the complete state of the ocean at t = 0, $\mathbf{x}_0(x, y, z)$?



The inner product:

For convergence reason, instead of using the \mathcal{L}_2 inner product, we use the \mathcal{W}_2 inner product defined in \mathbf{x}_0 by

$$(\eta, \eta') = \int_{\Omega} \mathbf{x}_0 \nabla \Phi \cdot \nabla \Phi',$$

with $\Phi, \Phi' \text{ s.t.} \qquad \eta = -\operatorname{div}(\mathbf{x}_0 \nabla \Phi), \qquad \eta' = -\operatorname{div}(\mathbf{x}_0 \nabla \Phi)$

Finally, the gradient of \mathcal{J}_W w.r.t. \mathcal{W}_2 inner product is, $(L^* \text{ is the adjoint model})$,

grad
$$\mathcal{J}_W(\mathbf{x}_0) = -\operatorname{div}\left(\mathbf{x}_0 \nabla \left(\sum_i L^*[t_i, \mathbf{x}_0].\psi_i\right)\right)$$

ASSIMILATION OF NON-PROBABILITY MEASURE VARIABLES

In the case where $\mathcal{H}_i(\mathbf{x}_0)$ and \mathbf{y}_0 are probability measures, but not \mathbf{x}_0 , it is not possible to write the background term \mathcal{J}^b with \mathcal{W}_2 . For example, for the Shallow-water system

In variational data assimilation this is done by minimizing a cost function \mathcal{J} defined as

$$\mathcal{J}(\mathbf{x}_{0}) = \underbrace{\frac{1}{2} \sum_{i} d\left(\mathcal{H}_{i}(\mathbf{x}_{0}), \mathbf{y}_{i}^{o}\right)^{2}}_{\approx d\left(\mathbf{x}_{0}, \mathbf{x}_{0}^{t}\right)^{2}} + \frac{\gamma}{2} d\left(\mathbf{x}_{0}, \mathbf{x}_{0}^{b}\right)^{2}, \qquad (1)$$

It is common for the distance d to be a weighted \mathcal{L}_2 distance. Yet, with this \mathcal{L}_2 distance, variational data assimilation cannot well cope with displacement errors. By choosing distances d that take into account the datas more in its entirety, we hope get a more realistic variational data assimilation. That is why we are interested in the Wasserstein distance $d = \mathcal{W}_2$.

ADVANTAGES OF OPTIMAL TRANSPORTATION IN CASE OF DISPLACEMENT ERROR

Knowing two densities $\rho_0(x)$ and $\rho_1(x)$, the **Wasserstein distance** $\mathcal{W}_2(\rho_0, \rho_1)$ is defined as

$$\mathcal{W}_{2}^{2}(\rho_{0},\rho_{1}) := \inf_{ \substack{ (\rho(t,x), v(t,x)) \\ \partial_{t}\rho + \operatorname{div}(\rho v) = 0 \\ \rho(0,x) = \rho_{0}(x), \rho(1,x) = \rho_{1}(x) } \int_{[0,1]\times\Omega} \rho(t,x) |v(t,x)|^{2} \, \mathrm{d}t \mathrm{d}x.$$

Example of interpolation using Wasserstein and \mathcal{L}_2 distances. (Interpolation in the sense of minimizing $d(\rho, \rho_0)^2 + d(\rho, \rho_1)^2$.

$$ho_0$$
 ho_1

$$\begin{cases} \partial_t h + \operatorname{div}(hu) = 0\\ \partial_t u + (u \cdot \nabla)u + g\nabla h = \mu \Delta u \end{cases}$$
(2)

where (h_0, u_0) are to be assimilated using observations of h only. The cost function writes

$$\mathcal{J}(h_0, u_0) = \frac{1}{2} \sum_i \mathcal{W}_2^2 \Big(\mathcal{H}_i(h_0, \boldsymbol{u}_0), h_i^o \Big) + \gamma \mathcal{J}^b(h_0, u_0).$$

As it is impossible to write $\mathcal{W}_2^2(u_0, u_0^b)$, we rather use the following background term

$$J^{b}(h_{0}, u_{0}) = \inf_{\substack{\partial_{t}\rho + \operatorname{div}(\rho v) = 0 \\ \partial_{t}u + \operatorname{div}(\rho w) = 0 \\ \rho(0, x) = h_{0}, \rho(1, x) = h_{0}^{b} \\ u(0, x) = u_{0}, u(1, x) = u_{0}^{b}} \int_{[0,1]\times\Omega} (|v|^{2} + |w|^{2}) \rho \, \mathrm{d}t \mathrm{d}x,$$
(3)

Using this, the interpolation of two shifted (h_0, u_0) and (h_1, u_1) is in-between:





For Wasserstein distance to be defined, one needs $\rho_0 \ge 0$, $\rho_1 \ge 0$ and $\int_{\Omega} \rho_0 = \int_{\Omega} \rho_1$.

GRADIENT DESCENT WITH THE WASSERSTEIN DISTANCE

Assuming $\mathbf{x}_0, \, \mathcal{H}_i(\mathbf{x}_0)$ and \mathbf{y}_i^o are probability measures, the cost function \mathcal{J}_W is

$$\mathcal{J}_W(\mathbf{x}_0) = \frac{1}{2} \sum_i \mathcal{W}_2^2 \Big(\mathcal{H}_i(\mathbf{x}_0), \mathbf{y}_i^o \Big)$$

The gradient descent algorithm for minimizing \mathcal{J}_W consists in using the following iterative algorithm

$$n \perp 1$$
 n $1 \neq (n)$

The inner product to be chosen for computing the gradient will be

 $\left(\begin{pmatrix}\eta\\v\end{pmatrix},\begin{pmatrix}\eta'\\v'\end{pmatrix}\right) = \int_{\Omega} h_0 \nabla \Phi \cdot \nabla \Phi' + \int_{\Omega} h_0 \nabla \Psi \cdot \nabla \Psi'$

with Φ , Φ' , Ψ , Ψ' such that

$$\eta = -\operatorname{div}(h_0 \nabla \Phi), \qquad \eta' = -\operatorname{div}(h_0 \nabla \Phi')$$
$$v = -\operatorname{div}(h_0 \nabla \Psi), \qquad v' = -\operatorname{div}(h_0 \nabla \Psi').$$

DIFFICULTIES OF USING THE WASSERSTEIN DISTANCE

• The Wasserstein distance is only defined for probability measures.

• When $\rho_0, \rho_1 \approx 1$, the \mathcal{W}_2 interpolation looks like \mathcal{L}_2 interpolation...

• When $\mathcal{J}(h_0^n) \to \min_{h_0} \mathcal{J}(h_0)$, one only has $h_0^n \rightharpoonup \arg\min_{h_0} \mathcal{J}(h_0)$.

• The computing time is larger for \mathcal{W}_2 than for \mathcal{L}_2 [Peyré, Papadakis, Oudet, 2013].

RESULTS AND PROPECTS

With some tests on the assimilation of (2), we compare the error $||h_0 - h_0^t||_2^2$ by using $d = \mathcal{L}_2$ and $d = \mathcal{W}_2$ in the cost function. The behaviors seem correct.

 $\mathbf{x}_0^{n+1} = \mathbf{x}_0^n - \alpha \operatorname{grad} \mathcal{J}_W(\mathbf{x}_0^n)$

so that $\mathcal{J}_W(\mathbf{x}_0^{n+1}) < \mathcal{J}_W(\mathbf{x}_0^n)$. To get the gradient, the **differential** must be computed and also an **inner product**. Indeed, grad $\mathcal{J}(\mathbf{x}_0)$ is such that for all η , $(\eta, \operatorname{grad} \mathcal{J}(\mathbf{x}_0)) = D\mathcal{J}[\mathbf{x}_0].\eta.$

The differential:

Let's differentiate \mathcal{J}_W by computing $\mathcal{J}_W(\mathbf{x}_0 + \epsilon \eta)$:

• $\mathcal{H}_i(\mathbf{x}_0 + \epsilon \eta) = \mathcal{H}_i(\mathbf{x}_0) + \epsilon L[t_i, \mathbf{x}_0].\eta$ with L the tangent model.

• $\frac{1}{2}\mathcal{W}_2^2(\mathbf{y}+\epsilon\mu,\mathbf{y}') = \epsilon\langle\mu,\psi\rangle_2$ with ψ the **Kantorovich potential** of optimal transportation between \mathbf{y} and \mathbf{y}' .

Then,

$$\mathcal{D}\mathcal{J}_W[\mathbf{x}_0].\eta = \sum_i \left\langle L[t_i, \mathbf{x}_0].\eta , \psi_i \right\rangle_2$$

with ψ_i the Kantorovich potential of the optimal transportation between $\mathcal{H}_i(\mathbf{x}_0)$ and \mathbf{y}_i^o .



The background term (3) is still to be implemented.

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